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中国古近纪构造—地层区划及地层格架

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摘要: 古近纪是地球演化进程中最为关键的时段之一, 如印度与欧亚板块碰撞和青藏高原初始隆升、太平洋板块向西往欧亚大陆俯冲伴随大规模弧后裂陷、古新世—始新世极热和渐新世初大降温、哺乳动物和被子植物大辐射等, 因此, 建立中国古近纪构造—地层区划及地层对比格架, 是进一步深入研究古近纪时期地球表层岩石圈—水圈—大气圈—生物圈等多圈层耦合演化的基石, 意义十分重大. 中国古近纪构造—地层区划主要反映古近纪阶段中国范围内的构造—地层的一级和二级单元的划分与各单元内的主要盆地构造背景与分布特征. 本次通过对中国古近纪大地构造、沉积盆地(群)、火山岩、隆起带和大型变形构造等的综合研究, 共厘定出一级构造—地层单元(大区)10个, 二级构造—地层单元27个(区), 包含各类沉积盆地129个. 通过对各个地层分区内盆地的盆地类型、构造控制背景、沉积序列及地层接触关系的研究, 建立了中国古近纪地层对比格架, 并对中国古近纪时期沉积—构造耦合历史进行了系统总结.

关键词: 古近纪; 构造—地层区划; 地层对比; 沉积盆地; 中国.

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Paleogene Tectonic-Stratigraphic Realms and Sedimentary Sequence in China

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Abstract: Paleogene is one of the key intervals in the earth history and a critical period of many biotic and geologic events, including the collision between Indian and Eurasian plates as the resultant formation of the Tibetan plateau, the westward subduction of the Pacific plate beneath Eurasian plate and the accompanying large scale back-arc rifting, the Paleocene-Eocene thermal maximum and Eocene-Oligocene climate transition, and rapid radiation of mammals and angiosperms, etc. Thus, it is of great significance to establish the Paleogene tectonic-stratigraphic divisions and stratigraphic correlation framework in China, which is the cornerstone of further study on the coupling evolution of lithosphere, hydrosphere, atmosphere and biosphere of the earth in the Paleogene. Paleogene tectonic-stratigraphic divisions in China mainly reflects the division of the first and second-order tectonic-stratigraphic units and the tectonic background and distribution characteristics of the main basins in each unit. Through the comprehensive study of Paleogene tectonics, sedimentary basins (groups), volcanic rocks, uplifting belts and large deformation structures in China, 10 first-order tectonic-stratigraphic units (super-realms) and 27 second-order tectonic-stratigraphic units (realms) are devised, including 129 basins. Through the research of the types of basins in each unit and their characteristics of tectonic settings, sedimentary sequence and stratigraphic contact, the Paleogene stratigraphic correlation framework in China and

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summarized the sedimentary-tectonic history is established.

Key words: Paleogene; tectonic-stratigraphic division; stratigraphic correlation; sedimentary basin; China.

古近纪是新生代的第一个纪,也是地质历史时期最为关键的时段之一(Scotese, 2014; Wang *et al.*, 2019). 古近纪期间印度板块与欧亚板块碰撞及其后青藏高原的隆升(Molnar and Tapponnier, 1975; Yin and Harrison, 2000),以及太平洋板块向西往欧亚大陆之下俯冲(Northrup *et al.*, 1995; 潘桂棠等, 2016)共同奠定了东亚现今的构造地貌格局. 此外,古近纪期间发生了多次重要的全球性生物和气候事件,生物事件包括白垩纪—古新世之交的全球性生物大灭绝、早始新世哺乳动物类的辐射演化等(图 1; Feduccia, 1995); 古气候事件包括古新世—始新世极热事件(PETM)、早始新世大暖期(EECO)、中始新世大暖期(MECO)和始新世—渐新世之交气候转变事件(EOCT)等(图 1; Zachos *et al.*, 2001),其中 EOCT 时期地球经历了新生代首次的大规模快速降温和南极大陆冰盖的快速扩张. 因此,建立中国古近纪构造—地层区划及地层对比格架,是进一步深入研究古近纪时期地球表层岩石圈—水圈—大气圈—生物圈等多圈层耦合演化的基石,意义十分重大.

自 20 世纪初李四光和尹赞勋等编制中国区域地层表的工作开始,众多学者对包括新生代在内的中国各个时代的地层区划与地层对比开展了大量卓有成效的工作(汪啸风和陈孝红, 2005; 张克信等, 2015). 李云通(1984)、叶得泉等(1993)和郑家坚等(1999)专门针对中国古近纪—新近纪(先前统称为第三纪)地层进行了详细的地层区划和地层对比工作,为之后中国古近纪—新近纪地层区划和地层对比研究奠定了重要基础. 近 20 多年来,随着全国各省(市、自治区)新一轮的 1:25 万和 1:5 万地质填图成果的提交,以及涉及中国新生代沉积大地构造区划与演化专题研究相关论著与论文的陆续发表,特别是全国新一轮各省区域地质志的陆续更新出版,在古近纪地层沉积方面积累了大量的宝贵资料(张克信等, 2013, 2017). 因此,非常有必要对中国古近纪各类地层资料进行系统梳理,进而建立新的中国古近纪构造—地层区划和地层对比格架.

1 区划原则

本文对中国古近纪构造地层区划遵循我们先

前订立的地层区划原则(Zhang *et al.*, 2010; 张克信等, 2015, 2017), 主要是: (1) 区内所包括的盆地具有大体一致的构造演化背景和形成机制; (2) 区内所包括的盆地具有大体一致的构造—地貌景观; (3) 区内所包括的盆地具有相对一致的沉积充填序列和大体相近的沉积环境演化过程; (4) 区内各盆地具有大体一致的古地理格局和古环境条件; (5) 地层区划的边界识别标志需明确,如构造—地层大区的边界通常以对接带、构造大相、主要断裂、地层类型、生物大区等边界为其边界; 而地层区的边界通常以结合带、重要断裂、构造相、地层类型、生物区等的边界为其边界. 以上述划分原则为指导,在前人对中国古近纪地层区划研究工作的基础上,系统查阅全国各省(市、自治区)1:25 万和 1:5 万地质填图以及全国最新版区域地质志中新生代地层资料,并结合近年来前人已发表成果的基础上,划分出中国古近纪各类盆地 129 个(图 2), 归属为西北压陷区(1)、羌塘—三江压陷—走滑构造区(2)、冈底斯陆缘弧(3)、喜马拉雅汇聚挤压构造区(4)、秦祁昆挤压变形区(5)、华北伸展断陷区(6)、东北裂陷区(7)、华南隆起区(8)、中国东部陆缘裂谷区(9)和东南亚多岛弧盆系(10)共 10 个构造—地层大区,进一步细分为 27 个构造—地层区(图 2).

