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海南荔枝沟中三叠世酸性火山岩年代学、地球化学特征及其构造意义

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摘要:海南岛位于印支与华南陆块的交界地带, 构造演化复杂, 因此海南岛的拼合历史和构造属性一直备受争议。在海南岛三亚荔枝沟地区新识别出的酸性火山岩对于探讨海南岛早中生代的构造属性具有重要指示意义。LA-ICP-MS 锆石 U-Pb 定年结果显示该套酸性火山岩中流纹岩的成岩年龄为 241 ± 6 Ma (MSWD=0.9)。地球化学特征表明该套火山岩样品的 SiO_2 含量介于 72.34%~77.83%, 具较高的 Al_2O_3 (10.51%~13.53%)、 K_2O (2.85%~4.85%)、 Na_2O (1.75%~3.79%) 含量和 A/CNK 比值 (0.99~2.07), 为过铝质钙碱性酸性火山岩。样品稀土元素含量为 $\sum\text{REE}=87 \times 10^{-6} \sim 177 \times 10^{-6}$, 富集轻稀土元素, 显示右倾岛弧型稀土元素配分模式, Eu 负异常明显。样品富集大离子亲石元素, 亏损高场强元素, $\epsilon_{\text{Nd}}(t)$ 变化于 -12.1~-11.3 之间, 可被解释为变沉积物部分熔融后经历分离结晶作用的产物。区域对比研究表明该套酸性火山岩可与三江地区的金沙江-哀牢山-松马构造带的中三叠世岩浆岩进行类比, 应形成于印支与华南陆块汇聚的构造背景。

关键词: 中三叠世; 酸性火山岩; 年代学; 地球化学; 碰撞后; 海南岛。

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Geochronological and Geochemical Constraints on Lizhigou Middle Triassic Felsic Volcanic Rocks in Hainan and Its Tectonic Implications

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Abstract: Hainan Island, which is located between the Indochina and South China blocks, has undergone multiphased structural overprinting and complex tectonic evolution. However, the Late Paleozoic tectonic regime and amalgamation process of the South China with Indochina blocks are still controversial. The newly-identified felsic volcanic rocks in the Lizhigou area are important for understanding the Early Mesozoic tectonic evolution in Hainan. LA-ICP-MS zircon U-Pb dating results show that the representative rhyolite samples yield a weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age of 241 ± 6 Ma (MSWD=0.9). These samples contain SiO_2 ranging from 72.34% to 77.83%, Al_2O_3 from 10.51% to 13.53%, K_2O from 2.85% to 4.85% and Na_2O from 1.75% to 3.79%. The A/CNK values range from 0.99 to 2.07. These samples have REE concentrations of 87×10^{-6} to 177×10^{-6} , and Eu/Eu* of 0.49~0.65. They are characterized by subparallel right-sloping REE pattern and enrichment in LREEs and LILEs, and depletion in HFSEs, similar to those of arc volcanics. Their $\epsilon_{\text{Nd}}(t)$ values range from -12.1 to -11.3. These felsic volcanic rocks are the products of partial melting of metasediments with significant fractionation crystallization. In combination with other available data, it is inferred that the Lizhigou felsic volcanic rocks can be compared with the Middle-Late Triassic igneous

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rocks along the Jinshajiang-Ailaoshan-Songma zone, and formed during the assemblage of the South China with Indochina blocks related to East Paleotethyan consumption.

Key words: Middle Triassic; felsic volcanic rock; geochronology; geochemistry; post-collisional setting; Hainan Island.

0 引言

东南亚地区保存了大量与古特提斯构造演化有关的岩浆岩及沉积记录,是研究古特提斯洋形成、演化、俯冲及消亡的天然实验室。海南岛地处太平洋、菲律宾板块和华南、印支陆块的交界部位,也是古太平洋构造域和古特提斯构造域的叠置区(图 1a),对于研究古特提斯洋的演化以及华南和印支陆块的聚合具有重要的意义(汪啸风等,1991;李孙雄等,2005;陈新跃等,2011,2014;唐立梅等,2013;温淑女等,2013;He *et al.*,2017,2018;Li *et al.*,2018)。

已有的研究表明海南岛北部地区存在石炭纪时期的岛弧及弧后盆地岩浆记录,并可以与金沙江—哀牢山—松马缝合带内相关的石炭纪岩浆岩进行对比(图 1a;Li *et al.*,2002,2018;He *et al.*,2017)。同时,岛内还广泛分布有大量的二叠纪—三叠纪岩浆岩(夏邦栋等,1990;李孙雄等,2005;谢才富等,2005;Li *et al.*,2006;Zhang *et al.*,2011;陈新跃等,2011;温淑女等,2013)。作为研究的热点,这些二叠纪—三叠纪岩浆岩的岩石成因及其构造背景众说纷纭,与之对应的二叠纪—三叠纪的构造演化史也未形成统一的定论。谢才富等(2005)在海南岛中部识别出了钾玄质侵入岩,并认为华南陆块与印支陆块的碰撞拼贴发生于 287~278 Ma,且华南陆块和印支陆块拼贴之间的缝合带可能是从越南松马地区,沿北部湾和云开大山北缘连接至武夷山地区。而其他学者结合海南岛早三叠世 NWW 向的构造变形(陈新跃等,2006;Zhang *et al.*,2011)和相关的三叠纪碱性岩、A 型花岗岩、高钾钙碱性花岗岩和镁铁质侵入岩等岩石组合,认为海南岛地区的 250 Ma 岩浆岩形成于同碰撞阶段,并在 240 Ma 之后进入到碰撞后阶段(谢才富等,2005;唐立梅,2010;唐立梅等,2013;陈新跃等,2014)。

最近,笔者在海南岛三亚荔枝沟地区新识别出了一套酸性火山岩组合(包括流纹岩、凝灰岩、火山集块岩)。因此,本文选取该套酸性火山岩开展 LA-ICP-MS 锆石 U-Pb 年代学、全岩主、微量元素及 Sr-Nd 同位素研究,并结合区域地质资料研究该套酸性火山岩的岩石成因及其大地构造背景,进而为探讨古特提斯洋的构造演化提供重要依据。

1 区域地质背景与岩石学特征

海南岛地处中国东南部,以琼州海峡与华南大陆相隔。海南岛由北向南发育王五—文教、昌江—琼海、尖峰—吊罗、九所—陵水等多条近东西向的断裂(广东省地质矿产局,1988;沈宝云等,2016);由西向东发育了戈枕、白沙等北东向断裂(夏邦栋等,1990,1991;汪啸风等,1991)。海南岛主要出露古生界,其次为元古宇和中新生界(马大铨等,1998;龙文国等,2002)。中元古代长城纪的抱板群和青白口纪的石碌群被认为是海南岛最古老的地层,主要出露于抱板、冲卒岭、昌江等地(广东省地质矿产局,1988;汪啸风等,1991)。下古生界寒武系和奥陶系由一套浅变质页岩、砂岩、粉砂岩和板岩组成(夏邦栋等,1990;汪啸风等,1991)。志留系为浅海相砂岩和一套过渡类型的碎屑岩、灰岩和板岩组合(胡宁等,2002;龙文国等,2007)。分布于九所—陵水断裂以北的上古生界主要出露有泥盆系砂岩、石炭系板岩和变火山岩、下二叠统灰岩以及中二叠统砂岩(夏邦栋等,1990;汪啸风等,1991;胡宁等,2002;龙文国等,2007)。中生界在岛内仅出露上三叠统碎屑岩沉积和白垩系砂岩(夏邦栋等,1990;汪啸风等,1991)。岛内新生界发育完整,主要分布于海南岛北部(唐立梅,2010)。

