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中国西部新生代沉积盆地演化

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摘要: 新生代期间中国西部发生了多次强烈的构造运动, 经历了复杂的构造—地貌演化历史。地质构造背景和地球动力学过程则控制了中国西部大陆新生代期间的构造—地貌演化。盆—山系统是中国西部新生代构造的基本格局, 盆—岭体系是中国西部新生代的主要地貌单元。根据盆地的几何学、动力学与构造演化特征, 中国西部新生代盆地可以划分为压陷盆地、断陷盆地、走滑拉分盆地以及残留海—前陆盆地 4 类。这些新生代封闭盆地均被造山带所围限, 而盆地与山脉之间由挤压型活动断裂(逆冲断层和走滑断层)所分割。新生代以来印度板块与欧亚板块的碰撞以及其后印度板块的向北俯冲挤压, 对中国西部新生代沉积盆地的发育和演化产生了重大影响。中国西部新生代盆地构造岩相古地理演化与板块运动和构造隆升之间存在明显的耦合。

关键词: 中国西部; 沉积; 构造; 盆地演化; 新生代。

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Evolution of Cenozoic Sedimentary Basins in Western China

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Abstract: During the Cenozoic, strong multi-stage tectonic events occurred in western China, which experienced a complex tectonic-geomorphic evolution history. Tectonic setting and geodynamic process determined the tectonic-geomorphic evolutions in the continent of western China. Sedimentary basin-orogen belt systems are the dominant basic structural pattern in western China during the Cenozoic. Basin-range pattern is the major geomorphic unit in western China. According to the geometric, dynamic features of tectonic evolution of basins, the Cenozoic sedimentary basins in western China fall into four tectonic categories: contractional basin, faulted basin, strike-slip pull-apart basin and remnant sea-foreland basin. These closed geomorphic basins in western China were surrounded by orogenic belts, which are divided by active compressive faults (thrust fault and strike-slip fault). Both the Indian-Eurasian plate collision and the intracontinental subduction play significant role in the formation and evolution of sedimentary basins in western China. There is a close relationship among the tectonic lithofacies paleogeographic evolution of sedimentary basins, plate movement and tectonic uplift.

Key words: western China; sedimentology; tectonics; basin evolution; Cenozoic.

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0 引言

中国大陆位于欧亚板块东南部、印度板块北缘,东部与太平洋板块以沟—弧—盆体系相隔,其地质演化历史受古亚洲构造域、特提斯构造域、滨太平洋构造域不同阶段构造演化的影响,具有复杂的构造演化阶段和过程,是基于不同时期的众多陆块和缝合带拼合而成(黄汲清等,1977;潘桂堂等,1997,2009;刘光鼎,2007;Wan,2010;Zheng *et al.*,2012)。中国西部地区地貌格局总体表现为,在较为稳定的陆块或地块基底之上发育一系列被造山带所环绕的封闭盆地,即造山带和盆地呈镶嵌分布(Carroll *et al.*,2010)。小地块拼合、多旋回运动和陆内构造活动强烈等控制了西部盆—山系统的形成和演化(刘树根等,2003)。新生代以来,中国西部大陆造山带和盆地之间的耦合过程与印度板块、欧亚板块碰撞引起的大陆内缩短挠曲和走滑作用密切相关。通过对中国西部新生代沉积盆地的构造背景、演化特征及沉积序列的分析,笔者更加深入地了解了与中国西部新生代区域构造演化相关联的盆—山耦合历史。本文对沉积盆地类型的划分与鉴别特征采纳了张克信等(2007,2014)的划分与论述,在此不再重述。

1 中国西部前新生代大地构造背景

在中国的造山带和沉积盆地的发展中,印支运动期是一个关键的转折期,中国大陆主体在印支运动中拼合完成,但陆内构造变形剧烈、广泛且持久(任纪舜,1984)。不同时期特提斯古大洋板块的扩张增生与俯冲消减对中国西部前新生代区域构造演化和古地理环境变迁具有重要的控制作用(潘裕生等,1998;Golonka,2007)。三叠纪末期,古特提斯洋盆的闭合导致了陆块的碰撞、拼合及增生,形成南欧亚大陆古特提斯碰撞造山系(许志琴等,2012)。中国西部大型走滑断裂包括南祁连断裂带、东昆仑断裂带和阿尔金断裂带,均形成于印支期(李海兵等,2001;许志琴等,2011)。印支期碰撞造山过程改变了亚洲大陆的古构造格局,中侏罗世之后中国大陆主体已处于陆地环境,进入到陆内造山和陆内变形的演化时期(王鸿祯,1985;Dong *et al.*,2007),因此本文探讨印支期以来的大地构造背景对理清中国西部沉积盆地的发展和演化具有重要的意义。

其中,塔里木铁克里克、柯坪等地中二叠世进入周缘前陆盆地演化阶段。中生代初塔里木陆块进入陆内盆地演化阶段,总体以挤压背景为主,形成一系列压陷盆地。西昆仑三叠纪发育弧后—弧前陆盆地,侏罗—白垩世发育断陷盆地。

晚二叠世—三叠纪,由于古特提斯大洋板块的俯冲消减,东昆仑南部形成海西—印支期中酸性岩浆侵入、褶皱变形、韧性剪切和浅变质作用(吴珍汉等,2009)。侏罗纪—白垩纪为缓慢抬升及小型陆相盆地形成阶段。

柴达木在晚二叠世—三叠纪由先前的弧后裂陷转为弧后挤压。晚三叠世由于受鄂拉山弧后陆缘裂陷影响,柴达木和柴北缘地区发育了陆内裂陷盆地。侏罗纪—白垩纪柴达木经历了一个不完整的拉张挤压旋回,其中早、中侏罗世表现出明显拉张特征,在柴达木盆地北缘发育一系列断陷盆地;而晚侏罗世—白垩世则具挤压特征,以发育压陷盆地为特征(汤良杰等,2000)。

随着三叠纪末期印支早幕挤压构造运动受松潘—甘孜地块南北向的挤压和龙门山向东构造的逆冲推覆,川西地区在中三叠世脱离海相沉积并进入周缘前陆盆地演化阶段(Li *et al.*,2003;Weislogel *et al.*,2010)。早—中侏罗世川西以坳陷盆地为主,晚白垩世以来在坳陷盆地的基础上发育了压陷盆地。

青藏地区晚三叠世—中侏罗世是特提斯洋的被动陆缘和陆内坳陷发育阶段,印支晚期古特提斯洋盆的闭合导致了青藏地区的最大海退。晚侏罗世(170 Ma)班公湖—怒江洋双向俯冲,早白垩世(约99 Ma)闭合完成,拉萨地块与羌塘地块碰撞拼合;以雅鲁藏布江缝合带为洋壳残迹的大洋板块自中侏罗世开始向北俯冲于羌塘板块之下,65 Ma 印度大陆与亚洲大陆南缘初始碰撞(莫宣学等,2009)。印度与欧亚碰撞不是在某个时间点完成的,其碰撞起始时间约65 Ma,陆—陆碰撞全面完成时间约45 Ma,碰撞方式(新特提斯洋关闭)为由东向西迁移(Zhang *et al.*,2013)。