2 各构造—地层大区特征简介

2.1 西北压陷区

西北压陷区根据盆地类型和盆地沉积充填序列特征,划分为阿尔泰隆起区、准噶尔—吐哈压陷区、天山—北山逆冲区、塔里木压陷区和阿尔金—阿拉善隆起区(图 2、图 3).

阿尔泰隆起区古近纪期间主体为剥蚀区,没有沉积记录.

准噶尔—吐哈压陷区在中生代系列断陷—坳陷盆地的基础上,在挤压背景之下古近纪时期以发育压陷盆地为主,包含布尔津盆地、塔城盆地、伊宁盆地、准噶尔盆地、三塘湖盆地和吐哈盆地. 以准噶尔盆地为例,该盆地新生代受到欧亚板块和印度板块的碰撞远程效应影响,盆地南北两侧冲断活动强烈,准噶尔盆地全面发育新生代沉积序列. 古近纪,准噶尔盆地南部整体表现为较稳定的沉积期,为均

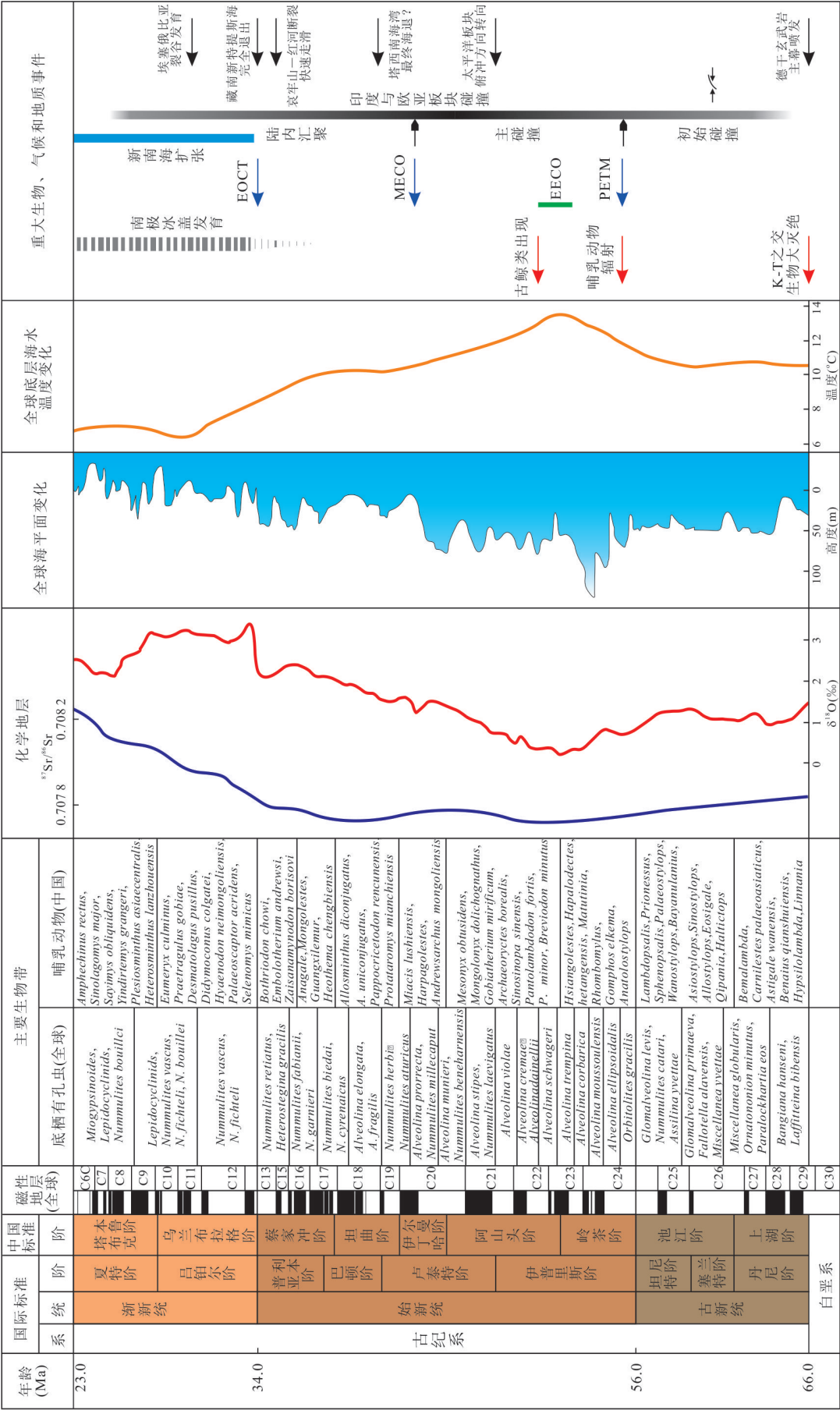


图 1 中国古近纪年代—磁性—生物—化学—气候等综合地层格架与全球地质事件序列
Fig.1 Paleogene multidisciplinary stratigraphic correlation in China and global geo-events
磁性地层据 Gradstein *et al.* (2012); 主要生物带据 Vandenbergh *et al.* (2012); Wang *et al.* (2019); 化学地层据 Zachos *et al.* (2001); McArthur *et al.* (2012); 全球海平面变化据 Miller *et al.* (2005); 全球底层海水温度变化据 Lear *et al.* (2000); 重大生物、气候和地质事件据 Feduccia (1995); Zachos *et al.* (2001); Bosboom *et al.* (2014); Li *et al.* (2014); Wang *et al.* (2014); Ding *et al.* (2017); Hu *et al.* (2017); 王二七 (2017); Wang and Jian (2019)