由于强烈的岩浆活动,海南岛内岩浆岩广泛分布,占全岛面积 51% 的侵入岩可划分为海西、印支和燕山期花岗岩,其中海西与印支期花岗岩分布最广,其次为燕山晚期的花岗岩类。喷出岩以王五—文教断裂以北的新生代玄武岩为主(广东省地质矿产局,1988;夏邦栋等,1990;李献华等,2000;He *et al.*,2017;Li *et al.*,2018)。目前的资料表明:岛内二叠纪至三叠纪出现的侵入岩体包括五指山、育才、通什、道票岭、禄马、尖峰、袁水、立才、六连岭和万宁等多个超单元(广东省地质矿产局,1988;温淑女等,2013)。

本次研究的酸性火山岩采自海南省九所—陵水断裂以南的三亚市荔枝沟地区(图 1b~1c),区内还分布有三叠纪和白垩纪花岗岩。本文所研究的火山岩出露于奥陶系南碧沟组中。野外剖面(自东向西)可依次观察到砾岩、流纹岩(图 2a)、石英砂岩、黑云

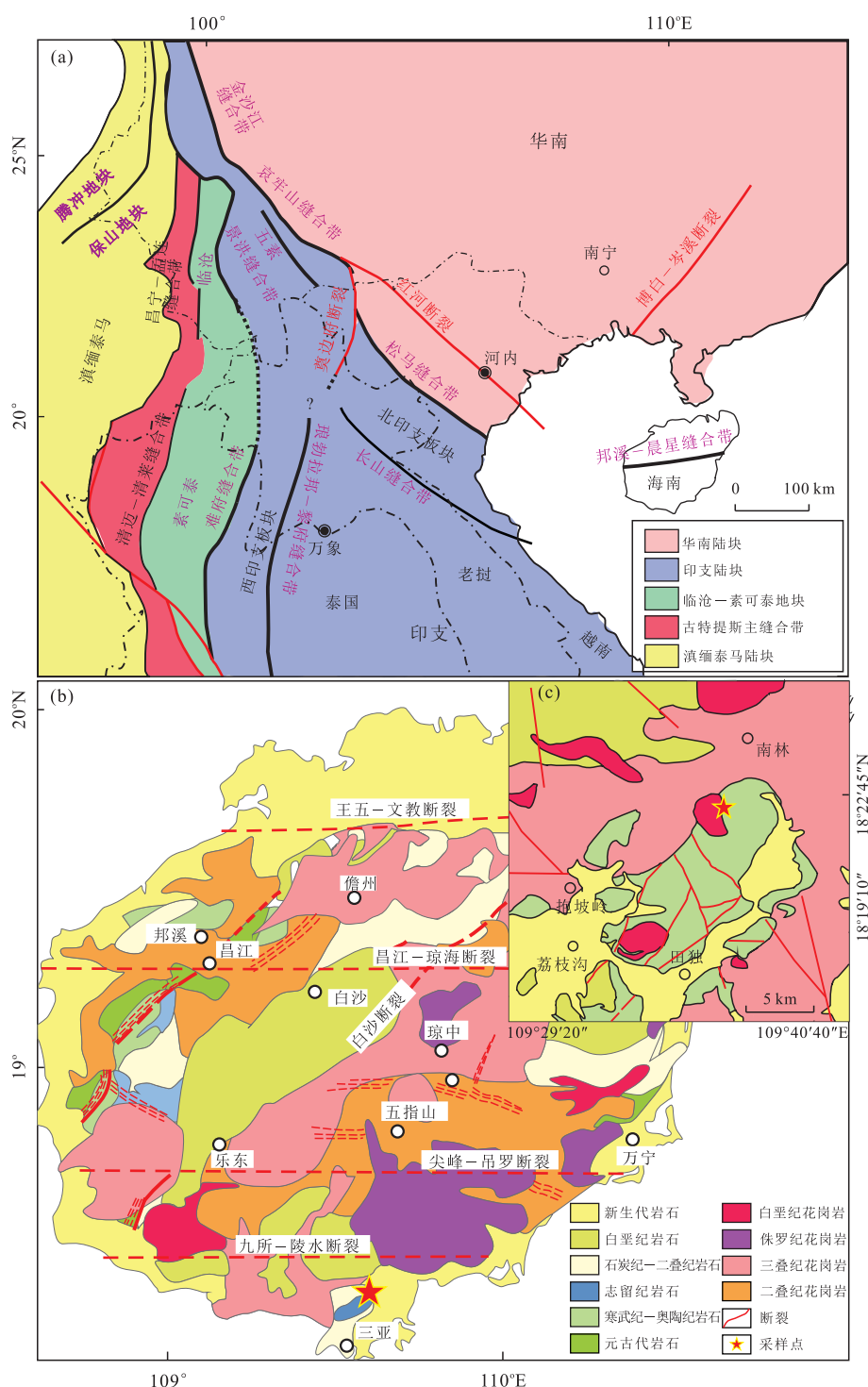


图 1 东南亚地区构造划分(a)、海南岛地质简图(b)和研究区地质概况及采样点位置(c)

Fig.1 Tectonic outline of Southeast Asia (a), geological sketch of Hainan Island (b) and geological sketch of the study area showing the sampling location (c)

图 a 据 Wang *et al.* (2010, 2018) 修改; 图 b 据 Zhang *et al.* (2011) 修改

母花岗岩、火山集块岩(图 2e)、凝灰岩(图 2c)以及层间砾岩,其中的酸性火山岩被上覆的砾岩角度不整合覆盖。野外观察流纹岩为浅红褐色,颗粒较细,易破碎;凝灰岩为浅紫灰色,风化较强;火山集块岩

为深灰黑色,出露体积较大。镜下观察表明流纹岩中的矿物斑晶主要为石英和长石,含少量黑云母;凝灰岩中可见石英和长石碎屑;火山集块岩的斑晶主要为石英和长石,基质为玻璃质。长石斑晶可见不同程

