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大别山北淮阳带东段石榴斜长 角闪岩石炭纪变质作用的测定

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摘 要:北淮阳带位于大别碰撞造山带北部,相对于南部三叠纪超高压变质带来说,通常被认为是一个相对低级变质的构造 岩石单位.以商(城)一麻(城)为界,分为西段和东段.其中,东段未发现与西段相对应的古生代大洋俯冲记录,也缺乏相关的 古生代岩浆作用和变质作用方面报道.因而严重地阻碍了对大别造山带形成和演化的深刻理解和准确认识.北淮阳带东段岩 石类型丰富,主要包括3类(套)岩石:(1)变质岩,如变质花岗岩(花岗片麻岩)、(石榴)斜长角闪岩、大理岩、云母石英片岩和变 质复理石等;(2)岩浆岩,主要为中生代花岗岩、正长岩和火山岩等;(3)盆地沉积岩,主要为中新生代沉积及少量石炭纪沉积. 然而,该区一直未发现古生代变质作用的岩石学记录.为此,对北淮阳带东段金寨县铁冲一带与大理岩相共生的石榴斜长角 闪岩开展了锆石 SHRIMP U-Pb 定年和初步的岩石学研究.结果表明,该区石榴斜长角闪岩经历了355±5 Ma变质作用,因而 首次发现研究区经历了石炭纪变质作用,为秦岭一桐柏造山带的东延以及大别山碰撞造山带的古生代构造演化过程提供了新 的制约.

关键词: 锆石; SHRIMP U-Pb定年; 石榴斜长角闪岩; 石炭纪变质作用; 北淮阳带东段; 大别造山带; 地球化学.
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Dating of Carboniferous Metamorphism for Garnet Amphibolite from the Eastern Beihuaiyang Zone in the Dabie Orogen

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Abstract: In comparison with the Triassic ultra-high-pressure metamorphic belt in the southern Dabie orogen, the Beihuaiyang zone (BZ) is generally regarded to be a relatively low-grade metamorphic unit located in the northern part of the orogen. The BZ can be subdivided into eastern and western segments separated by Shangcheng-Macheng fault zone, respectively. Among of them, the eastern segment of the BZ has no record related to Paleozoic oceanic subduction similar to those in the western segment. Also, there are rare reports on the Paleozoic magmatism and metamorphism in the region. As a result, these have greatly impeded the deep understanding and precise recognition on the formation and evolution of the Dabie orogen. The eastern segment of the BZ is mainly composed of three groups of rocks: (1) metamorphic rocks, e.g., meta-granitoid (granitic gneiss), (garnet) amphibolite,

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marble, mica-quartz schist and meta-flysch; (2) igneous rocks mainly including Mesozoic granite, syenite and volcanics; (3) basin sediments mainly composed of Mesozoic-Cenozoic and subordinate Carboniferous rocks. However, no petrological record of the Paleozoic metamorphism has been reported in the region. Therefore, SHRIMP zircon U-Pb dating and preliminary petrographic observations on the garnet amphibolite coexisting with marble from Tiechong, Jinzhai in the eastern segment of the BZ are carried out in this study. The results suggest that the studied rocks experienced a strong metamorphism with the peak at 355 ± 5 Ma. It is reported for the first time on the Carboniferous metamorphism in the region, and thus provides new constraints on the eastern extension of the Qinling-Tongbai orogen and the Paleozoic tectonic evolution of the Dabie collisional orogen.

Key words: zircon; SHRIMP U-Pb dating; garnet amphibolite; Carboniferous metamorphism; eastern segment of the Beihuaiyang zone; Dabie orogen; geochemistry.