衡沉降区,以紫泥泉子组、安集海河组和沙湾组下部沉积为代表(图 3;李云通,1984;Ji *et al.*,2008);盆地北部古近纪地层序列由老到新包括乌伦古河组和铁尔斯哈巴合组(图 3;叶捷等,2001)。

天山—北山逆冲区包含焉耆盆地,该盆地在中生代残留盆地基础上,古近纪时期整体处于天山南北两侧对冲所造成的双向挤压构造格局之中,最终压缩定型(袁正文,2003),发育了一套以鄯善群为代表的粗碎屑沉积。

塔里木压陷区包含塔西南海湾和塔里木盆地,均为挤压背景下的构造挠曲压陷盆地。古近纪时期塔西南海湾为副特提斯海的一部分,以半封闭的海湾滨浅海沉积为主(图 2),从老至新岩石地层单位依次是阿尔塔什组、齐姆根组、卡拉塔尔组、乌拉根组、巴什布拉克组和克孜洛依组下段(图 3),各组间均呈整合接触(郑家坚等,1999;Bosboom *et al.*,2014)。受塔中巴楚隆起的隔断,塔里木北部地区包括库车坳陷及塔北隆起,塔北地区形成了另一套沉积序列,古近纪地层序列由老到新分别为库姆格列木群、苏维依组和吉迪克组下部(图 3、图 4;贾承造等,2004)。库车坳陷新生代地层的沉积时代一直存在较大的争议,最近 Zhang *et al.*(2014)通过高精度磁性地层学研究对库车坳陷的新生代地层沉积时代进行了重新厘定,将库姆格列木群、苏维依组和吉迪克组的沉积时代分别限定为 65~38 Ma,38~36 Ma 和 36~13 Ma(图 4)。

阿尔金—阿拉善隆起区主体为剥蚀区,仅发育索尔库里盆地、花海—金塔盆地和雅布赖盆地 3 个小型山间盆地。以索尔库里盆地为例,该盆地位于阿尔金山中段,为阿尔金走滑断裂控制下的拉分盆地,其古近纪沉积序列与其东侧的柴达木盆地一致(陈正乐等,2003)。

2.2 羌塘—三江压陷—走滑构造区

羌塘—三江压陷—走滑构造区根据盆地类型和盆地沉积充填序列特征,划分为羌塘压陷区和三江压陷—走滑拉分区(图 2、图 3)。

羌塘压陷区古近纪盆地包含黑石北湖盆地、那底岗日盆地、木孜塔格盆地、西可可西里盆地、松西盆地、鲁玛江东措盆地、拜惹布措盆地、温泉湖盆地、查多岗日盆地、强措盆地、石坪顶盆地、玛尔果茶卡盆地、鲸鱼湖盆地、多格措仁盆地、可可西里盆地、沱沱河盆地、米提将古木错盆地、唐古拉盆地、莫云盆地、日土盆地、扎普盆地、丁则盆地、改则盆

地、洞措盆地、北双湖盆地、尼玛盆地、南双湖盆地、伦坡拉盆地、索县盆地和类乌齐北盆地。以伦坡拉盆地和可可西里—沱沱河盆地为例细述如下。伦坡拉盆地是在燕山期褶皱基底上发育的新生代陆相断拗盆地,其形成、演化与班公湖—怒江断裂的活动密切相关,沿班公湖—怒江断裂东西向展布(图 2)。伦坡拉盆地古近纪沉积厚度巨大,生烃地质条件良好,由老到新分别为牛堡组和丁青湖组(图 3;赵政璋等,2001)。其中牛堡组发育在盆地初始断陷时期,而丁青湖组主要发育在盆地拗陷期(马鹏飞等,2013)。对于伦坡拉盆地古近纪地层的沉积时代,先前主要基于介形类等微体化石来进行限定,近年来随着哺乳动物化石、火山熔岩测年等研究手段的开展,将该盆地丁青湖组中下段至上段的沉积时代限定为晚渐新世—早中新世(Deng *et al.*,2012;Sun *et al.*,2014)。可可西里—沱沱河盆地为受近东西向断裂影响的压陷盆地,位于巴颜喀拉地块西段和羌塘地块的北部,北部以昆仑南缘缝合带为界,南部以唐古拉断层为界(Wang *et al.*,2008),其中可可西里盆地和沱沱河盆地为该盆地中的两个主要次级盆地。可可西里—沱沱河盆地内古近纪地层序列由老到新分别为风火山群和雅西错组(图 3、4;Wang *et al.*,2008)。可可西里盆地古近纪地层的沉积时代目前仍存在较大争议(Liu *et al.*,2003;Staisch *et al.*,2014;Jin *et al.*,2018),而我们最新的微体古生物学研究显示雅西错组的沉积时限包括晚始新世—早渐新世(宋博文等,2019)。

三江压陷—走滑拉分区古近纪盆地的形成演化与印度—欧亚板块自初始碰撞以来的陆内变形、地壳缩短、逆冲挤压—走滑等构造响应密切相关(Zhang *et al.*,2010;张克信等,2013;周江羽和王江海,2019),表现为发育一系列中小型的古近纪压陷—走滑拉分盆地和断陷盆地,包括巴干盆地、曲麻莱盆地、杂多盆地、玉树盆地、石渠盆地、囊谦—下拉秀盆地、甘孜—理塘盆地、白玉盆地、贡觉盆地、丁青盆地、八宿盆地、左贡盆地、芒康盆地、乡城盆地、兰坪—剑川盆地、丽江盆地、木里西盆地、盐源盆地、西昌—攀枝花盆地、楚雄盆地、大理盆地、昌宁盆地、保山盆地、耿马盆地、临沧东盆地、思茅盆地、昆明盆地、马关盆地和百色盆地。以囊谦—下拉秀盆地和兰坪—剑川盆地为例细述如下。