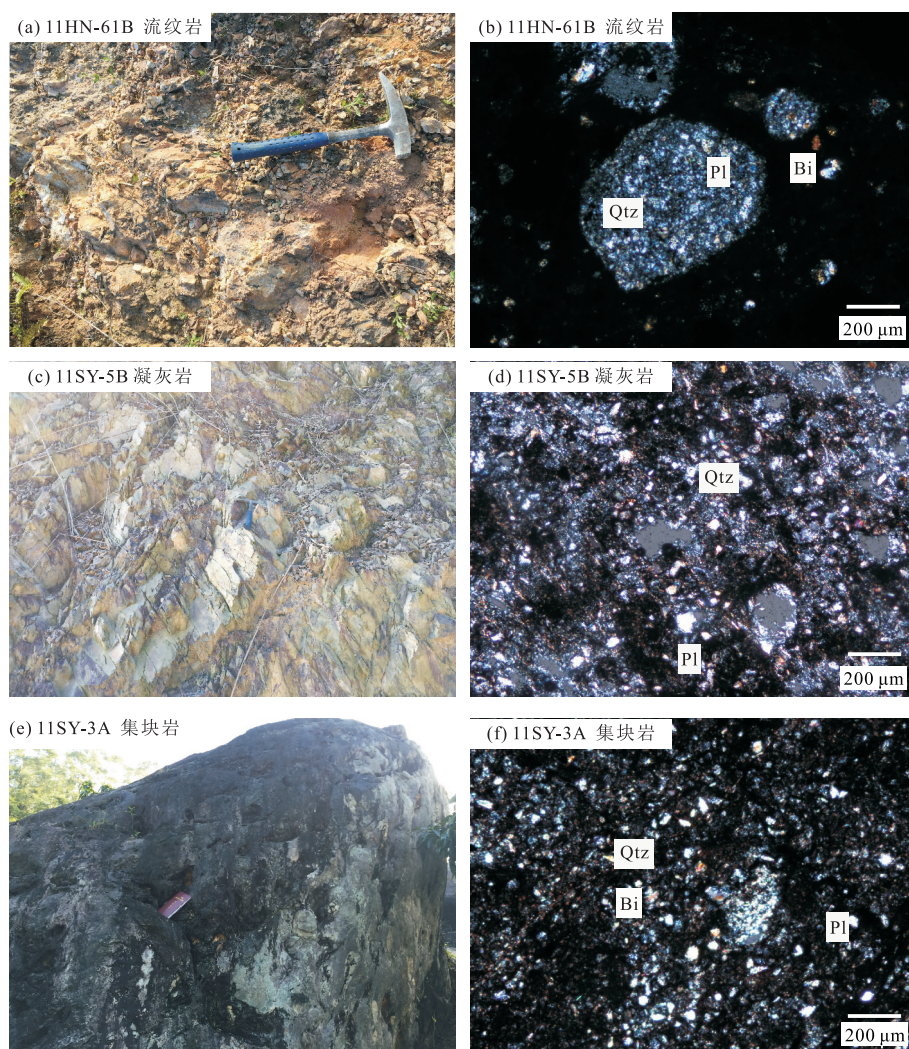


图 2 海南荔枝沟酸性火山岩野外照片(a,c,e)和岩石显微照片(b,d,f)

Fig.2 Field photographs (a,c,e) and photomicrographs (b,d,f) for the Lizhigou felsic volcanic rocks in Hainan

Qtz.石英; Pl.斜长石; Bi.黑云母

度的绢云母化。

2 样品分析方法

通过人工重砂法和电磁选,从新鲜的样品中选出无磁性重矿物样品.在双目镜下挑选出表面平整光洁,无裂隙、无包体、透明干净、形态较好的锆石颗粒,将其粘在双面胶上,并用混合固化剂的环氧树脂胶结,固化后细磨抛光制成样品靶.通过阴极发光(CL)探查锆石样品的内部结构,确定锆石的成因类型.锆石阴极发光图像由中山大学的 Carl ZEISS Σ IGMA 场发射扫描电镜完成。

锆石 U-Pb 同位素分析由中国科学院广州地球化学研究所同位素地球化学国家重点实验室的 ICP-MS Elan6100DRC 激光离子探针完成.详细的

分析步骤和数据处理方法参照 Xia *et al.* (2004),数据处理软件使用 GLITTER4.4,年龄计算及绘制相应图件使用 Ludwig(2003)的 Isoplot 宏插件。

将岩石样品在无污染环境下粉碎至 200 目,用于主微量元素和 Sr、Nd 同位素分析.全岩粉末样的主微量元素和 Sr-Nd 同位素组成分析在中国科学院广州地球化学研究所同位素地球化学国家重点实验室完成.主量元素分析采用 Rigaku RIX 2000 型荧光光谱仪(XRF)分析,分析精度优于 1%~5%,详细分析方法见 Li *et al.* (2005).微量元素分析采用 PE Elan 6000 型电感耦合等离子体质谱仪(ICP-MS),仪器分析精度优于 2%~5%,具体分析流程和方法见刘颖等(1996).Sr 和 Nd 同位素组成分析采用 Neptune 多接收等离子质谱仪(MC-ICP-MS)进行⁸⁷Sr/⁸⁶Sr 和¹⁴³Nd/¹⁴⁴Nd 比值测试.测定过程中

表 1 海南荔枝沟酸性火山岩样品(11SY-3A) LA-ICP-MS 锆石 U-Pb 定年结果

Table 1 LA-ICP-MS U-Pb zircon geochronological data for the Lizhigou felsic volcanic rocks in Hainan

样品编号	Th	U	Th/U	$^{207}\text{Pb}/^{206}\text{Pb}$	1 σ	$^{207}\text{Pb}/^{235}\text{U}$	1 σ
11SY-3A-01	574	898	0.64	0.053 76	0.001 63	0.286 10	0.008 73
11SY-3A-02	450	1 023	0.44	0.056 24	0.001 70	0.298 14	0.009 07
11SY-3A-03	192	872	0.22	0.058 33	0.002 03	0.301 60	0.010 89
11SY-3A-04	386	942	0.41	0.061 80	0.001 87	0.328 12	0.010 00
11SY-3A-05	432	733	0.59	0.055 28	0.001 67	0.276 96	0.008 39
11SY-3A-06	521	803	0.65	0.052 33	0.001 57	0.285 16	0.008 71

样品编号	$^{206}\text{Pb}/^{238}\text{U}$	1 σ	$^{207}\text{Pb}/^{206}\text{Pb}$	1 σ	$^{207}\text{Pb}/^{235}\text{U}$	1 σ	$^{206}\text{Pb}/^{238}\text{U}$	1 σ
11SY-3A-01	0.038 62	0.001 18	361	64	255	7	244	7
11SY-3A-02	0.038 45	0.001 16	461	67	265	7	243	7
11SY-3A-03	0.037 43	0.001 17	543	76	268	8	237	7
11SY-3A-04	0.038 56	0.001 18	733	60	288	8	244	7
11SY-3A-05	0.036 37	0.001 11	433	73	248	7	230	7
11SY-3A-06	0.039 53	0.001 21	298	70	255	7	250	7

的质量分馏效应分别采用 $^{87}\text{Sr}/^{86}\text{Sr}=0.119\ 4$ 和 $^{143}\text{Nd}/^{144}\text{Nd}=0.721\ 9$ 进行校正.详细分析测试流程见 Yang *et al.*(2006).

3 LA-ICP-MS 锆石 U-Pb 年代学测试结果

分析的锆石样品 11SY-3A 为海南荔枝沟流纹岩,相关的同位素比值结果和年龄结果列于表 1 中.锆石 U 和 Th 的含量分别为 $733\times 10^{-6}\sim 1\ 023\times 10^{-6}$ 和 $192\times 10^{-6}\sim 574\times 10^{-6}$;锆石 Th/U 比值为 $0.22\sim 0.65$,大部分大于 0.4,其 CL 图像显示锆石形态完整,有明显的振荡韵律环带结构(图 3a),为典型的岩浆成因锆石.其中 6 个锆石测点的 $^{206}\text{Pb}/^{238}\text{U}$

谐和年龄在 $230\sim 250\ \text{Ma}$ 之间,均靠近谐和曲线(图 3b).其加权平均年龄为 $241\pm 6\ \text{Ma}$,MSWD=0.9,时代为中三叠世,代表了该套酸性火山岩的形成时间.