2 中国西部新生代沉积盆地分布特征

本文共划分出中国西部新生代盆地及盆地群共87个(图1)。对于盆地类型的划分,本文在前人研究的基础上(刘和甫,1993,2001;张克信等,2007;Zhang *et al.*,2013),将中国西部新生代盆地划分为压陷—坳陷盆地、断陷(裂陷)盆地、走滑—拉分盆地

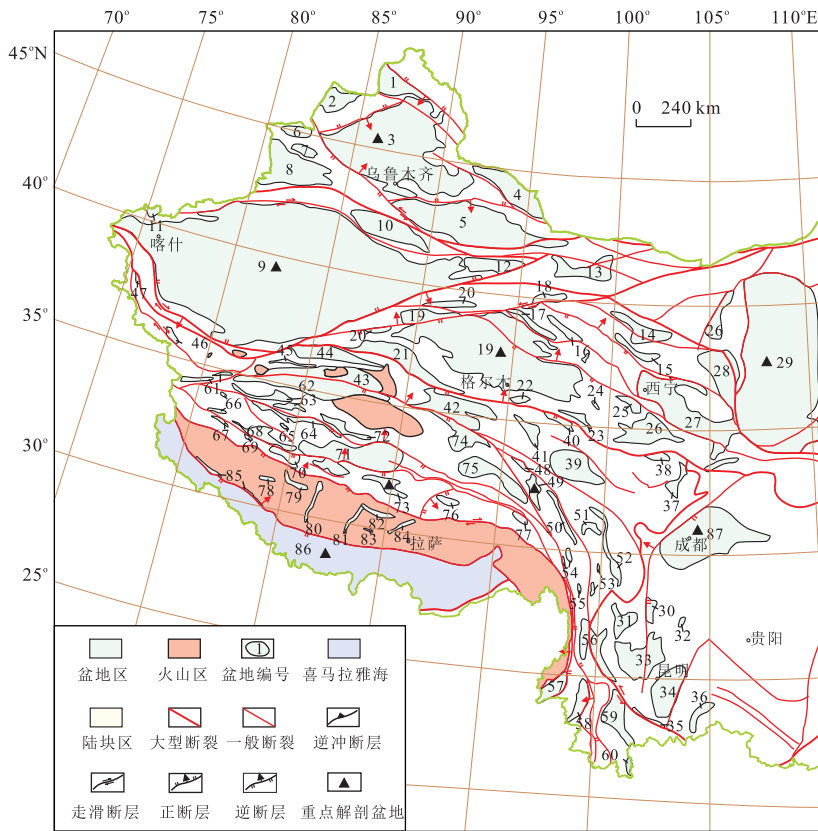


图 1 中国西部新生代沉积盆地分布

Fig.1 Distribution of Cenozoic sedimentary basins in western China

- 1.布尔津断陷盆地;2.塔城坳陷盆地;3.准噶尔压陷盆地;4.三塘湖压陷盆地;5.吐哈压陷盆地;6.博乐坳陷盆地;7.博罗科努凹陷盆地;8.伊宁压陷盆地;9.塔里木压陷盆地;10.焉耆坳陷盆地;11.吐尔尕特坳陷盆地;12.北山压陷盆地;13.包尔乌拉凹陷盆地;14.阿拉善右旗压陷盆地;15.张掖—武威压陷盆地;16.疏勒南山压陷盆地;17.党河南山压陷盆地;18.玉门西走滑拉分压陷盆地;19.柴达木压陷盆地;20.瓦石峡南压陷盆地;21.阿奇克库勒坳陷盆地;22.秀河压陷盆地;23.冬给措纳湖坳陷盆地;24.兴海坳陷盆地;25.共和坳陷盆地;26.同仁坳陷盆地;27.西宁—兰州压陷盆地;28.宁夏压陷盆地;29.鄂尔多斯断陷盆地;30.西昌断陷盆地;31.木里坳陷坳陷盆地;32.昭通坳陷盆地;33.元谋坳陷盆地;34.昆明断陷—坳陷盆地;35.元江压陷盆地;36.明江上游坳陷盆地;37.百色断陷—坳陷盆地;38.松潘坳陷盆地;39.玛曲压陷盆地;40.石渠坳陷盆地;41.曲麻莱坳陷盆地;42.五道梁压陷盆地群;43.五道梁压陷盆地群;44.木夜塔格断陷盆地;45.那底岗日断陷盆地;46.喀喇崑崙山山口压陷—拉分盆地;47.塔尔库尔干压陷盆地;48.治多压陷盆地;49.囊谦拉分盆地;50.生达—巴塘走滑拉分盆地群;51.白玉坳陷盆地;52.甘孜—理塘走滑拉分盆地群;53.乡城—中甸走滑拉分盆地群;54.芒康走滑拉分盆地;55.德钦压陷盆地;56.丽江走滑拉分盆地;57.腾冲断陷盆地;58.昌宁断陷盆地;59.思茅压陷—拉分盆地;60.勐腊走滑拉分盆地;61.松西—鄂扎错压陷盆地;62.拜惹布错—多格错仁强错断陷盆地;63.温泉湖—查多岗日压陷盆地;64.鱼鳞山—玛依岗日压陷盆地;65.丁固坳陷盆地;66.鲁玛江冬错—扎普坳陷盆地;67.丁则压陷盆地;68.强堆—日玛压陷盆地;69.擦咪坳陷盆地;70.改则坳陷盆地;71.尼玛—双湖压陷盆地;72.玛尔果茶卡压陷盆地;73.伦坡拉断陷—坳陷盆地;74.唐古拉坳陷盆地;75.杂多坳陷盆地;76.那曲坳陷盆地;77.丁青压陷盆地;78.隆格爾压陷—拉分盆地;79.措勤压陷—拉分盆地;80.当惹雍错断陷盆地;81.谢通门断陷盆地;82.申扎走滑拉分盆地;83.南木林断陷盆地;84.羊八井断陷盆地;85.冈仁波齐—罗波峰压陷盆地;86.喜马拉雅残留海;87.成都压陷盆地

和残留海—前陆盆地 4 大类。

压陷—坳陷盆地主要分布在中国西部青藏高原周边及高原内部,盆地的边界逆冲断层控制着盆地的发展.压陷盆地的基底通常为刚性较大的陆块、地块或褶皱基底,同时盆地的展布方向通常与造山带一致(潘桂堂等,1990),其包括塔里木盆地、准噶尔盆地、吐鲁番—哈密盆地、伊宁盆地等,以及前期的造山带复活后发育的柴达木盆地、可可西里盆地、张