囊谦—下拉秀盆地是在印度与欧亚板块碰撞的构造背景下,在中生代基底经历长期隆升剥蚀的

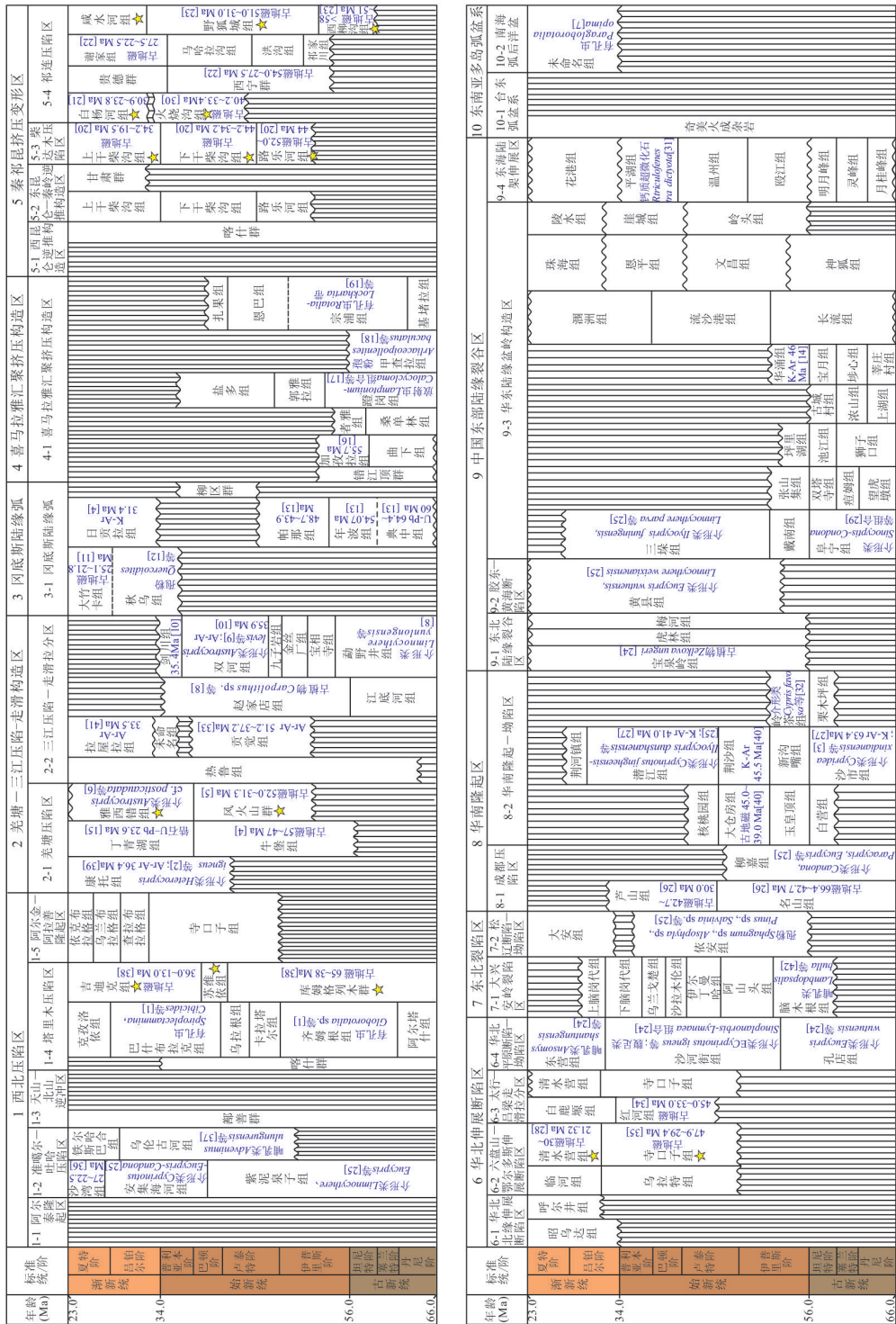


图3 中国古近纪构造—地层区划与地层格架对比

Fig.3 Paleogene tectonic-stratigraphic realms and correlation of lithostratigraphic units in China

图中数值表示: [1]郗家坚等,1999;[2]姜高磊等,2014;[3]张守安,1989;[4]赵政璋等,2001;[5]Wang *et al.*,2008;[6]宋博文等,2019;[7]Wang *et al.*,2003;[8]云南省地质矿产局,1996; [9]Wu *et al.*,2018;[10]Gourbet *et al.*,2017; [11]Ai *et al.*,2019;[12]李建国,2004;[13]Zhou *et al.*,2004;[14]董月霞等,2006;[15]Sun *et al.*,2014;[16]Hu *et al.*,2016;[17]Liang *et al.*,2012;[18]Li *et al.*,2005;[19]万晓樵,1987;[20]Ji *et al.*,2017;[21]宋春晖,2006;[22]Fang *et al.*,2019;[23]Yue *et al.*,2001;[24]汪啸风和陈孝红,2005;[25]李云通,1984;[26]申琪,2018;[27]徐论勋等,1995; [28]Liu *et al.*,2020; [29]江苏省地质矿产局,1997;[30]Dai *et al.*,2005;[31]王建平,1992;[32]张显球和李茜,2010;[33]Spurim *et al.*,2005;[34]鹿化煜等,2018;[35]房建军等,2008;[36]Ji *et al.*,2008;[37]叶捷等,2001;[38]Zhang *et al.*,2014;[39]钟华明等,2008;[40]马小林,2012;[41]Zhang *et al.*,2004; [42]Wang and Jian,2019