4 主微量元素及 Sr-Nd 同位素特征

荔枝沟地区酸性火山岩的主、微量元素和同位素分析结果列于表 2 中.11 个荔枝沟地区酸性火山岩样品烧失量为 $0.55\%\sim 2.25\%$, SiO_2 含量为 $72.34\%\sim 77.83\%$,具有低含量的 TiO_2 ($0.34\%\sim 0.54\%$), MgO ($0.54\%\sim 1.27\%$) 以及较高含量的 Al_2O_3 ($10.51\%\sim 13.53\%$), K_2O ($2.85\%\sim 4.85\%$) 和 Na_2O ($1.75\%\sim 3.79\%$) 的含量较高,全碱含量($\text{K}_2\text{O}+\text{Na}_2\text{O}$)为 $5.83\%\sim 8.26\%$, $\text{Mg}^\#$ 值为 $11\sim$

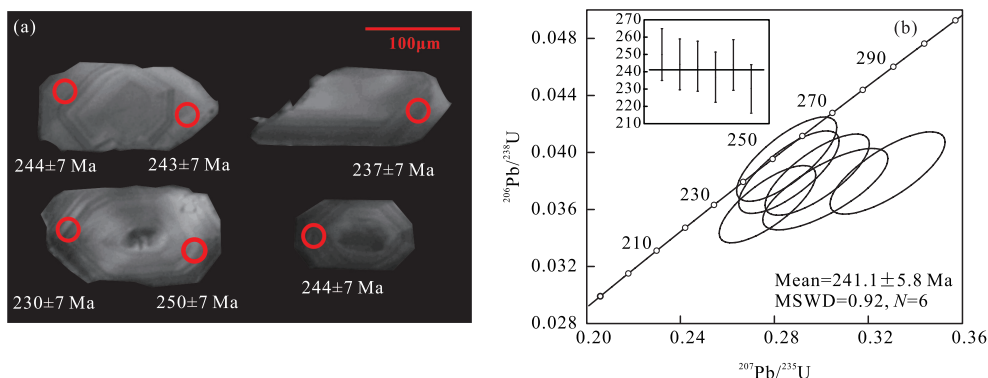
图 3 荔枝沟酸性火山岩(11SY-3A)锆石 CL、测试点位及 $^{206}\text{Pb}/^{238}\text{U}$ 表观年龄(a)和锆石 U-Pb 谐和图(b)Fig.3 CL images, dating spots and $^{206}\text{Pb}/^{238}\text{U}$ apparent age (a) of the zircons and U-Pb concordia diagram for the Lizhigou felsic volcanic rocks (11SY-3A) (b)

表 2 海南荔枝沟地区酸性火山岩主量元素(%)、微量元素(10^{-6})和 Sr-Nd 同位素组成分析结果

Table 2 Major oxides (%), trace elements (10^{-6}) and Sr-Nd isotopic analytical results for the Lizhigou felsic volcanic rocks in Hainan

样品编号	11HN-61B	11SY-3A	11SY-3B	11SY-3C	11SY-3D	11SY-3E	11SY-3F	11SY-3G	11SY-5B	11SY-5C
SiO ₂	76.92	74.83	71.61	70.51	74.03	71.06	71.31	75.56	74.01	74.34
TiO ₂	0.35	0.36	0.46	0.39	0.49	0.51	0.43	0.33	0.53	0.51
Al ₂ O ₃	11.70	10.38	11.38	10.95	12.07	12.74	10.99	10.31	13.07	13.43
Fe ₂ O ₃	2.20	2.97	3.20	3.48	2.60	3.17	2.83	2.65	3.47	3.50
MgO	1.73	3.08	3.54	3.68	2.61	3.35	3.40	2.74	2.32	2.37
CaO	0.62	2.66	3.19	4.16	1.13	1.70	3.69	2.54	0.29	0.32
K ₂ O	4.73	2.79	3.40	3.01	4.61	4.76	4.02	2.98	3.78	3.80
Na ₂ O	0.53	0.89	0.76	0.76	0.97	0.85	0.91	0.89	1.25	0.87
MnO	0.01	0.03	0.04	0.05	0.02	0.02	0.04	0.03	0.01	0.01
P ₂ O ₅	0.04	0.04	0.06	0.06	0.06	0.06	0.08	0.05	0.10	0.09
L.O.I	0.98	1.31	1.68	2.25	0.79	1.10	1.65	1.30	0.60	0.55
Total	99.81	99.35	99.31	99.3	99.38	99.33	99.36	99.37	99.44	99.79
Mg [#]	32	32	37	32	30	42	35	39	33	11
A/CNK	1.65	1.22	1.12	1.01	1.45	1.30	0.99	1.25	2.05	2.07
A/NK	1.81	2.82	2.74	2.91	2.16	2.27	2.23	2.67	2.60	2.88
Sc	7.04	5.09	6.32	5.97	7.33	9.80	7.33	6.23		
V	39.7	25.6	29.8	26.3	33.8	46.0	36.8	29.8		
Cr	20.7	27.1	34.0	25.3	34.6	30.6	36.1	23.0		
Co	3.99	6.33	5.69	5.41	5.20	7.23	6.38	6.00		
Ni	5.78	12.80	13.50	11.20	12.30	16.50	19.60	23.00		
Rb	124.0	84.6	113.0	93.8	115.0	139.0	105.0	97.0		
Sr	82.9	76.6	95.9	84.9	62.0	67.3	89.3	76.9		
Y	9.2	13.2	22.9	22.9	33.0	20.8	23.4	23.3		
Zr	186	178	246	205	412	221	219	227		
Nb	9.10	7.19	9.64	8.19	10.50	11.20	9.10	8.66		
Ba	562	709	675	543	1 090	800	989	474		
La	20.2	19.1	25.7	29.1	37.4	26.9	26.9	23.1		
Ce	41.2	35.5	54.8	60.0	68.4	55.2	55.7	52.6		
Pr	3.75	3.99	6.19	6.90	8.52	6.28	6.54	5.88		
Nd	13.3	15.4	24.9	26.1	32.6	24.9	26.0	24.5		
Sm	2.10	3.17	5.27	5.33	6.83	5.05	5.67	5.12		
Eu	0.40	0.64	0.88	0.91	1.09	0.77	0.86	0.74		
Gd	1.54	2.82	4.35	4.78	6.16	3.73	4.29	3.86		
Tb	0.25	0.43	0.75	0.81	1.00	0.63	0.63	0.63		
Dy	1.23	2.50	4.43	4.70	6.17	3.71	3.78	3.64		
Ho	0.29	0.46	0.84	0.88	1.13	0.79	0.73	0.78		
Er	1.01	1.60	2.56	2.66	3.59	2.43	2.31	2.39		
Tm	0.17	0.26	0.40	0.39	0.53	0.36	0.32	0.36		
Yb	1.37	1.76	2.50	2.61	3.50	2.44	2.25	2.35		
Lu	0.26	0.30	0.41	0.42	0.55	0.39	0.37	0.36		
Hf	4.58	5.38	6.68	6.05	11.60	6.14	5.76	5.72		
Ta	0.50	0.82	0.71	0.66	0.83	0.78	0.62	0.56		
Pb	10.7	15.2	15.8	11.3	16.2	16.2	19.1	15.1		
Th	7.86	7.96	9.75	8.87	12.10	10.00	8.36	8.04		
U	1.89	1.81	2.57	2.30	3.12	2.41	2.37	2.09		
Eu/Eu [*]	0.65	1.38	1.72	2.04	1.98	1.54	1.59	1.74		
(La/Yb) _N	26.6	18.8	35.1	39.5	39.3	31.9	41.9	43.8		
(Gd/Yb) _N	0.58	0.56	1.18	1.33	1.36	1.01	1.18	1.39		
⁸⁷ Rb/ ⁸⁶ Sr		3.209		3.209			3.417			
¹⁴⁷ Sm/ ¹⁴⁴ Nd		0.124		0.124			0.132			
⁸⁷ Sr/ ⁸⁶ Sr		0.751 057		0.746 798			0.751 993			
2σ		6		11			11			
¹⁴³ Nd/ ¹⁴⁴ Nd		0.511 947		0.511 922			0.511 919			
2σ		6		11			11			
(⁸⁷ Sr/ ⁸⁶ Sr) _i		0.740 1		0.735 8			0.740 3			
ε _{Nd} (<i>t</i>)		−11.3		−11.7			−12.1			
<i>T</i> _{DM2} (Ga)		1.67		1.71			1.73			