掖—武威盆地等.断陷(裂陷)盆地主要发育在伸展构造背景下,控制盆地形成和发展的最基本构造为盆缘正断层.盆地演化通常经历了初始褶皱断陷(裂陷)期—断陷(裂陷)扩展沉降期—萎缩封闭期的断陷(裂陷)旋回.主要为东西向展布的伦坡拉盆地、那曲盆地、尼玛—双湖盆地、南北向的当惹雍错盆地、谢通门盆地、羊八井盆地等.走滑—拉分盆地主要沿区域性大断裂或沿特提斯洋陆转换形成的巨型结合

带分布(Wang *et al.*, 2011; 张克信等, 2013). 在中国西部主要分布于横断山地区, 均呈现 SN 向、NWW 向展布, 通常伴随着各类火山岩浆喷发和溢流(潘桂堂等, 1990). 主要为甘孜—理塘盆地、乡城—中甸盆地、芒康盆地、丽江盆地、申扎盆地等. 残留海—前陆盆地主要分布于西藏南部地区, 其形成与古—始新世期间印度大陆与亚洲大陆的碰撞以及新特斯提洋的闭合有关.

在分布格局方面, 盆—山系统是中国西部新生代构造的基本格局, 而盆—岭地貌是中国西部新生代的主要地貌单元.

3 中国西部新生代主要盆地构造—岩相古地理演化

笔者重点描述的中国西部新生代大型盆地主要有塔里木盆地、准噶尔盆地、鄂尔多斯盆地、柴达木盆地、羌塘盆地、囊谦盆地、成都盆地和喜马拉雅残留海等(图 1, 图 2). 以上盆地新生代期间充填序列和演化特征如下.

3.1 塔里木盆地

塔里木盆地是在克拉通陆壳基底上发育的大型复合叠合盆地. 新生代以来, 受欧亚板块和印度板块碰撞远程效应的影响, 天山、昆仑山、阿尔金山受挤压及扭张应力的影响不断上升和变形, 塔里木相对沉降, 经过复杂的过程形成现今大型压陷盆地格局. 由于塔中巴楚隆起的分隔, 塔里木盆地南北发育两套不同的沉积序列.

塔里木盆地南缘地区新生代地层十分发育, 由老到新分别为阿尔塔什组(E_{1a})、齐姆根组(E_{1-2q})、卡拉塔尔组(E_{2k})、乌拉根组(E_{2w})、巴什布拉克组(E_{3b})、克孜洛伊组(N_{1k})、安居安组(N_{1a})、帕卡布拉克组(N_{1p})、阿图什组(N_{2a})和西域组($N_2 Qp^1x$) (李云通, 1984; 蒋显庭, 1995; 贾承造等, 2004; Zhang *et al.*, 2010), 各组间均呈整合接触(巴什布拉克组和克孜洛伊组之间不是). 古新世—始新世沉积与下伏白垩纪海相地层为连续沉积, 为新特斯提残留海的一部分, 以半封闭的海湾滨浅海沉积为主(郝诒纯等, 2002)(图 3). 古新世初开始沉积的阿尔塔什组底部普遍发育白色巨厚层状石膏层, 为干旱炎热气候条件下发育的萨布哈相沉积. 之后发生大范围海侵, 呈现向上变深的海侵序列, 早期的封闭—半封闭的萨布哈—局限台地相沉积环境被打破, 随

之沉积的是大范围分布的碳酸盐岩开阔台地相, 以齐姆根组、卡拉塔尔组和乌拉根组的生屑灰岩、鲕粒灰岩、牡蛎介壳灰岩和灰绿色、暗红色泥岩为代表, 其中盛产有孔虫、介形虫和双壳类等海相动物群(张克信等, 2010, 2013). 始新世晚期叶城海湾内的内源生物—化学堆积逐步被陆源碎屑沉积所替代(唐天福等, 1992; 邵龙义等, 2006). 渐新世时期塔里木盆地西部西昆仑、喀喇昆仑和东部的阿尔金地区隆起, 同时西南部为浅海环境, 东南部为河湖环境, 两者被塔中隆起所隔(郝诒纯等, 2002).

新近纪印度板块进一步向北运动, 天山、西昆仑山加剧隆升和变形, 塔里木盆地也最终封闭形成现今格局. 塔里木盆地西南部在渐新世末期结束海相沉积, 进入大型内陆盆地演化阶段(图 3).

塔里木北部地区包括库车凹陷及塔北隆起, 毗邻天山褶皱系, 北缘是南天山构造带. 受塔中巴楚隆起的隔断, 塔北地区形成了另一套沉积序列(图 3), 由老到新分别为塔拉克组(E_{1t})、小库孜拜组(E_{2x})、苏维依组(E_{2-3s})、吉迪克组(N_{1j})、康村组(N_{1-2k})和库车组(N_{2k}) (腾志宏等, 1997; 邓秀芹等, 1998; 贾承造等, 2004; Huang *et al.*, 2006). 古近纪时期, 库车凹陷延续了白垩纪时期的沉积环境和气候, 发育红色碎屑岩系. 新近纪以来, 库车凹陷受喜马拉雅末期构造运动的显著影响, 形成天山山前大型逆冲褶皱系及一系列逆冲断层. 中新世库车拗陷北缘发生台阶状逆断层由北向南逆冲, 南缘伸展构造带的基底断层受到盆地北缘的逆冲作用再度复活, 使得库车盆地进入再生前陆盆地演化阶段(王清晨等, 2007).

3.2 准噶尔盆地

准噶尔盆地是自晚古生代至中生代以来在华力西期褶皱基底上发展演化而成的新生代陆相大型压陷沉积盆地(李丕龙等, 2010). 新生代受到欧亚板块和印度板块的碰撞远程效应, 盆地周缘山系抬升, 受北天山、阿尔泰山、博格达山构造带控制, 准噶尔盆地全面发育新生代沉积序列. 由于南部山系抬升强烈, 盆地向南倾斜, 北部地区早古近纪地层缺失. 在准噶尔盆地南部、北部分别发育两套不同的沉积序列.