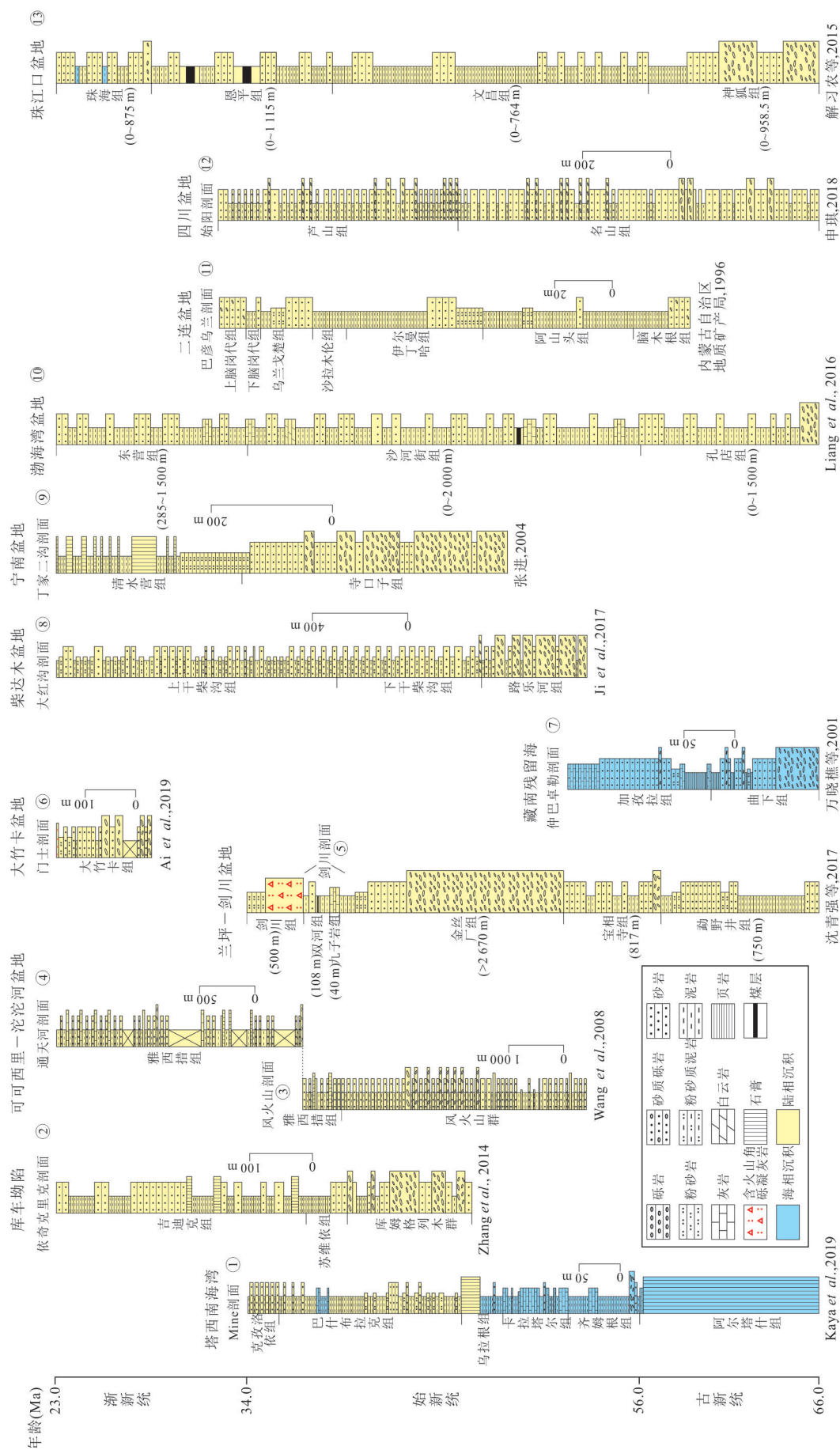


图4 中国古近纪构造—地层区代表性盆地沉积充填序列

Fig.4 Sedimentary filling sequence of representative basins in Paleogene tectonic-stratigraphic realms of China

于青藏高原东南缘,沿哀牢山—红河断裂呈 NNW—SSE 向展布,先前对于该盆地新生代特别是古近纪地层的沉积时代多有争议,近年来学者们基于盆地构造、古生物学和盆内新生代火山岩等的综合研究,对兰坪—剑川盆地新生代地层的沉积时代进行了重新厘定:古近纪地层序列由老到新依次为勐野井组、宝相寺组、金丝厂组、九子岩组、双河组和剑川组(图 3、图 4; Gourbet *et al.*, 2017; 沈青强等, 2017; Wu *et al.*, 2018)。

2.3 冈底斯陆缘弧

冈底斯陆缘弧位于印度河—雅鲁藏布江结合带的北侧,班公湖—双湖—怒江对接带的南侧。古新世—始新世期间,冈底斯陆缘弧主要由巨厚的钙碱性—高钾钙碱性林子宗群火山岩地层构成(Mo *et al.*, 2008),从下向上分为典中组、年波组和帕那组,典中组最底部火山岩的 Ar/Ar 年龄为 64.43 Ma (Zhou *et al.*, 2004),帕那组时代介于 48.7~43.9 Ma (Zhou *et al.*, 2004)。林子宗群与下伏的白垩纪地层之间呈角度不整合接触关系,说明印度与欧亚板块沿雅鲁藏布江带的初始碰撞发生在白垩纪与古新世之交,主碰撞结束的时间约为 44 Ma。在该陆缘弧及其两侧发育古近纪沉积盆地,包含措勤盆地、尼玛盆地、申扎盆地、班戈南盆地、当雄盆地和大竹卡盆地(图 2、图 3)。以大竹卡盆地为例细述之。渐新世晚期,随着印度板块高角度向欧亚板块之下俯冲,导致上地壳拉张,在拉张背景之下沿雅鲁藏布缝合带形成了系列断陷盆地(Wang *et al.*, 2013),沿雅鲁藏布江缝合带呈东西向带状分布,统称为大竹卡盆地。大竹卡盆地古近纪地层即为大竹卡组(图 3、图 4; Aitchison *et al.*, 2009; Ai *et al.*, 2019)。同时在冈底斯陆缘弧南侧日喀则—拉孜一带可见柳区群(柳区砾岩)发育并沿雅江缝合带呈东西向延展,而目前学界对于柳区群(柳区砾岩)具体的沉积时代依然存在争议(方爱民等, 2005; Hu *et al.*, 2016)。

2.4 喜马拉雅汇聚挤压构造区

古新世—始新世,喜马拉雅汇聚挤压构造区具有明显的沉积相带分异,从北向南,海水由浅海到深海再到浅海(Zhang *et al.*, 2010; 张克信等, 2013)。北带的海陆过渡—浅海分布带很窄,也较局限,见于仲巴和巴噶地区,称错江顶群,包括下部的曲下组 and 上部的加孜拉组(图 3; 万晓樵等, 2001; Zhang *et al.*, 2010)。南带的浅海分布带宽,分布区域较大,见于岗巴—定日一带,地层层序自下而上为基堵拉