经典板块构造理论认为,碰撞造山带的形成通 常涉及到威尔逊旋回(Wilson cycle)的洋盆的打开 和关闭(Dewey and Spall, 1975),并伴随着与俯 冲一碰撞相关的变形、变质作用和最终因陆一陆碰 撞作用而形成造山带(Wilson, 1966; Dewey, 1969; Brown, 2009). 通常,形成于汇聚板块边缘的 造山带可以划分为增生型和碰撞型两大类.其中, 增生造山带主要形成于大洋板块的俯冲时期;而碰 撞造山带则形成于大洋板块俯冲结束后的陆一陆 碰撞阶段,最典型的例子是阿尔卑斯和喜马拉雅造 山带.然而,陆一陆碰撞之前往往涉及到洋盆的关 闭,所以对于大陆碰撞造山带而言,碰撞之前的大 洋俯冲及伴生的岛弧岩浆作用可能是必然的板块 构造过程(O'Brien, 2001).因此,这种复合造山作 用常造成碰撞构造叠加在早期增生构造之上 (Brown, 2007, 2009),从而为识别和重建两期造山 作用的精细过程增加了巨大困难.

中国中部近东西向延伸、长约2000 km 的秦 岭一桐柏一大别一苏鲁造山带,主要是由华北和华 南(扬子)两大陆块碰撞形成,并在陆一陆碰撞之前 曾经历了长期的大洋俯冲、岛弧增生和弧一陆碰撞 过程(张国伟等,1988,2001;Xu et al., 1992a;徐树 桐等, 1994, 2002; 杨经绥等, 2002; Dong et al., 2011; Wu and Zheng, 2013; Dong and Santosh, 2016),形成了南、北分带的增生造山体系和碰撞造 山体系(刘晓春等,2015). 然而,由于沿造山带横向 上构造过程的复杂性、多期性、复合性、叠置性和穿 时性(许志琴等,2015),不同地段出露的构造岩石 单元差别较大.其中,秦岭一桐柏一红安造山带均 保留了明显的古生代洋壳俯冲的证据,如北秦岭商 丹蛇绿混杂岩和古生代岛弧成因的岩石(如张国伟 等,1988,2001;孙卫东等,1995;裴先治等,2009; Dong et al., 2011, 2016; Liu et al., 2016)、桐柏一 红安北缘古生代复理石(Liu et al., 2004, 2011b)以 及熊店、胡家湾和苏家河古生代洋壳成因榴辉岩 (Sun et al., 2002; Cheng et al., 2009; Wu et al., 2009). 然而,东部大别-苏鲁造山带中却鲜见古生 代大洋俯冲的记录和证据:一方面可能与三叠纪 陆一陆强烈碰撞改造、燕山期热事件叠加以及多期 构造作用和破坏等有关,进而影响了人们认识华北 与华南板块之间的古生代一中生代演化的横向分 布;另一方面也与研究程度有关.秦岭一大别一苏 鲁造山带,又称中央造山带(杨经绥等,2002),晚元 古代以来,从冈瓦纳大陆分离的南、北中国板块,经 过原特提斯洋和古特提斯洋的演化以及板块多次 离散、汇聚和碰撞,形成显生宙以来以原特提斯和 古特提斯为主体的复合构造格架,以及以古生代和 印支期为主体的秦岭一大别一苏鲁复合造山系 (如, Mattauer et al., 1985; Hsü et al., 1987; 张国伟 等, 1988, 2001; 杨 经 绥 等, 2002; Ratschbacher et al., 2003, 2006; 刘良等, 2013; 许志琴等, 2015; Liu et al., 2016). 实际上,北秦岭古生代造山带及 商丹洋在大别山及相邻地区的东延问题至今仍没 有解决. Dong et al.(2011)根据区域地质背景分析, 认为商丹洋/商丹缝合带可能通过北秦岭向东延 伸,但至大别山之后,如何衔接,仍是值得研究的问 题;刘晓春等(2015)根据桐柏造山带的研究,推测 古生代商丹洋可能向东延伸到信阳、乃至商城地 区.然而,北淮阳带(尤其商一麻断裂以东;图1)尚 缺乏与之相对应的古生代大洋俯冲、岩浆作用和变 质作用等方面的岩石学记录及相关文献报道.为 此,本文开展了北淮阳带东段金寨一带与大理岩相 共生的石榴斜长角闪岩的锆石 SHRIMP U-Pb 定年 和初步的岩石学研究,证明研究区经历了石炭纪变 质作用,从而为秦岭--桐柏造山带的东延以及大别 山碰撞造山带的古生代构造演化过程提供了新的 制约.