古近纪时期, 准噶尔盆地南部整体表现为较稳定的沉积期, 为均衡沉降区, 以紫泥泉子组(E_{1-2z})和安集海河组(E_{2-3a})为代表的古近纪沉积在盆地分布相对均匀(图 2), 紫泥泉子组与白垩系沉积呈不整合接触. 新近纪以来准噶尔盆地发生强烈挤压和走滑变形, 奠定了盆地最终的构造格局. 新近纪

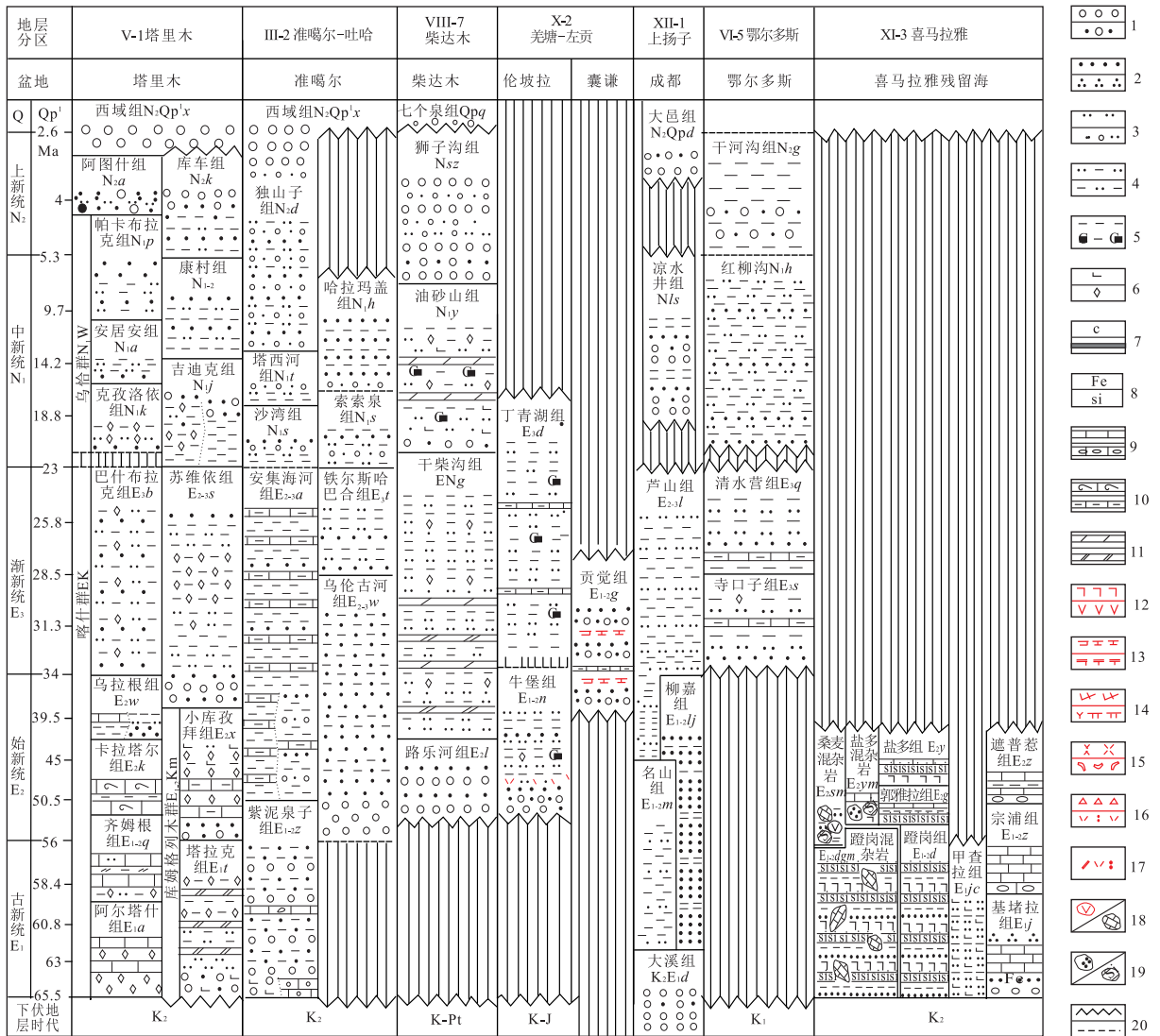


图 2 中国西部主要大型盆地新生代充填序列及其对比

Fig.2 Correlation and sedimentary sequence of the main Cenozoic basins in western China

1.砾岩/含砾砂岩;2.砂岩/石英砂岩;3.粉砂岩/含砾粉砂岩;4.泥质粉砂岩/粉砂质泥岩;5.泥岩/油页岩;6.钙质/石膏;7.炭质/煤;8.铁质/硅质;9.灰岩/灰质砾岩;10.介壳灰岩/含泥质灰岩;11.泥灰岩/泥质白云岩;12.玄武岩/安山岩;13.粗面岩/粗安岩;14.英安岩/白云榴榴岩;15.流纹岩/火山集块岩;16.火山角砾岩/凝灰岩;17.晶屑凝灰岩;18.火山岩岩块/灰岩岩块;19.砂岩岩块/硅质岩岩块;20.角度不整合/平行不整合.据以下参考文献:王开发等,1975;李云通,1984;蒋显庭,1995;青海省地质矿产局,1997;四川省地质矿产局,1997;腾志宏等,1997;邓秀芹等,1998;杨景林和沈一新,2000;赵政璋等,2001;王凤林等,2003;贾承造等,2004;曾宜君等,2004;李国彪等,2005;王成善等,2005;Huang *et al.*,2006;宋春晖,2006;Fang *et al.*,2007;Wang *et al.*,2007;张克信等,2007,2010,2013;Yin *et al.*,2008;Lu and Xiong,2009;李丕龙等,2010;周江羽等,2011

印度板块与亚洲板块碰撞远程效应使天山造山带再次活跃,天山地区发生快速地壳缩短和强烈的逆冲推覆,北天山快速大幅度隆升,并向盆地冲断产生巨大的构造负荷,在挤压地球动力学背景之下,盆地南缘再次挠曲下沉形成压陷盆地(再生前陆盆地)(Hendrix *et al.*,1994;方世虎等,2006).发育了一套以沙湾组(N₁t)、塔西河组(N₁s)和独山子组(N₂d)为代表的内陆河湖相沉积(蒋显庭,1995),沉积相和

厚度均表现为明显的南粗厚、北细薄的不对称性。
准噶尔盆地北部地层序列由老到新包括乌伦古河组(E₂₋₃w)、铁尔斯哈巴合组(E₃t)、索索泉组(N₁s)、哈拉玛盖组(N₁h)(杨景林和沈一新,2000),其新生代早期地层缺失。
3.3 柴达木盆地
柴达木盆地是我国西部大型含油气盆地之一,地处古亚洲构造域南缘附近,其南邻特提斯-喜马