组、宗浦组、恩巴组和扎果组(图 3; 万晓樵, 1987; Hu *et al.*, 2016)。从东向西,半深海—深海沉积沿江孜—萨嘎—郭雅拉—桑麦一线分布。中带东部的江孜发育以甲查拉组为代表的半深海斜坡扇沉积建造(图 3; Li *et al.*, 2005),中带西部为桑单林组、者雅组、蹬岗组、郭雅拉组和盐多组为代表的半深海—深海含放射虫硅质岩和混杂岩建造(图 3; Zhang *et al.*, 2010; Liang *et al.*, 2012; Hu *et al.*, 2016)。

2.5 秦祁昆挤压变形区

秦祁昆挤压变形区根据盆地类型和盆地沉积充填序列特征,划分为西昆仑逆推构造区、东昆仑—秦岭逆推构造区、柴达木压陷区和祁连压陷区(图 2、图 3)。

西昆仑逆推构造区古近纪主体为剥蚀区,仅在甜水海地区有小型山间盆地发育,即甜水海盆地。

东昆仑—秦岭逆推构造区包含秀河盆地、玛多盆地和玛曲盆地。

柴达木压陷区主要包含柴达木盆地。柴达木盆地位于青藏高原北缘,新生代自早始新世以来一直处于南祁连山、阿尔金山和东昆仑山三大构造带的相互作用形成的挤压—走滑的区域应力环境中,其根本动力来自南侧印度板块向北的俯冲挤压。柴达木盆地古近纪的沉积演化直接受盆地周缘三大主控断裂带即阿尔金断裂、东昆仑断裂及南祁连断裂带的控制,盆—山耦合十分显著(Yin *et al.*, 2008)。柴达木盆地古近纪地层序列由老至新依次为路乐河组、下干柴沟组 and 上干柴沟组中下段(图 3、图 4; Lu and Xiong, 2009; Ji *et al.*, 2017)。此外,必须指出的是目前柴达木盆地新生代磁性地层学年龄约束存在较大争议(Wang *et al.*, 2017; Nie *et al.*, 2020),本文中我们仍然遵循柴达木盆地传统地层年龄认识(宋春晖, 2006; Zhang *et al.*, 2010)。

祁连压陷区主要包含阿克塞盆地、疏勒南山盆地、酒泉盆地、花海—金塔盆地、武威盆地、西宁—兰州盆地、临夏—定西盆地、合作盆地、天水盆地 and 宁南盆地。以酒泉盆地和西宁—兰州盆地为例细述如下。

酒泉盆地位于青藏高原北缘的河西走廊西段,其新生代沉积演化受阿尔金走滑断裂和北祁连断裂带控制,是在古生代褶皱基底上发育的新生代压陷盆地(宋春晖, 2006)。该盆地古近纪地层序列由老到新依次为火烧沟组和白杨河组(图 3; Dai *et al.*, 2005; 宋春晖, 2006)。古近纪期间西宁盆地和兰州盆

地应为相互连通的两个次级盆地,统称为西宁—兰州盆地(图 2)。其中西宁盆地位于青藏高原东北缘中祁连造山带、东昆仑造山带和西秦岭构造带的交汇部位,其北部和南部分别为达坂山和拉脊山所围限(Zhang *et al.*, 2016)。该盆地古近纪地层序列由老到新依次为祁家川组、洪沟组、马哈拉沟组和谢家组(图 3; Fang *et al.*, 2019)。兰州盆地是西宁—兰州盆地的次级盆地之一,位于青藏高原东北缘祁连造山带东段。该盆地古近纪地层序列由老到新依次为西柳沟组、野狐城组和咸水河组下段(图 3; Yue *et al.*, 2001; 张克信等, 2013)。兰州盆地古近纪地层的磁性时代同样依然存在争议(Zhang *et al.*, 2014; Wang *et al.*, 2016)。

2.6 华北伸展断陷区

华北伸展断陷区根据盆地类型和盆地沉积充填序列特征(图 3),划分为华北北缘伸展断陷区、六盘山—鄂尔多斯伸展断陷区、太行—吕梁走滑—拉分区和华北平原断陷—坳陷区(图 2、图 3)。

华北北缘伸展断陷区古近纪期间主体为剥蚀区。

六盘山—鄂尔多斯伸展断陷区包含河套盆地和银川盆地。以银川盆地为例细述之。银川盆地位于鄂尔多斯地块与阿拉善地块之间,其南缘与青藏高原东北缘弧形构造带相接,是新生代以来在其主边界断裂控制之下发育的断陷盆地(Shi *et al.*, 2020)。银川盆地古近纪地层序列与宁南盆地一致,由老到新依次为寺口子组和清水营组(图 3; Liu *et al.*, 2020)。关于寺口子组的具体沉积时代,由于未见有哺乳动物化石发现,目前尚存在较大争议(Jiang *et al.*, 2007; 房建军等, 2008; Wang *et al.*, 2011),本文中依然沿用传统年龄认识,其沉积时代为中始新世—早渐新世。

太行—吕梁走滑—拉分区古近纪盆地主要为渭河盆地。渭河盆地位于鄂尔多斯地块南缘与秦岭造山带的交接部位,总体呈 EW 向展布(图 2),为一受其南北主控正断层控制之下形成的新生代断陷盆地(Shi *et al.*, 2020)。渭河盆地古近纪地层序列由老到新依次为红河组和白鹿塬组(图 3; 鹿化煜等, 2018)。

华北平原断陷—坳陷区包含周口盆地、渤海湾盆地和合肥盆地。以渤海湾盆地为例细述之。渤海湾盆地位于华北克拉通东部,是在中国东部环太平洋构造域岩石圈拉伸变薄的背景下形成的断陷—

坳陷型新生代裂谷盆地(许淑梅等, 2014)。渤海湾盆地古近纪为其主要裂陷期(Allen *et al.*, 1997),古近纪地层序列由老到新依次为孔店组、沙河街组和东营组(图 3、图 4; Liang *et al.*, 2016)。