BZ. 北淮阳带;NDZ. 北大别杂岩带;CDZ. 中大别超高压变质带;SDZ. 南大别低温榴辉岩带;SZ. 宿松变质带;HMZ. 浒湾混杂岩带;HZ. 红安 低温榴辉岩带;DC. 角闪岩相大别杂岩;XMF. 晓天一磨子潭断裂;WSF. 五河一水吼断裂;HMF. 花凉亭一弥陀断裂;TSF. 太湖一山龙断裂; TLF. 郯庐断裂;SMF. 商城一麻城断裂;DWF. 大悟断裂

1 地质背景

大别山是秦岭造山带的东延部分,东端被郑一 庐断裂带切割(图1). 郑一庐断裂带以东的苏鲁造 山带是大别山东延并位移了的部分.在地质位置 上,它位于华北和华南两个大陆板块之间,是华南 板块向华北板块之下俯冲形成的三叠纪大陆碰撞 造山带(Xu et al., 1992a, 1992b; Li et al., 1993). 而且,发育了与大陆俯冲和碰撞过程相关的、不同 变质等级的构造岩石单位(徐树桐等, 2002; Liu et al., 2007). 从南到北,大别山可分为宿松变质带、南 大别低温榴辉岩带、中大别超高压变质带、北大别 杂岩带及北淮阳带等构造岩石单位(图1). 但根据 区域地质背景分析和沉积地层的物源区研究等,大 别山印支期陆一陆碰撞之前应该存在已经消失的 古大洋(Xu et al., 1992a;徐树桐等,1994,2002; Li et al., 2004;李任伟等,2005;李双应等,2011).

北淮阳带西段(商一麻断裂以西),主要由二郎 坪群、原"信阳群"南湾组(变质复理石)和龟山组 (又称龟山杂岩)、原"苏家河群"中定远组、原商城 群"歪庙组"(可能对应于"二郎坪群";徐树桐等, 1994,2002)和原石炭纪梅山群等岩石单位以及变 质火成岩、古生代闪长岩等花岗岩类岩石和中生代 岩石等组成(图1).其中,原岩时代为新元古代晚期 (~630 Ma)(刘贻灿等,2006;Liu et al., 2017)的变 质(橄榄)辉长岩沿苏家河一八里畈断裂的北侧,从 千斤河棚乡王母观向西经苏家河至信阳南部西双 河和桐柏一带呈大小不等的岩块或岩片出露,千斤



Fig.2 Schematic geological map of the Tiechong area, Jinzhai in the eastern segment of the Beihuaiyang zone

河棚乡向东经吴陈河乡至八里畈乡一带也有类似 岩石断续分布.其围岩为原"定远组"变质火山岩, 目前表现为含石榴子石绿帘云母石英片岩.二者之 间为断层接触,统称为定远变质火山岩带(刘贻灿 等,2006)或肖家庙一八里畈构造混杂岩带(刘晓春 等,2015).其南、北分别与浒湾混杂岩带和"南湾 组"变质复理石等构造岩石单位相邻,再向南为新 县超高压变质带.原苏家河群"浒湾组"中既有石炭 纪洋壳俯冲成因榴辉岩(如熊店),又有三叠纪陆壳 俯冲成因榴辉岩,它们的原岩时代分别为晚古生代 和新元古代(Sun et al., 2002; Cheng et al., 2009; Wu et al., 2009),因而称之为浒湾混杂岩带 (Ratschbacher et al., 2006;刘贻灿等,2006)或浒湾 高压榴辉岩带(刘晓春等,2015),它的变质相及形成时代完全不同于岛弧成因的定远变质火山岩即 峰期表现为绿帘角闪岩相变质作用和原岩时代为 早古生代奥陶纪(Li et al., 2001;刘贻灿等,2006).