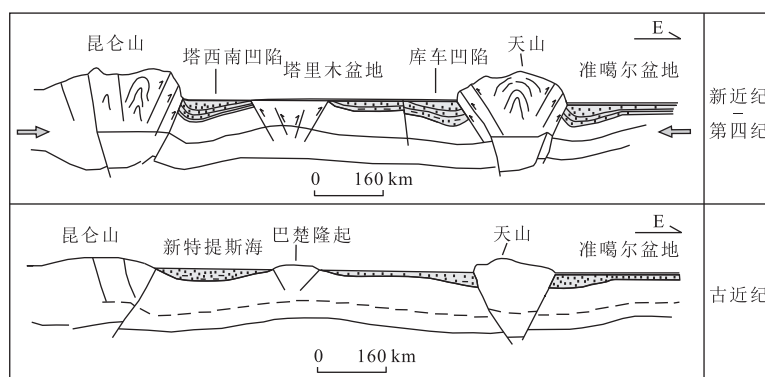


图 3 塔里木地区新生代地球动力学演化

Fig.3 Geodynamic evolution of the Tarim region in the Cenozoic

据以下参考文献: Hendrix *et al.*, 1994; Li, 1996; 杨树锋等, 2003; 卢华复等, 2005; 方世虎等, 2006; Sobel *et al.*, 2006; 王清晨等, 2007; 陈正乐等, 2008

拉雅构造域. 新生代以来, 柴达木盆地一直处于祁连山、阿尔金山、昆仑山三大体系相互作用形成的挤压一走滑的区域应力环境中, 其根本动力来自南侧印度板块向北的俯冲挤压(吕宝凤等, 2011). 柴达木盆地新生代岩石地层序列从老至新为路乐河组(E_2l)、干柴沟组(ENG)、油砂山组(N_1y)、狮子沟组(N_5z)和七个泉组(Qpq)(李云通, 1984; 青海省地质矿产局, 1997; 宋春晖, 2006; Lu and Xiong, 2009; Zhang *et al.*, 2010). 柴达木盆地新生代的沉积演化直接受盆地周缘三大主控断裂, 即阿尔金断裂、东昆仑断裂及南祁连断裂的控制, 盆—山耦合明显(Fang *et al.*, 2007; 陈宣华等, 2010; Yin *et al.*, 2008). 古新世, 柴达木盆地为无沉积的隆起区. 始新世早期柴达木盆地受印度—欧亚板块碰撞远程效应影响, 发生初始拗陷, 自约 53.5 Ma 开始接受与下伏白垩系地层不整合接触的路乐河组沉积, 沉积中心位于当时盆地西部阿尔金山前一带, 其他地区广泛分布洪泛和河流相红色碎屑沉积(赵加凡等, 2004). 始新世晚期开始, 盆地更大范围拗陷, 湖水面积扩大, 开始形成比路乐河组分布范围更广、几乎遍及全盆地的干柴沟组细碎屑沉积. 渐新世—中新世早期为柴达木盆地稳定沉降阶段, 沉积厚度巨大. 柴北缘发育一系列湖泊—三角洲相沉积, 大江沟剖面渐新统一早中新统地层古水流和物源资料显示南祁连造山带是其主要物源区(宋博文等, 2010; Song *et al.*, 2013a, 2013b). 中新世后期, 由于印度板块不断向北俯冲, 受青藏高原快速隆升的影响, 柴达木古湖盆开始萎缩, 地层中大量的石膏晶体和岩盐层表明盆地开始出现消亡的迹象. 上新世之后, 盆地沉积面积缩小, 接受粗碎屑的狮子沟组和七个泉组沉积.

物探和钻孔资料表明新生代以来盆地沉积中心具由西向东的迁移趋势(Wang *et al.*, 2006; 吕宝凤等, 2008), 始新世沉积中心位于盆地西南部的油砂山一带, 渐新世沉积中心开始向东迁移, 到达现代盆地的中部, 上新世盆地沉积中心位于一里坪一带, 上新世以来沉积中心继续向东南方向迁移(Metivier *et al.*, 1998), 至第四纪才基本与现代一致. 上述沉积中心的迁移表明阿尔金山至少在渐新世就已经开始了显著的成山作用, 并对柴达木盆地的沉积迁移起到了明显的控制作用.

3.4 伦坡拉盆地

伦坡拉盆地位于青藏高原中部、拉萨地体与羌塘地体间的班公湖—怒江缝合带之上(王剑等, 2004), 是在燕山褶皱带基础上, 以古生界地层和中生界海相碳酸盐岩、碎屑岩、基性火山岩和火山碎屑岩为基底而发展起来的, 具有走滑特征的新生代断拗盆地. 早始新世, 盆地开始形成并接受湖泊相沉积. 伦坡拉盆地新生代沉积在盆地内均不整合于老地层之上, 并广泛分布, 最大沉积厚度超过 4000 m, 沉积序列由老到新分别为牛堡组(E_{1-2n})和丁青湖组(E_3d), 呈局部平行不整合接触(王开发等, 1975; 赵政璋等, 2001).

早始新世, 班公湖—怒江缝合带在北西—南东向伸展作用下发生活化, 上地壳发生裂陷作用, 基底快速沉降, 主要在滨浅湖相—半深湖环境中沉积了牛堡组地层(马鹏飞等, 2013); 进入渐新世之后, 盆地的沉降速率开始放缓, 在半深湖—深湖环境中沉积了丁青湖组地层. 渐新世末期, 受印度板块向北俯冲造成的区域上南北向挤压的影响, 区域应力场逐渐由伸展转变为挤压, 盆地热沉降作用逐渐被抵消, 盆地基底开始抬升并遭受剥蚀(张克银等, 2000; 马

鹏飞等,2013).

3.5 囊谦盆地

囊谦盆地是在印度板块与欧亚板块碰撞以及中生代基底经历长期隆升剥蚀的基础上,经早期逆冲—推覆和晚期走滑—拉分作用形成的(潘桂棠,1990;潘裕生等,1998;张克信等,2007),是青藏高原东缘第三纪走滑拉分盆地的典型代表。囊谦盆地新生代的构造演化主要可以划分为以下3个阶段(姜勇彪等,2011):早期挤压推覆前陆盆地、中期走滑拉分盆地和晚期走滑挤压前陆盆地。始新世晚期—渐新世受印度与欧亚板块碰撞远程效应的影响,囊谦盆地在周缘逆冲推覆断层的控制之下沉积了一套以河流相为主的贡觉组($E_{1-2}g$)红色粗碎屑岩系。随后,印度板块继续向北俯冲,在青藏高原东部发育大规模的走滑拉分断层,使得早期的前陆盆地演变为走滑—拉分盆地,沉积了一套红色湖相泥岩、膏岩沉积,并伴有大规模的火山活动(周江羽等,2002;姜勇彪等,2011)。随着印度板块与欧亚板块碰撞进一步加剧,原先走滑拉分应力转变为走滑挤压应力,拉分盆地发育结束,形成晚期的褶皱冲断带及前陆盆地,沉积了盆地上部的粗碎屑岩系(姜勇彪等,2011)。周江羽等(2011)通过对盆地地层内部细碎屑岩孢粉和古植物、火山碎屑岩和侵入岩的 $^{40}Ar/^{39}Ar$ 年代学的研究,揭示出盆地充填沉积物形成于38~28 Ma。