2.7 东北裂陷区

东北裂陷区根据盆地类型和盆地沉积充填序列特征(图 3),划分为大兴安岭裂陷区和松辽断陷—坳陷区(图 2、3)。

大兴安岭裂陷区包含二连盆地。二连盆地是在内蒙古—大兴安岭造山带基底上和燕山晚期拉张构造应力场作用下发育起来的中新生代断陷—坳陷盆地(李洪军等, 2012)。二连盆地古近纪地层序列由老到新依次为脑木根组、阿山头组、伊尔丁曼哈组、沙拉木伦组、乌兰戈楚组、下脑岗代组和上脑岗代组(图 3、图 4; 内蒙古自治区地质矿产局, 1996; Wang *et al.*, 2019)。

松辽断陷—坳陷区包含松辽盆地。松辽盆地处于古亚洲洋构造域与古太平洋—太平洋构造域的复合交切部位,为晚中生代以来在拉张背景下的发育的大型陆内断陷—坳陷盆地(Ren *et al.*, 2002),该盆地古近纪地层主要为依安组和大安组下段(图 3; 李云通, 1984)。

2.8 华南隆起区

华南隆起区根据盆地类型和盆地沉积充填序列特征(图 3),划分为成都压陷区和华南隆起—坳陷区(图 2、图 3)。

成都压陷区包含成都盆地。成都盆地位于青藏高原东缘龙门山造山带以东,扬子克拉通西缘,为挤压构造背景下发育的构造挠曲压陷盆地(Zhang *et al.*, 2010),该盆地古近纪地层由老到新依次为名山组和芦山组(图 3、图 4; 申琪, 2018)。

华南隆起—坳陷区包含南襄盆地、江汉—洞庭盆地、鄱阳湖盆地和衡阳盆地。以江汉—洞庭盆地为例细述之。江汉—洞庭盆地是在南(江南—雪峰造山带)、北(秦岭—大别造山带)两套体系夹持及西部黄陵隆起的制约下,在逆冲推覆基底上发育的中、新生代陆内断陷盆地(王德良等, 2018)。江汉—洞庭盆地古近纪地层由老到新依次为沙市组、新沟嘴组、荆沙组、潜江组和荆河镇组(图 3; 李云通, 1984; 徐论勋等, 1995)。

2.9 中国东部陆缘裂谷区

中国东部陆缘裂谷区根据盆地类型和盆地沉积充填序列特征(图 3),划分为东北陆缘裂谷区、胶

东—黄海断陷区、华东陆缘盆岭构造区和东海陆架伸展区(图2、图3)。

东北陆缘裂谷区包含依兰—伊通盆地、三江盆地和辽源东盆地。三江盆地是在佳木斯地块前侏罗纪褶皱基底上发育的中新生代断陷—坳陷盆地(刘志宏等,2011),盆地内古近纪沉积主要为宝泉岭组(图3;汪啸风和陈孝红,2005)。

胶东—黄海断陷区包含胶东盆地和宿迁盆地。

华东陆缘盆岭构造区包含苏北盆地、芜湖—九江盆地、绍兴—金华盆地、南雄盆地、三水盆地、茂名盆地、河源盆地、文昌盆地、北黄海盆地、南黄海盆地、东海陆架盆地、珠江口盆地、北部湾盆地、莺歌海盆地、琼东南盆地和台湾东盆地。以苏北盆地和南海北部次级盆地群为例细述之。苏北盆地位于郯庐断层带以东的扬子克拉通东北部,北邻苏鲁造山带,是在晚白垩世—古近纪西太平洋板块NW向的斜向俯冲使中国东部大陆构造体制由挤压转为拉张的背景之下发育的断陷—坳陷盆地(杨力,2014)。苏北盆地古近纪地层序列由老到新依次为阜宁组、戴南组和三垛组(图3;李云通,1984;江苏省地质矿产局,1997)。南海北部次级盆地群分别为珠江口盆地、琼东南盆地、北部湾盆地和莺歌海盆地,4个盆地均充填古近纪地层序列(解习农等,2015)。南海北部次级盆地群以珠江口盆地为代表,该盆地是在古生代及中生代复杂褶皱基底之上发育的呈NE走向展布的新生代断陷—坳陷盆地(林畅松等,2018),其古近纪地层序列由老到新依次为神狐组、文昌组、恩平组和珠海组(图3、图4;解习农等,2015)。

东海陆架伸展区包含东海陆架盆地。东海陆架盆地位于欧亚板块、西太平洋板块及澳大利亚—印度板块交接地带,是一个发育在克拉通基底之上的中、新生代断陷—坳陷陆缘裂谷盆地(张建培等,2014),其古近纪地层序列由老到新依次为月桂峰组、灵峰组、明月峰组、毆江组、温州组、平湖组和花港组(图3;Su *et al.*,2014)。

2.10 东南亚多岛弧盆系

东南亚多岛弧盆系根据盆地类型和盆地沉积充填序列特征(图3),划分为台东弧盆系和南海弧后洋盆(图2、图3)。

台东弧盆系古近纪时期主要发育奇美火成杂岩(郑家坚等,1999);南海弧后洋盆古近纪地层目前研究程度甚低,相关地层资料主要通过国际大洋

钻探计划获取(Wang *et al.*,2003)。

3 中国古近纪沉积—构造耦合演化历史

通过对中国古近纪地层区开展详细区划(图2),以及对各个构造—地层区内盆地类型及展布、岩石地层序列和地层接触关系(图3、图4)的讨论,本文在综合我们前期研究(宋博文等,2014;徐亚东等,2014)的基础之上,对中国古近纪沉积—构造耦合演化历史进行系统总结如下。