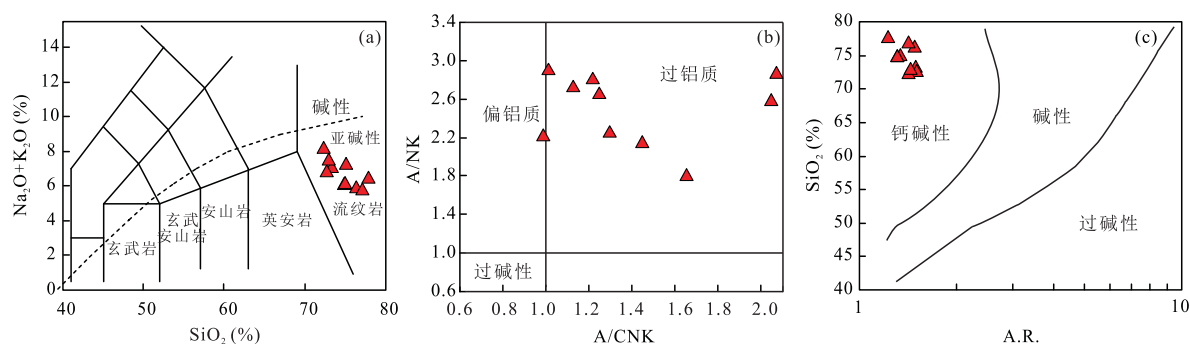
图 4 荔枝沟酸性火山岩 TAS(a)、A/CNK—A/NK(b)和 A.R.—SiO₂ 图解(c)

Fig.4 SiO₂ vs. Na₂O+K₂O(TAS) (a), A/NK vs. A/CNK (b) and A.R. vs. SiO₂ (c) classification diagrams for the Lizhigou felsic volcanic rocks

图 a 据 Maitre *et al.*(1989);图 c 据 Wright (1969)

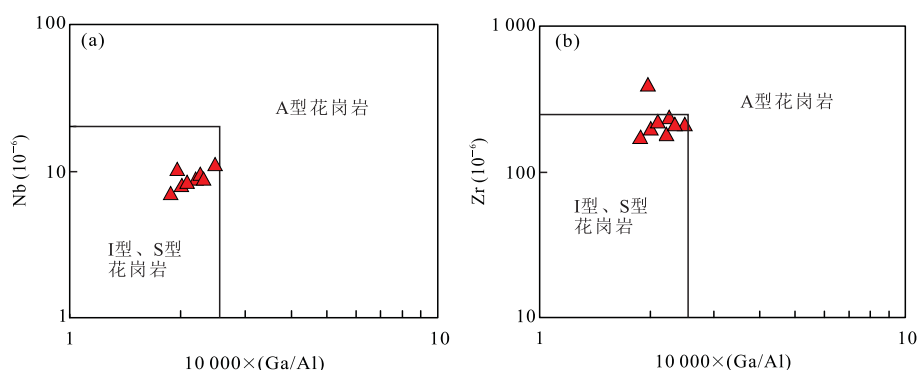


图 5 荔枝沟酸性火山岩 10 000×Ga/Al—Nb (a)和 10 000×Ga/Al—Zr 图解(b)

Fig.5 10 000×Ga/Al vs. Nb (a) and 10 000×Ga/Al vs. Zr (b) diagrams for the Lizhigou felsic volcanic rocks

据 Whalen *et al.*(1987)

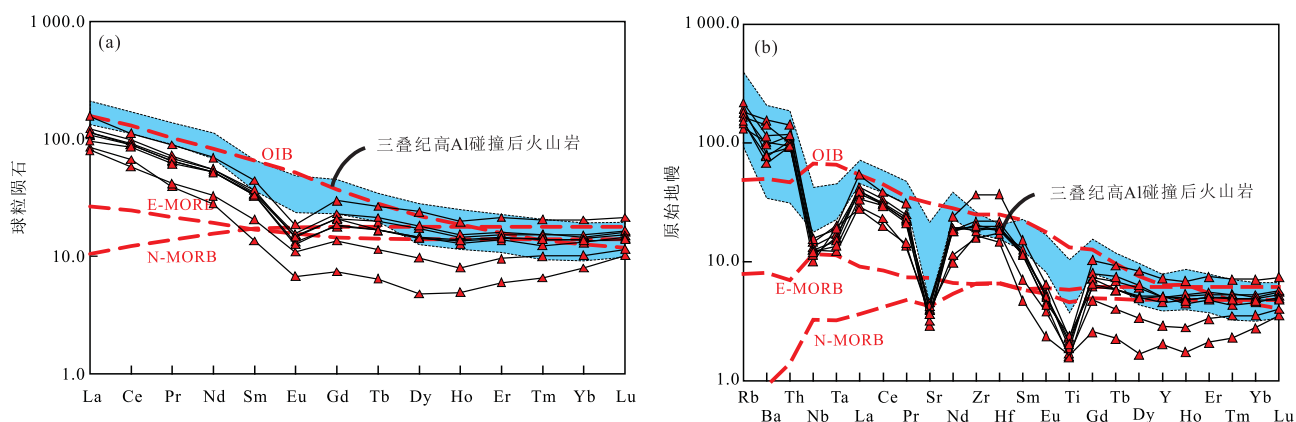


图 6 荔枝沟酸性火山岩球粒陨石标准化稀土元素配分图(a)和标准化微量元素蛛网图(b)

Fig.6 Chondrite-normalized REE pattern(a) and primitive mantle-normalized spidergram (b) for the Lizhigou felsic volcanic rocks