3.6 喜马拉雅残留海

古新世—始新世期间,藏南残留海以北为以林子宗群火山岩为主的冈底斯陆缘火山弧带(简称冈底斯带),火山弧总体呈东西向展布,其上广泛分布火山洼地。在林子宗群火山喷发间歇期,火山洼地内充填了河湖相碎屑岩系(张克信等,2010)。冈底斯带南侧为覆盖雅鲁藏布—喜马拉雅—印度板块北缘的藏南新特提斯残留海。古新世—始新世,喜马拉雅海具有明显的沉积相带分异,从北向南,海水由浅海到深海再到浅海。北带的浅海分布带很窄,也较局限,见于仲巴和巴噶地区。南带的浅海分布带宽,分布区域较大,见于岗巴—定日一带,地层层序自下而上为基堵拉组(E_1j)、宗浦组($E_{1-2}z$)和遮普惹组(E_2z)。从东向西,半深海—深海沉积沿江孜—萨嘎—郭雅拉—桑麦一线分布。中带东部的江孜为甲查拉组(E_{1jc})为代表的半深海斜坡扇沉积建造,中带西部为蹬岗组($E_{1-2}d$)、郭雅拉组(E_2g)和盐多组(E_2y)为代表的半深海—深海含放射虫硅质岩和混杂岩建造(蹬岗混杂岩($E_{1-2}dgm$)、盐多混杂盐(E_2ym)和桑麦混杂岩(E_2ym)),中带的海水东浅西深(张克信

等,2007;Zhang *et al.*,2010;Liang *et al.*,2012)。始新世末期,新特提斯残留海全面向西退去,从而全面结束了该带的海相沉积地层(李国彪等,2005;王成善等,2005)。该带连同冈底斯陆缘弧的大部分地区被构造抬升为剥蚀区,印度与欧亚板块沿雅鲁藏布江带完成了全面碰撞,进入陆内演化阶段。

3.7 成都盆地

成都盆地位于青藏高原东缘龙门山褶皱冲断带以东、扬子克拉通西缘,为挤压构造背景下的构造挠曲压陷盆地。古新世—始新世受龙门山断裂活动影响,成都盆地基底形态为西陡东缓,盆底向西倾斜,沉降中心位于盆地西缘紧靠龙门山冲断带一侧的雅安—名山区。形成的沉积地层厚度特征是西厚东薄。成都盆地沉积序列从老至新为大溪组(K_2E_1d)、名山组($E_{1-2}m$)、柳嘉组($E_{1-2}lj$)、芦山组($E_{2-3}l$)、凉水井组(N_1s)和大邑组(N_2Qpd) (四川省地质矿产局,1997;王凤林等,2003;曾宜君等,2004),与下伏白垩纪陆相地层为连续沉积。盆地内新生界岩石地层序列在渐新统与中新统、中新统与上新统之间均存在沉积间断与角度不整合面。

古近纪时期,成都盆地是受龙门山南段推覆构造带控制的压陷盆地。龙门山南段推覆构造带大致由3条走向北东的巨大断裂,以及数个巨型断片构成,其推覆构造的发展与压陷盆地演化密切相关(刘树根,1993;陶晓风,1999)。盆地周缘造山带自晚燕山期以来呈隆升状态。盆地西缘因受龙门山断裂影响地势较陡,盆地东缘较平缓。

进入新近纪以来,成都盆地发生明显的抬升并遭受不同程度的剥蚀。在盆地西部龙门山前主要受晚新生代青藏高原构造抬升的东逸挤压作用影响,同时也受中央造山带和扬子板块在东缘的抵挡作用,形成了压陷盆地。盆地沉积中心位于龙门山断裂东南侧,盆地呈西北深、东南浅的格局。

3.8 鄂尔多斯盆地

鄂尔多斯盆地是在华北中元古代—古生代克拉通浅海台地基础上发育起来的大型陆内叠合型残留盆地(Ren and Xiao,2002)。作为中国东西两大构造单元的交汇地带,鄂尔多斯盆地自中生代以来,由于受到构造体制转换和多次构造运动的影响,在其周缘地带形成了复杂的盆—山耦合系统(王锡勇等,2010)。鄂尔多斯盆地新生代沉积序列由老到新分别为寺口子组(E_3s)、清水营组(E_3q)、红柳沟组(N_1h)和干河沟组(N_2g)。晚白垩世—始新世,鄂尔多斯盆地在区域挤压应力的作用下整体抬升,使得盆

地周邻地区的断陷盆地发生构造反转,表现为挤压构造背景下的隆升剥蚀,盆地主体部分普遍缺失古近纪沉积(刘池洋等,2006;郑孟林等,2006).构造热年代学研究显示鄂尔多斯盆地快速抬升和由剥蚀作用引起的快速冷却事件主要发生在 23 Ma 之前,抬升速率在盆地西部为约 23 m/Ma,在盆地东部为约 34 m/Ma(任战利,1995).同时该次构造事件导致盆地内部古近系和新近系之间的地层不整合接触.渐新世时期,鄂尔多斯盆地在 EW-SE 向引张应力作用下在盆地周缘发育地堑式断陷盆地(张岳桥和廖昌珍,2006),沉积了以寺口子组(E_3s)和清水营组(E_3q)为代表的内陆河湖相沉积;中新世时期,盆地周缘受到 NE-SW 至 NNE-SSW 向引张应力作用,原来的断陷盆地恢复区域性热沉降(王锡勇等,2010),广泛发育以红柳沟组(N_1h)和干河沟组(N_2g)为代表的河湖相沉积序列.进入晚上新世,盆地内大量发育红土堆积,至第四纪初,风成黄土堆积开始发育,并逐渐形成现代地貌.

4 中国西部新生代盆地演化与板块运动—构造隆升耦合

晚白垩世末—古新世初,印度/欧亚板块碰撞以

来,受印度洋板块朝北运移、挤压和碰撞的影响,中国大陆西部发生了巨大的变形(图 4)(Molnar and Tapponnier, 1975; Tapponnier *et al.*, 1990, 2001; Wang *et al.*, 2008; Yin, 2010; 潘桂堂等, 2013).印度板块与欧亚板块的碰撞以及其后印度板块的向北俯冲挤压,对中国西部新生代沉积盆地的发育和演化产生了重大影响.