北淮阳带东段(商一麻断裂以东),主要由"佛 子岭群"变质复理石(对应于秦岭的"刘岭群"和桐 柏的"南湾组")和庐镇关杂岩(原"庐镇关群")和原 石炭纪梅山群(梅山煤系)等构造岩石单位及中新 生代岩浆岩和盆地沉积组成(徐树桐等,1994, 2002;Chen et al., 2003;吴元保等,2004;江来利等, 2005;刘贻灿等,2006,2010;刘景波等,2013)(图 1).其中,庐镇关杂岩主要包括原"小溪河组"新元 古代变质花岗岩和变基性岩等和"仙人冲组"大理 岩及相伴生的(石榴)斜长角闪岩等.然而,由于研究区可能因印支期大陆的强烈碰撞改造和构造叠加,早期一些岩石单位被破坏或者被盆地沉积所掩盖(Xu et al., 1992a;徐树桐等,1994,2002),造成至今未发现确切的与古生代大洋俯冲相关的岩石单位或记录.然而,直到最近,刘贻灿等在金寨县西部查明存在古生代花岗岩(其时代为457±2 Ma,未发表资料;图 2),而且证明研究区经历了石炭纪变质作用(见后文).

本文研究的样品为采集于李桥南(图2)原"仙 人冲组"与大理岩相共生的石榴斜长角闪岩.其中, 石榴斜长角闪岩以构造透镜体的形式产出于大理 岩中,二者共同发生了复杂的变质、变形作用(图3).

2 样品描述和分析方法

岩相学研究表明,石榴斜长角闪岩的主要矿物 有石榴子石、斜长石、单斜辉石、角闪石和少量的石 英、方解石和榍石等(图4).其峰期变质矿物主要为 石榴子石+斜长石+单斜辉石+石英等,至少达到 麻粒岩相、甚至可能达到石英榴辉岩相条件(见后 文分析),至于具体变质过程和峰期温压条件,还有 待于进一步研究(将另文讨论).

岩石样品经破碎、筛选、电磁及重液分选分离

出锆石,而后在体视显微镜下对未蚀变、透明等的 颗粒进行挑选,再将其和标准锆石 TEMORA(年龄 为417 Ma)一起制成样品靶,磨至一半后抛光并在 北京离子探针中心进行透射光、反射光和阴极发光 (CL)显微照相.其中,锆石的分选工作由河北省区 域地质矿产调查研究所实验室完成;锆石微区 U-Pb 定年测定在北京离子探针中心 SHRIMP II 离子探 针 仪器上进行,详细测定程序见有关文献



- 图 3 北淮阳带东段金寨县李桥与大理岩相伴生的石榴斜 长角闪岩野外照片
- Fig.3 A field photograph showing garnet amphibolite coexisting with marble in the eastern segment of the Beihuaiyang zone at Liqiao, Jinzhai



图 4 北淮阳带东段石榴斜长角闪岩的显微照片 Fig.4 Photomicrographs of garnet amphibolite from the eastern Beihuaiyang zone Grt. 石榴子石; Pl. 斜长石; Cpx. 单斜辉石; Hbl. 角闪石; Ttn. 榍石; Cc. 方解石



图 5 代表性锆石的阴极发光图像 Fig.5 Cathodoluminescene (CL) images of representative zircons



(Compston *et al.*, 1992; Williams, 1998; 宋彪等, 2002). 测试时所用的标准锆石为 M257 和 TEMO-RA,前者用于标定 U含量,后者用于校正²⁰⁶Pb/²³⁸U 比值(亦即²⁰⁶Pb/²³⁸U年龄). 束斑直径约为 20 μm. 测定结果用²⁰⁸Pb 进行普通 Pb 校正,详细数据处理 过程见 Liu *et al.*(2017). 测试数据列于表 1 中.