球粒陨石及原始地幔数据来自 Sun and McDonough (1989), 三叠纪高 Al 碰撞后火山岩数据来自 Wang *et al.*(2010)

42.在 TAS 图解中(图 4a),样品皆落于亚碱性流纹岩区域内.样品的铝饱和指数(A/CNK)变化较大为 0.99~2.07,大部分样品 A/CNK>1.1(图 4b).在 A.R.—SiO₂ 图解中(图 4c),样品点都位于钙碱性区

域范围内,因此该套酸性岩浆岩属于钙碱性过铝质酸性火山岩.另外,样品的 10 000×(Ga/Al)比值在 1.87~2.48 的范围内,且大部分样品点落于 I 或 S 型花岗岩的区域内(图 5;Whalen *et al.*,1987),这

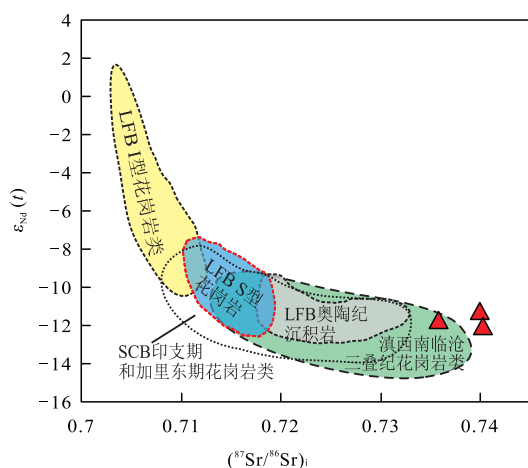


图 7 荔枝沟酸性火山岩 $(^{87}\text{Sr}/^{86}\text{Sr})_i - \epsilon_{\text{Nd}}(t)$ 图解

Fig.7 $(^{87}\text{Sr}/^{86}\text{Sr})_i$ vs. $\epsilon_{\text{Nd}}(t)$ diagram for the Lizhigou felsic volcanic rocks

拉克兰褶皱带(LFB)数据,华南陆块(SCB)印支期、加里东期花岗岩以及滇西南临沧三叠纪花岗岩数据来自 Healy *et al.* (2004); Wang *et al.* (2007, 2013, 2015) 和 Peng *et al.* (2013)

些特征表明所研究样品的性质与 S 型花岗岩接近,同时分异指数 DI ($DI = Q + \text{Or} + \text{Ab} + \text{Ne} + \text{Lc} + \text{Kp}$) 为 $85.5 \sim 92.4$, 表明岩浆经历了高度分异。

荔枝沟酸性火山岩样品稀土元素 ΣREE 变化于 $87 \times 10^{-6} \sim 177 \times 10^{-6}$ 。在球粒陨石标准化图解(图 6a)上,稀土元素配分曲线显示出 LREE 富集的右倾型配分模式,样品轻、重稀土分异明显, $(\text{La}/\text{Yb})_{\text{N}} = 7.05 \sim 10.58$, $(\text{Gd}/\text{Yb})_{\text{N}} = 0.93 \sim 1.58$, Eu 负异常明显 ($\text{Eu}/\text{Eu}^* = 0.49 \sim 0.65$), 其配分特征类似澜沧江晚三叠世高 Al 碰撞后火山岩 (Wang *et al.*, 2010)。在微量元素蛛网图上(图 6b),荔枝沟酸性火山岩样品富集大离子亲石元素(如 Rb、Ba 等),亏损高场强元素,具有 Nb、Ta 和 Ti 的负异常,其配分模式类似于岛弧火山岩。3 个代表性样品均

具较低的全岩 Nd 同位素比值 ($^{143}\text{Nd}/^{144}\text{Nd} = 0.511919 \sim 0.511947$)。根据流纹岩形成年龄 (241.1 Ma) 计算得到的 $(^{87}\text{Sr}/^{86}\text{Sr})_i$ 初始值为 $0.7358 \sim 0.7403$, $\epsilon_{\text{Nd}}(t)$ 值为 $-11.3 \sim -12.1$, Nd 的二阶模式年龄 $T_{\text{DM}2}$ 为 $1.67 \sim 1.73$ Ga。在图 $(^{87}\text{Sr}/^{86}\text{Sr})_i - \epsilon_{\text{Nd}}(t)$ 中(图 7),这些样品均靠近滇西地区临沧三叠纪花岗岩区 (Peng *et al.*, 2013), 同时样品的 $\epsilon_{\text{Nd}}(t)$ 值与金沙江—哀牢山构造带内的晚三叠世酸性岩浆岩类似 ($\epsilon_{\text{Nd}}(t) = -11.4 \sim -10.6$, Liu *et al.*, 2015; $\epsilon_{\text{Nd}}(t) = -10.66 \sim -9.61$, Wang *et al.*, 2014)。

5 讨论

5.1 酸性火山岩岩石成因

所研究的海南三亚荔枝沟地区酸性火山岩烧失量 (LOI) 小于 3% ($0.55\% \sim 2.25\%$), 表明岩浆喷出后未经受强烈的风化和剥蚀。在稀土元素配分图和微量元素蛛网图中,样品具有明显的 Ba、Sr、Ti 和 Eu 负异常特征,说明岩浆在演化过程中可能存在碱性长石、斜长石和黑云母的分离结晶作用。另外,样品点在 La—La/Sm 图解(图 8a)中的趋势显示该套酸性火山岩具有分离结晶作用,在 Sr—Rb 图解(图 8b)和 Sr—Ba 图解(图 8c)中均显示出一定程度的黑云母和斜长石分离结晶的趋势。

如前所述,大部分样品的 A/CNK 值较高 ($A/\text{CNK} > 1.1$), 为过铝质岩石。现有资料表明,富硅的过铝质岩浆岩的形成原因有: (1) 富铝变质岩(如片岩和片麻岩)的部分熔融 (Wang *et al.*, 2007); (2) 角闪岩在富水条件下部分熔融 (Ellis and Thompson, 1986); (3) 贫铝质岩浆的分异 (Zen, 1986)。然而,由 (2) 和 (3) 形成的岩浆岩具有偏铝质、富集 Na

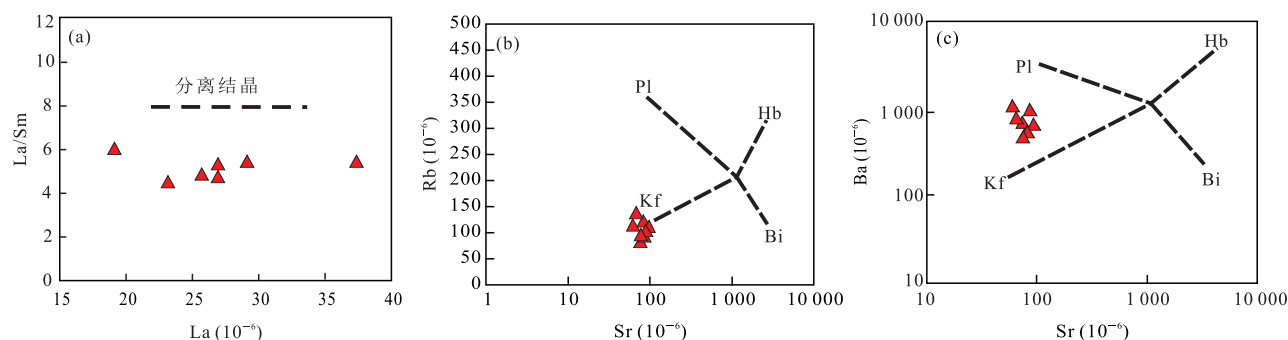


图 8 荔枝沟酸性火山岩 La—La/Sm (a)、Sr—Rb (b) 和 Sr—Ba 图解 (c)