古近纪印度与欧亚板块的碰撞及之后青藏高原的形成是中国新生代最重大的构造事件之一(图1;Yin,2010;潘桂棠等,2016),深刻影响了中国构造地貌格局乃至全球新生代气候环境的演变。而目前有关印度与欧亚板块碰撞发生的具体时间依然存在较大争议,从70~34 Ma(Zhang *et al.*,2010;Wang *et al.*,2014;Ding *et al.*,2017;Hu *et al.*,2017;王二七,2017)。古近纪印度与欧亚板块碰撞的具体表现为:藏南冈底斯带林子宗群底部不整合面(约65 Ma)广布标志着印度与欧亚板块陆—陆碰撞造山作用开始(潘桂棠等,2016);雅鲁藏布江缝合带两侧沉积环境与物源在65 Ma发生显著变化(Ding *et al.*,2017)。初始碰撞之后的陆—陆碰撞造成了地壳的抬升和新特提斯海沿雅鲁藏布江缝合带退出闭合。而最高海相层位在亚洲南缘及藏南的分布显示新特提斯海的退出表现为明显的穿时性(Ding *et al.*,2017;Hu *et al.*,2017)。晚始新世34 Ma左右之后,新特提斯海在藏南全面退出,整个雅江缝合带和喜马拉雅地区缺乏海相沉积地层记录(Wang *et al.*,2002),该带连同冈底斯陆缘弧的大部分地区被构造抬升为剥蚀区。而副特提斯海从塔里木盆地最终退出的时限目前仍不确定,Bosboom *et al.*(2014)基于生物地层学与磁性地层学综合研究将副特提斯海在塔西南退出的时间限定在晚始新世(~41 Ma),而国内学者郝诒纯等(1982)和郭宪璞等(2002)则认为中新世以后塔里木盆地才彻底变为陆相盆地。随着陆—陆碰撞和俯冲作用的持续,在印度与欧亚板块碰撞及其远程效应影响下,高原北部古天山、昆仑山、祁连山、阿尔金等造山带在古近纪时期依次复活并构造抬升,发生强烈陆内造山作用(Wang *et al.*,2011;曹凯等,2018)。同时可可西里盆地、柴达木盆地、西宁—兰州盆地等大型盆地均在早始新世普遍发生压陷并开始接受沉积,从而

在中国西部形成近 EW 或 NWW 的造山带和内陆盆地相间分布的构造—地貌景观(宋博文等, 2014)。青藏高原东部中始新世—渐新世初则表现为构造挤压隆升, 导致地壳强烈缩短和增厚, 发育一系列压陷—走滑拉分盆地, 并伴有大规模火山喷发和岩浆侵入(Spurlin *et al.*, 2005; 周江羽和王江海, 2019)。

古近纪时期中国大陆构造变形东西差异明显, 古新世以来随着西太平洋板块的俯冲和后撤引起地幔上涌、岩石圈减薄伸展(Northrup *et al.*, 1995; Li *et al.*, 2012), 导致中国东部发生了与俯冲—碰撞相关的旋转和弧后扩展(Zhang *et al.*, 2019)。在此背景下, 华北克拉通陆内裂陷盆地主要沿郯庐断裂带分布, 如渤海湾盆地、胶东盆地、合肥盆地等(图 2), 其中渤海湾盆地在古近纪伸展作用强烈, 在大型低角度正断层的控制下快速拉张裂陷, 盆地内发育一系列半地垒—半地堑构造, 沉积了巨厚的沉积地层(Allen *et al.*, 1997); 在华北克拉通以西的环鄂尔多斯地块形成了银川—河套裂谷系、东缘的汾渭裂谷系, 发育银川盆地、河套盆地、渭河盆地等一系列断陷盆地(Zhang *et al.*, 2003; Shi *et al.*, 2020); 东北地区古近纪处于地壳伸展、岩石圈减薄的构造环境, 主要表现为 NW—SE 向的伸展构造应力体制, 沿佳木斯—伊通和敦化—密山两条巨型断裂带发育依兰—伊通具右行走滑特征的断陷盆地(刘志宏等, 2011), 远离走滑断裂带的裂陷盆地体系呈 NNE 向分布, 如三江盆地、虎林盆地等(图 2); 扬子大陆裂谷系包括北黄海盆地、南黄海盆地、东海陆架盆地、苏北盆地、河源盆地、三水盆地和茂名盆地等(图 2), 以上诸裂陷—断陷盆地均在古近纪时期发育形成, 在区域拉张伸展构造应力场下, 盆地成因机制和盆内构造发育均具有相似性(索艳慧等, 2012)。南海北部陆缘盆地群古近纪发生显著断陷, 古新世—始新世中期为陆相沉积序列(解习农等, 2015); 中晚始新世(~43 Ma), 随着西太平洋板块的运动方向由 NNW 向转为 NWW 向, 板块俯冲速度增大, 华南内陆沿海隆起, 裂陷中心南移并发生顺时针旋转, 南海北部陆缘裂陷活动达到高峰(索艳慧等, 2012), 南海北部盆地群主要发育海陆交互相沉积序列(解习农等, 2015); 渐新世晚期开始南海北部陆缘盆地群持续沉降, 发育海相沉积序列建造(解习农等, 2015)。35 Ma 之后, 现今中国东部沟—弧—盆体系下的西太平洋型大陆边缘开始形成(Li *et al.*, 2018)。新南海在早渐新世(~33 Ma)发生初

始海底扩张(Li *et al.*, 2014), 岩石圈破裂、洋壳出露, 东部次级海盆开始发育, 新南海海盆由东向西开始逐步打开(Wang and Jian, 2019)。

4 结论

(1) 本文主要基于构造控盆、盆控岩相和岩石组合的基本原则, 通过对古近纪大地构造、沉积盆地(群)、火山岩、隆起带和大型变形构造等的综合分析研究, 共厘定出中国古近纪一级构造—地层单元(大区)10 个、二级构造—地层单元(区)27 个。

(2) 在中国古近纪地层区划的基础上, 通过对各个地层分区中盆地的类型、构造背景、沉积序列及地层接触关系的研究, 本文建立了中国古近纪地层对比格架。

(3) 通过对中国古近纪时期构造与盆地沉积耦合过程的总结, 揭示出印度与欧亚板块碰撞和太平洋板块向西往欧亚板块之下俯冲共同塑造了中国古近纪的构造格局, 并导致该阶段中国陆域的西压东张的构造环境。

需要强调的是, 目前中国古近纪构造—地层区划的划分和地层对比格架的建立是我们前期工作的阶段性成果, 依然存在很多问题尚需今后更深入的探讨。其中就包括中国西部众多新生代盆地磁性地层时代的限定问题, 目前依然争议较大, 今后还需做更深入细致的工作以取得共识。

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