古新世—早中始新世,中国西部地区最重要的构造事件是印度板块与欧亚板块的强烈俯冲碰撞.此次碰撞引发了中国西部大规模地壳挤压缩短,岩石圈增厚和抬升(钟大赉等,2001).古地磁学研究显示,在 58~50 Ma 之间,印度板块与亚洲板块的汇聚速率较大且保持相对稳定;50 Ma 之后,印度板块与亚洲板块的汇聚速率发生急剧降低(Lee and Lawver, 1995; Molnar and Stock, 2009; Copley *et al.*, 2010)(图 5),指示印度板块与欧亚板块的主碰撞期应为 58~50 Ma.而青藏高原热年代学研究表明 58~50 Ma 高原表现为强烈隆升阶段(Wang *et al.*, 2011; Zhang *et al.*, 2013),这与古地磁学揭示的主碰撞期一致(图 5).约 40 Ma 以来喜马拉雅沉积缺失,印度与欧亚板块全面完成碰撞,喜马拉雅残留海盆全面退缩(李国彪等,2005;张克信等,2007,2010,2013).碰撞之后高原的构造变形表现为近南北向的挤压缩短和近东西向的走滑挤出及其转

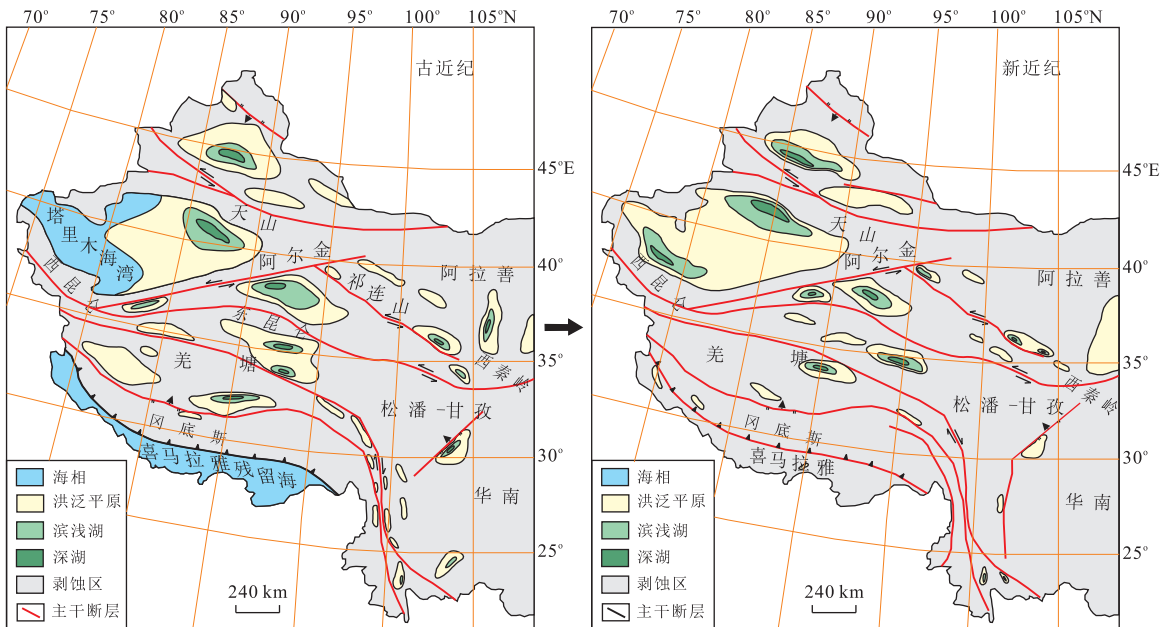


图 4 中国西部新生代构造—岩相古地理格局演替

Fig.4 Tectonic-lithofacies paleogeography evolution of the western China during the Cenozoic

据以下参考文献:王鸿祯,1985;叶得泉等,1993;陈发景和汪新胜,1997;邓万明,2003;刘训,2004;王成善等,2005;Yin *et al.*, 2008;张克信等, 2010,2013;Carroll *et al.*, 2010

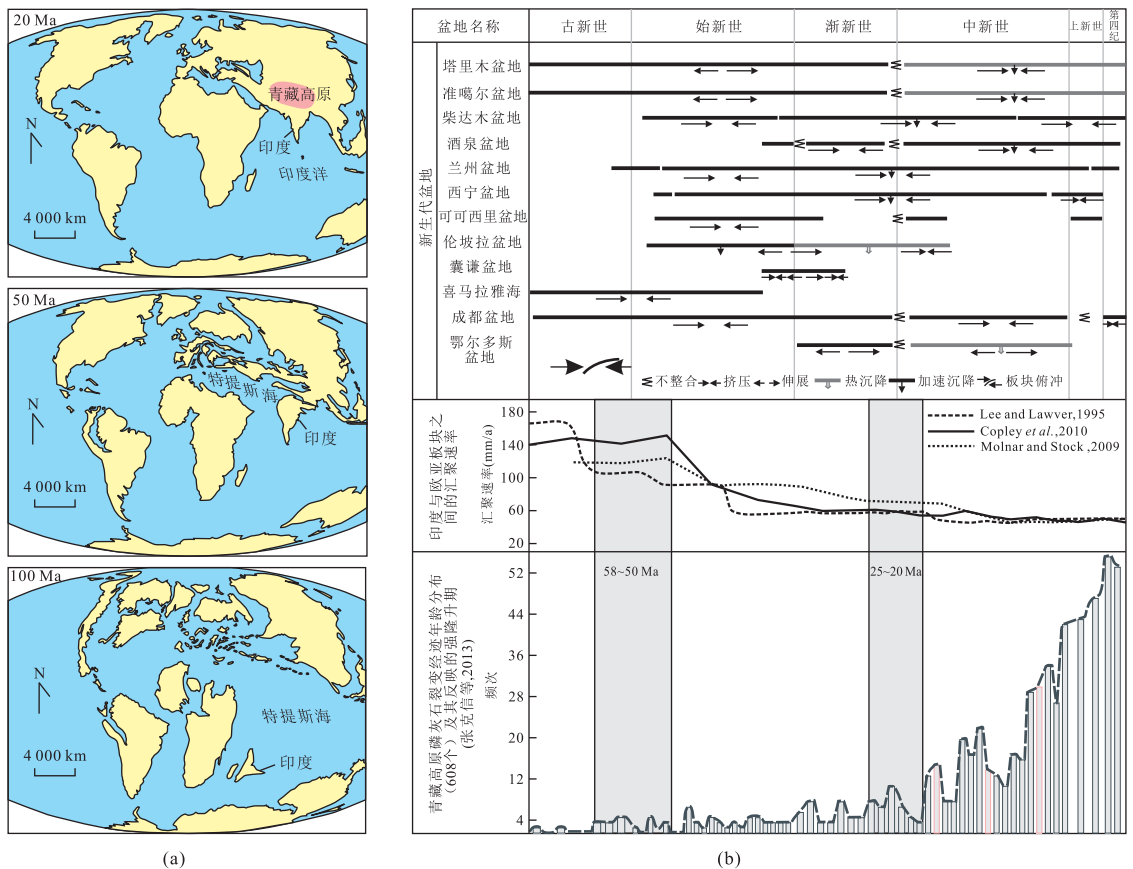


图 5 中国西部新生代盆地演化与板块运动—构造隆升耦合

Fig.5 Evolution of Cenozoic sedimentary basins in western China coupling with plate movement and tectonic uplift