Fig.8 La vs. La/Sm (a), Sr vs. Rb (b) and Sr vs. Ba (c) diagrams for the Lizhigou felsic volcanic rocks

和 Sr 的特征 (Zen, 1986), 这与所研究的荔枝沟酸性火山岩的过铝质和亏损 Sr 的特征不相符。一般情况下, 高级变质岩在富水条件下的深熔作用能产生 2%~5% 的熔体 (Zen, 1986; Patiño Douce *et al.*, 1990), 而这种情况形成的岩浆岩出露范围较小。所研究的样品具有与临沧三叠纪花岗岩相似的地球化学特征, 在 $(^{87}\text{Sr}/^{86}\text{Sr})_i - \epsilon_{\text{Nd}}(t)$ 图中接近滇西南临沧三叠纪 S 型花岗岩区域 (图 7)。已有研究认为, 临沧三叠纪花岗岩源自中上地壳变沉积岩的部分熔融 (Peng *et al.*, 2013), 因此, 推测荔枝沟酸性火山岩可能源于富铝变质岩的部分熔融。

酸性火山岩的成分与其源区密切相关。由于熔体中 $\text{CaO}/\text{Na}_2\text{O}$ 比值的大小主要依赖于斜长石/粘土的比值, 而与温度和压力无关, 因此 $\text{CaO}/\text{Na}_2\text{O}$ 比值能够较好地指示源区的成分 (Sylvester, 1998)。所研究样品的 $\text{CaO}/\text{Na}_2\text{O}$ 比值均大于 0.3, 符合富斜长石贫粘土源区产生的强过铝质花岗岩的 $\text{CaO}/\text{Na}_2\text{O}$ 比值范围 (>0.3), 非贫斜长石富粘土源区产生的强过铝质花岗岩 ($\text{CaO}/\text{Na}_2\text{O}$ 比值 <0.3 ; Patiño Douce and Johnsto, 1991)。同时, 在 $\text{Rb}/\text{Sr}-\text{Rb}/\text{Ba}$ 图解中 (图 9a), 样品全部落于贫粘土源岩的区域; 在 $(\text{Al}_2\text{O}_3 + \text{FeO}^{\text{T}} + \text{MgO} + \text{TiO}_2) - \text{Al}_2\text{O}_3/(\text{FeO}^{\text{T}} + \text{MgO} + \text{TiO}_2)$ 图解中 (图 9b), 部分样品点落于杂砂岩和角闪岩部分熔融的重叠区, 但大多数样品落入杂砂岩部分熔融的区域。该特征结合部分样品具有较高的 $\text{Mg}^{\#}$ (>30) 值, 可以推测荔枝沟酸性火山岩的源区主要为富斜长石贫粘土杂砂岩源区, 同时可能有一定量变火成岩的参与。

温度升高能够使含钛矿物 (如钛铁矿和黑云母) 更易分解, 导致更多的 TiO_2 进入岩浆, 因此 $\text{Al}_2\text{O}_3/\text{TiO}_2$ 比值能够很好地指示强过铝质花岗岩岩浆形成的温度 (Sylvester, 1998)。当 $\text{Al}_2\text{O}_3/\text{TiO}_2$ 比值较高时 (大于 100) 形成温度为低温 ($825\sim 900\text{ }^{\circ}\text{C}$), 当 $\text{Al}_2\text{O}_3/\text{TiO}_2$ 比值较低时 (小于 100) 形成温度为高温 ($900\sim 950\text{ }^{\circ}\text{C}$) (Sylvester, 1998)。荔枝沟酸性火山岩样品的 $\text{Al}_2\text{O}_3/\text{TiO}_2$ 比值均小于 100, 反映其部分熔融的温度为高温 ($>875\text{ }^{\circ}\text{C}$), 属高温型强过铝质酸性火山岩。高压型和高温型强过铝质花岗岩的主要区别在于前者形成的温度较低 ($<875\text{ }^{\circ}\text{C}$), 其诱发地壳物质熔融的热动力因素是地壳增厚, 而後者的形成温度较高 ($>875\text{ }^{\circ}\text{C}$), 是通过玄武质岩浆的底侵作用导致地壳物质重熔形成的 (钟长汀等, 2007)。综上所述, 荔枝沟酸性火山岩是中上地壳的变沉积岩受到玄武质岩浆底侵作用而发生部分熔融、后经历一定程度的分离结晶作用的产物。

5.2 构造意义

荔枝沟酸性岩浆岩为钙碱性岩, 具有富集大离子亲石元素、亏损高场强元素的特征。在构造环境判别图中 (图 10), 本次研究的样品点落于碰撞后火山岩的区域范围, 显示荔枝沟酸性火山岩形成于碰撞后的构造环境。除本文报道的海南三亚地区 241 Ma 酸性岩浆岩之外, 海南岛还分布有同一时期的中基性岩浆岩, 比如: 在三亚地区的石榴霓辉石正长岩 (244 Ma; 谢才富等, 2005); 万宁地区的辉长岩和辉绿岩脉 (240 Ma; 唐立梅, 2010); 分界洲地区的 A 型正长岩类 (231 Ma; 唐立梅, 2010); 陵水—龙滚断

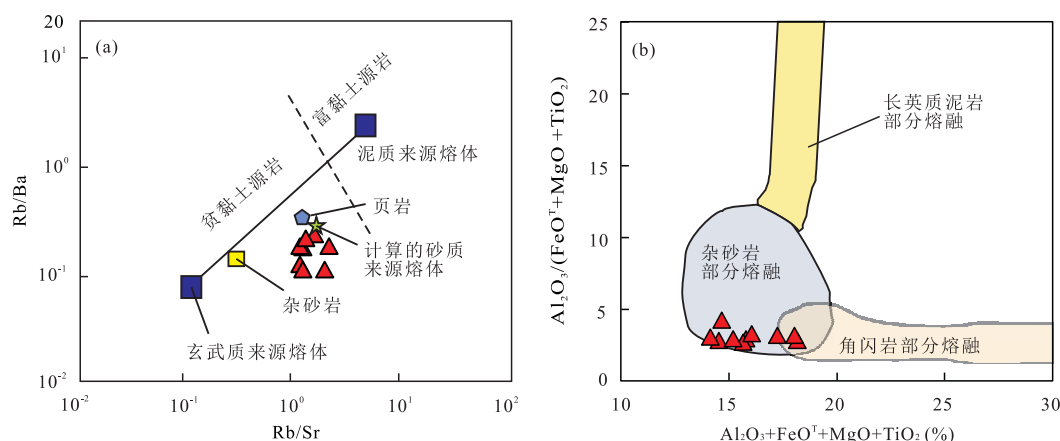


图 9 荔枝沟酸性火山岩 $\text{Rb}/\text{Sr}-\text{Rb}/\text{Ba}$ (a) 和 $(\text{Al}_2\text{O}_3 + \text{FeO}^{\text{T}} + \text{MgO} + \text{TiO}_2) - \text{Al}_2\text{O}_3/(\text{FeO}^{\text{T}} + \text{MgO} + \text{TiO}_2)$ 图解 (b)
Fig.9 Rb/Sr vs. Rb/Ba (a) and $(\text{Al}_2\text{O}_3 + \text{FeO}^{\text{T}} + \text{MgO} + \text{TiO}_2)$ vs. $\text{Al}_2\text{O}_3/(\text{FeO}^{\text{T}} + \text{MgO} + \text{TiO}_2)$ (b) diagrams for the Lizhigou felsic volcanic rocks