3 分析结果

阴极发光(CL)图像(图5)显示,样品中的锆石 都表现为无分带或区域分带及浑圆状外形等、无明 显岩浆结晶环带,结合它们的低 Th/U 值(0.02~0.31),应为变质生长锆石或变质重结晶锆石 (Vavra *et al.*, 1999; Pidgeon *et al.*, 2000; Rubatto *et al.*, 2001; Schmitz and Bowring, 2003).

锆石 U-Pb 同位素数据和年龄结果如表1和图6 所示.本文分析了16颗锆石,获得的年龄都为谐和 年龄,而且²⁰⁶Pb/²³⁸U年龄为346±15~365±12 Ma, 加权平均值为355±5 Ma(MSWD = 0.22).

因此,根据锆石的阴极发光图像特点和低的低 Th/U值,355±5 Ma应为石榴斜长角闪岩的变质 时代.

4 讨论

Mattauer et al.(1985)在商丹断裂南侧刘岭群 (武关杂岩)的变泥质岩中获得314±6 Ma的黑云 母⁴⁰Ar/³⁹Ar年龄, Zhai et al.(1998)在桐柏龟山杂岩 斜长角闪岩中获得两个角闪石⁴⁰Ar/³⁹Ar年龄,分别 为316±1 Ma和304±14 Ma,从而证实秦岭一桐柏 造山带中确实存在晚古生代构造事件.近年来,桐 柏一红安造山带北部的龟山杂岩被确定是秦岭大 陆弧与弧前复理石的混杂体,其在340~310 Ma经 历了角闪岩相变质过程(Liu et al., 2011a, 2013a, 2013b;曲玮等,2018).同时,对秦岭造山带商丹断 裂南侧从刘岭群中解体出来的武关杂岩的研究表 明,该套以沉积岩为主体的中级变质杂岩中含有大 陆弧成因变质火成岩,包括原岩时代分别为446±2 Ma的斜长角闪岩、368±3 Ma的安山质片麻岩和 351±2 Ma的糜棱岩化花岗岩脉,其变质时代为~ 320 Ma(Chen et al., 2014). 而且,已有研究表明,北 秦岭(包括秦岭、二郎坪和宽坪岩群等)的变质时代 至少包括两期即 450~420 Ma 和 310~350 Ma (Ratschbacher et al., 2003, 2006; Dong et al., 2011, 2016; Liu et al., 2011a; Chen et al., 2014;刘 晓春等,2015,及所引文献; Yan et al., 2016),经历 了早古生代(500~400 Ma)大洋俯冲、岛弧增生与 弧一陆碰撞及晚古生代(340~310 Ma)大洋俯冲与 增生等复杂的构造过程(王涛等,2009;裴先治等, 2007,2009;第五春荣等,2010;刘良等,2013;Wang et al., 2013a, 2013b;张成立等, 2013;王晓霞等, 2015; Liu et al., 2016); 北淮阳带西段熊店榴辉岩 带的原岩应主要是洋壳成因,其形成时代主要为 425~410 Ma,并经历了石炭纪(~310 Ma)洋壳俯 冲及榴辉岩相高压变质作用(Sun et al., 2002;