据以下参考文献:唐天福等,1992;刘树根,1993;Lee and Lawver,1995;岳乐平等,2003;赵加凡等,2004;Dai et al.,2005,2006;王成善等,2005;李国彪等,2005;邵龙义等,2006;刘池洋等,2006;郑孟林等,2006;Haq,2007;王清晨等,2007;Wang et al.,2008;Molnar and Stock,2009;李丕龙等,2010;Copley et al.,2010;张克信等,2007,2010,2013;Zhang et al.,2008,2010,2013;周江羽等,2011;马鹏飞等,2013

换构造体系(Tapponnier and Molnar,1976).南北向的缩短构造在空间上普遍发育,主要包括边缘造山带及随后发育的一系列近东西向的逆冲推覆构造(Yin and Harrison,2000;吴珍汉等,2009).东西向走滑挤压及转换构造主要发育在高原内部,在高原北部尤为典型,以大规模的左行走滑边界断层为主,包括阿尔金断裂和东昆仑断裂等.与此同时,随着陆陆碰撞和俯冲作用的持续,在欧亚大陆与印度板块的碰撞及其远程效应影响下,古天山、昆仑山、祁连山、阿尔金造山带依次复活并构造抬升,发生陆内造山作用.中国西部主要大型盆地如柴达木盆地、酒泉盆地、可可西里盆地、兰州—西宁盆地、伦坡拉盆地均在早—中始新世期间形成(Dai et al.,2005,2006;宋春晖,2006;Wang et al.,2008;Zhang et al.,2013)(图5),在中国西部形成近EW或NWW向的造山带和内陆盆地相间分布的构造—地貌景观,而大型盆地与山脉之间以压性—压扭性活动断裂(逆

冲断层与走滑断层)为主要边界.古近纪期间羌塘构造带则呈现出伸展构造背景,伦坡拉盆地、尼玛盆地等一系列东西向展布盆地继续快速沉降.塔里木盆地和准噶尔盆地均为伸展盆地,古天山整体隆升,分别向南、向北为以上两个盆地提供物源.而塔里木盆地西南部在古近纪期间仍然为滨浅海相沉积,为新特提斯海的一部分,期间发生多次海侵(郝谄纯等,2002).

到新近纪,中国西部地区构造—地貌演化进入了一个新的阶段.25~20 Ma之间印度板块与亚洲板块之间的碰撞速率基本呈下降的趋势(图5),高原表现为强烈隆升,该时期是青藏高原重大构造变革期(Sobel et al.,2006;Xiao et al.,2012;王二七,2013).新近纪以来青藏高原进入强烈陆内汇聚挤压阶段,中国西部地区整体处于挤压构造背景之下.新近纪以来,塔里木盆地西部岩相古地理面貌发生重大转变,随着西昆仑和天山山脉的隆升,渐新世末期

塔里木盆地海相沉积结束,进入大型内陆盆地演化阶段(张克信等,2010);塔里木盆地、准噶尔盆地、酒泉盆地、成都盆地及鄂尔多斯盆地在挤压构造背景下在盆地周缘均发育冲断带,古近系地层与新近系地层之间均呈角度不整合接触(图 5);可可西里—沱沱河地区区域性的角度不整合面发育和盆地内古近纪地层发生抬升变形(Wang *et al.*, 2002, 2008);同时,受印度板块向北俯冲造成的区域上南北向挤压的影响,渐新世末伦坡拉盆地区域应力场逐渐由伸展转变为挤压(马鹏飞等,2013).喜马拉雅地区新近纪构造变形相当强烈,构成了一个规模巨大的逆冲断层系统.在东西向伸展作用的影响下,藏南发育了一系列南北向的断陷盆地(Blisniuk *et al.*, 2001),同时在川西—藏东地区发育一系列走滑—拉分盆地(潘桂棠等,1990;Zhang *et al.*, 2010).青藏高原北缘地区中新世中晚期强烈的构造抬升导致高原向北的扩展.西昆仑山系列向北逆冲断层强烈活动,蚀源区向北扩展,现代西昆仑山成形;阿尔金山断裂全面活动强烈,阿尔金山现代构造地貌格局形成;祁连山系列逆冲断层再次强烈活动,祁连山继续向北扩展(王国灿等,2010).上新世约 8 Ma 以来,青藏高原及周边普遍加速隆升,周缘山系快速抬升,中国西部地区现代盆—山地貌格局基本成型(张培震等,2006;Zheng *et al.*, 2006;Zhang *et al.*, 2013). 3.6 Ma 以来,青藏高原和天山山脉强烈隆升,在高原南北及天山山前发育大规模的磨拉石建造(李吉均等,1979;Li and Fang, 1999).

5 结论

(1)新生代以来,印度板块与欧亚板块的碰撞及其远程效应导致 SSW-NNE 向纵向挤压,在中国西部形成了近 EW 或 NWW 的造山带和内陆盆地相间分布构造—地貌的景观.在不同的盆地动力学背景下,又分别发育断陷盆地、压陷—坳陷盆地、走滑—拉分盆地和残留海—前陆盆地.

(2)盆—山系统是中国西部新生代构造的基本格局,而盆—岭地貌是中国西部新生代的主要地貌单元.造山带与沉积盆地之间存在明显的耦合关系,中国西部新生代盆地的发育与造山带的再活动密切相关.大型盆地与山脉之间以压性—压扭性活动断裂(逆冲断层与走滑断层)为主要边界,同时断裂控制了盆地的构造演化和沉积相的展布.

(3)在不同的盆地类型和构造背景下,发育不同

的沉积充填序列.断陷盆地和走滑—拉分盆地均以近源堆积的沉积体系为主,压陷盆地以远源堆积的沉积体系为主,而残留海—前陆盆地则兼具近源堆积和远源堆积两种沉积类型.

(4)印度板块与欧亚板块的碰撞以及其后印度板块的向北俯冲挤压,对中国西部新生代沉积盆地的发育和演化产生了重大影响.新生代期间中国西部地区最重要的构造事件是印度板块与欧亚板块的强烈俯冲碰撞及其后青藏高原的迅速隆升.中国西部特别是青藏高原地区新生代构造岩相古地理演化与板块运动和构造隆升之间存在明显的耦合.

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