图 a 据 Sylvester (1998); 图 b 据 Wang *et al.* (2016)

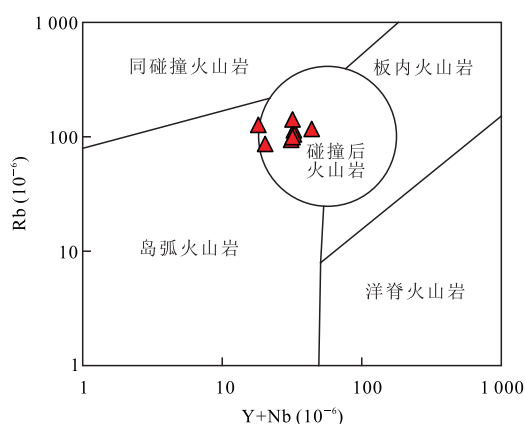


图 10 荔枝沟酸性火山岩 Y+Nb-Rb 构造判别图解

Fig.10 Y+Nb vs. Rb discrimination diagram for the Lizhigou felsic volcanic rocks

据 Pearce (1996)

裂两侧及白沙断裂东南侧中三叠世的正长岩和花岗岩组合(周佐民等, 2011). 这些关于碱性岩、辉长岩和辉绿岩的研究表明海南岛在 240 Ma 左右已进入后造山伸展阶段. 根据野外观察发现, 荔枝沟酸性火山岩被上覆的上三叠统砾岩角度不整合覆盖, 暗示了华南陆块和印支陆块之间的碰撞在晚三叠世之前已经结束(Jian *et al.*, 2009; Fan *et al.*, 2010; Liu *et al.*, 2014).

越来越多的研究表明, 在晚古生代至早中生代时期, 海南岛的构造演化与古特提斯洋的闭合及印支、华南陆块的拼合有着密切的联系. Li *et al.* (2002) 和 He *et al.* (2017) 对邦溪—晨星地区石炭纪变基性岩与金沙江—哀牢山—松马缝合带内的岩浆记录进行对比, 认为沿金沙江—哀牢山—松马带和邦溪—晨溪带发育了弧后盆地. Li *et al.* (2006) 和陈新跃等(2011) 在五指山地区获得了 267~262 Ma 与俯冲相关的 I 型花岗片麻岩. 唐立梅等(2013) 认为 238~234 Ma 的兴隆双峰式侵入岩的成因与古特提斯洋的俯冲闭合有关. 相应地沿金沙江—哀牢山—松马构造带分布有大量的晚二叠世至三叠纪的岩浆岩. Zi *et al.* (2012) 在白马雪山识别出了 250 Ma 的闪长岩和英云闪长岩组合. Liu *et al.* (2015) 在哀牢山构造带发现了 247 Ma 的狗头坡和通天阁淡色花岗岩, 其中通天阁花岗岩的 $\epsilon_{\text{Nd}}(t)$ 值为 -11.4~-10.6, 与所研究的荔枝沟酸性火山岩的 $\epsilon_{\text{Nd}}(t)$ 值接近. Liu *et al.* (2017) 还在下关地区识别出了形成于 233~229 Ma 碰撞后的强过铝质 S 型花岗质岩, 该套花岗质岩具有富集的 Sr-Nd 同位素组成, $\epsilon_{\text{Nd}}(t)$ 小于 -10. 最近 Wang *et al.* (2018) 研究表明,

金沙江—哀牢山—松马分支洋/弧后盆地的同碰撞和后碰撞时间分别为 247 Ma 和 237 Ma. 以上认识表明在中—晚三叠世时期, 金沙江—哀牢山—松马构造带已处于同碰撞—后碰撞的构造背景. 因此海南岛在二叠纪末—中三叠世的构造演化过程可与古特提斯构造域在滇西的演化进行对比. 除了岩浆记录外, 研究者们也对海南岛和印支陆块东北部地区的韧性剪切带进行了对比研究, 并在琼中和琼西公爱地区发现了 NW 向右旋走滑韧性剪切带, 其糜棱岩同构造期云母的 ^{40}Ar - ^{39}Ar 坪年龄为 242~245 Ma, 而在印支陆块北部松马、岷港—溪山和 Song Ca-Rao Nay 等地发现的韧性剪切带, 其构造特征和变形年龄 (254~240 Ma) 与前者一致 (Lepvrier *et al.*, 2004; 陈新跃等, 2006; Zhang *et al.*, 2011). 野外观察表明荔枝沟酸性火山岩被上覆的上三叠统砾岩角度不整合所覆盖, 而研究者在金沙江—哀牢山地区发现中至上三叠统的火山岩序列也被上三叠统一碗水组砾岩不整合覆盖 (Peng *et al.*, 2006; Zi *et al.*, 2012; Wang *et al.*, 2018). 荔枝沟地区酸性火山岩的地化特征也表明其源区为变沉积物, 具有碰撞后岩浆作用的主要特征, 其形成年龄 (241.1 Ma) 符合三江地区一般认为的碰撞后的岩浆作用时间, 与金沙江—哀牢山—松马构造带内碰撞后岩浆作用在时空上相吻合 (Liu *et al.*, 2014). 结合海南岛二叠纪俯冲相关的 I 型花岗岩, 现有的资料暗示了华南和印支陆块之间的俯冲于早二叠世已开始, 并在 240 Ma 左右进入到后碰撞环境. 此时期华南陆块和印支陆块碰撞后发生板片拆沉、软流圈上涌, 导致上地壳物质发生部分熔融, 从而形成了荔枝沟酸性火山岩 (Liu *et al.*, 2014, 2017; Wang *et al.*, 2018).

6 结论

(1) 海南岛荔枝沟地区酸性火山岩 LA-ICP-MS 锆石 U-Pb 年龄为 241.1 ± 5.8 Ma, 为中三叠世;

(2) 荔枝沟地区酸性火山岩具有高 Al_2O_3 和全碱含量、富集轻稀土元素和大离子亲石元素、亏损 Nb、Ta 和 Ti 等高场强元素等特征, 为变沉积岩部分熔融并经历一定程度的分离结晶的产物;

(3) 荔枝沟酸性火山岩形成于碰撞闭合后的伸展环境, 与碰撞后板片拆沉、软流圈上涌有关, 并可以与金沙江—哀牢山—松马带相关的中—晚三叠世岩浆岩进行对比.

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