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²⁰⁶Pb_c U Th ²³²Th ²⁰⁶Pb* ²⁰⁷Pb* ²⁰⁷Pb* ²⁰⁶Pb* ²⁰⁶Pb/²³⁸U 分析点号 10^{-6} 10^{-6} $/^{238}$ U 10^{-6} $(\frac{0}{0})$ $/^{206} Pb^* (\pm \%)$ $/^{235}U(\pm \%)$ $/^{238}U(\pm \%)$ 年龄(Ma) 1705TC9_1.1 4.17 6.3 1.30 0.21 0.31 0.0516 0.055.9 350 ± 13 16.00.39717.03.3 1705TC9_2.1 7.57 0.02 0.0620 34.0 0.470 34.0 347 ± 10 5.30.12 0.27 0.05522.51705TC9 3.1 8.93 36 0.11 0.03 0.19 0.041041.0 0.320 41.0 0.057 2 3.0 $359\!\pm\!19$ 1705TC9 4.1 4.19 6.3 1.40 0.24 0.32 0.073 5 12.0 0.581 12.0 0.057 3 2.2 $359\!\pm\!10$ 1705TC9_5.1 8.93 4.7 0.770.17 0.250.084.0 23.00.660 23.00.056 6 2.6 355 ± 12 1705TC9_6.1 2.508.9 2.00 0.23 0.45 0.0596 10.0 0.467 10.0 0.056.9 1.9 356 ± 7 1705TC9 7.1 3.83 15.0 2.30 0.16 0.77 0.0623 11.0 0.485 12.0 0.056 4 1.5 354 ± 6 1705TC9_8.1 15.0 4.60 0.31 0.044 9 7.6 0.347 7.8 $0.056\ 1$ 352 ± 6 2.06 0.76 1.5 1705TC9_9.1 9.13 4.2 0.61 0.15 0.22 0.067 0 21.0 0.520 21.0 0.056 3 2.9 $353\!\pm\!13$ 1705TC9_10.1 17.0 0.321 4.295.8 0.74 0.13 0.30 0.0403 17.00.0577 2.4 362 ± 10 1705TC9 11.1 0.0690 19.0 0.540 5.34 4 2 0.72 0.18 0.22 19.0 0.057.0 29 358 ± 12 1705TC9 12.1 3 20 6.0 1.00 0.17 0.30 0.058.0 19.00 457 19.0 0.057.0 2.3 357 ± 9 1705TC9 13.1 2.24 11.0 2.40 0.22 0.55 0.068 5 15.0 0.541 15.0 0.0573 1.9 359 ± 8 1705TC9_14.1 10.0 2.00 0.21 7.63 0.53 0.0600 28.00.480 28.0 0.058 2 2.2 365 ± 12 1705TC9_15.1 10.08 0.17 0.32 26.0 0.520 6.0 1.00 0.068.0 26.0 0.05522.5 346 ± 15 1705TC9 16.1 24.60 4.5 0.82 0.19 0.29 0.080.0 51.0 0.640 51.0 0.0577 3.9 $361\!\pm\!27$

表1 北淮阳带东段石榴斜长角闪岩的锆石 SHRIMP U-Pb 数据^(a)

Table 1 SHRIMP zircon U-Pb data for garnet amphibolite from the eastern Beihuaiyang zone

注:(a)误差为1σ, Pb_e和 Pb*分别表示普通铅和放射性成因铅;用²⁰⁸Pb进行普通铅扣除.

Cheng et al., 2009, 2016, 2018; Wu et al., 2009).

此外,刘贻灿等(未发表资料)最新野外地质调 查和锆石U-Pb定年结果表明,北淮阳带东段金寨 一带发育具有岛弧成因的早古生代(457±2 Ma)花 岗岩.该花岗岩的形成时代类似于北秦岭二郎坪地 区的满子营岛弧成因花岗岩(459.5±0.9 Ma;郭彩 莲等,2010)以及桐柏地区的马畈闪长岩(463.5± 3.4 Ma;马昌前等,2004)和"定远组"岛弧成因的变 质火山岩(464±7 Ma; 刘贻灿等, 2006)等. 由此证 明,北淮阳带东段存在早古生代(奥陶纪-志留纪) 与大洋俯冲相关的岩浆作用.结合本文石炭纪变质 时代的确定,证明北淮阳带东段的大洋俯冲可能从 早古生代奥陶纪(~450 Ma)即已开始,并在石炭纪 发生变质作用.

正如前文所述,本文研究样品的锆石具有一致 的阴极发光图像(图5)和低的Th/U值(0.02~0.31) (表1),为变质生长或变质重结晶锆石,偶尔见残留 的岩浆核(图 5a, 5c),证明该样品经历了较高温的 变质作用(Liu et al., 2015; 王程程等, 2018). 结合 石榴子石+斜长石+单斜辉石+石英等变质矿物 组合,推断该石榴斜长角闪岩的峰期变质作用至少 达到麻粒岩相、甚至榴辉岩相变质 P-T 条件, 可能 与泥盆纪晚期弧一陆碰撞有关.这也与北淮阳带石 炭纪沉积地层学记录相吻合:

(1) 金福全等(1987) 的古生物地层学研究表 明:①早石炭世杨山组砾石中含有形成时代主要为 晚奧陶世-早志留世的珊瑚类、介形类和牙形石类 等化石组合.特别是砾石中所含Heliolites cf. Anhuiensis(安徽日射珊瑚比较种),则见于扬子陆块的 安徽含山早志留世高家边组和三峡早志留世罗惹 坪组,证明灰岩砾石的源区为扬子陆块,也就是说, 北淮阳地区存在类似于扬子的早古生代地层;②晚 石炭世胡油坊组发现了丰富的小型 Protomonocarina(原单脊叶肢介)化石(仅见于华北陆块),指示晚 石炭世北淮阳带和华北陆块处于同一个古生物区 系,两者是联为一体的.因此,这不仅反映了北淮阳 带与扬子陆块之间在石炭纪没有分隔性大洋,而且 也证明扬子陆块和华北陆块在石炭纪以前已经 拼接.

(2)北淮阳带石炭系未变质或轻微变质的沉积 地层(如胡油坊组、杨山组等),记录了其物源信息 和古生代的构造演化(李双应等,2011).其中:①碎 屑岩的微量元素地球化学特征揭示其物源区的大 地构造属性为岛弧;②碎屑锆石U-Pb年龄和碎屑 白云母的Rb-Sr年龄也进一步确定了岛弧的时代为 古生代(400~500 Ma)(Li et al., 2004; 李任伟等, 2005; Chen et al., 2009; 杨栋栋等, 2012), 这与北秦 岭、桐柏及北淮阳带西段(定远变质火山岩)的岛弧

时代相一致.因此,石炭纪沉积地层的部分物源区 类似于北秦岭等地的古生代岛弧岩石,也就是说, 北淮阳带在早古生代也应发育岛弧体系及相关岩 石并为石炭纪盆地沉积提供了重要的物源.

综上所述,北淮阳带东段存在类似于北秦岭和 桐柏造山带的早古生代岩浆热事件和晚古生代变 质作用,而且,研究区向北的(大洋)板块俯冲作用 可能从早古生代开始,一直延续到石炭纪;同时,在 (晚泥盆纪一)早石炭纪时,发生了弧一陆碰撞作 用.由此证明北秦岭古生代造山带及商丹洋向东延 伸到大别山北淮阳带及相邻地区.然而,北淮阳带 东段石炭纪变质作用的大地构造背景、古生代的构 造演化过程和汇集过程与拼贴机制以及与北秦岭 和桐柏造山带的不同构造岩石单位之间相互关系 尚需进一步查明,尤其需要详细的野外地质调查及 系统的岩石学、年代学和地球化学等方面制约.

5 结论

(1)北淮阳带东段与大理岩相伴生的石榴斜长 角闪岩经历了 355±5 Ma变质作用.

(2)北淮阳带东段石炭纪变质作用的厘定,为 秦岭一桐柏造山带的东延以及大别山碰撞造山带 的古生代构造演化过程提供了新的制约.

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