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海南岛西北部早古生代安山岩的识别及其大地构造意义

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摘 要: 为阐明海南岛西北部早古生代构造演化特征, 对海南岛西北部邦溪-晨星地区出露的安山岩展开了详细的岩相学和锆石 U-Pb 年代学分析. 安山岩的主要矿物组成为斜长石、角闪石、黑云母和玻璃质. LA-ICP-MS 锆石 U-Pb 定年结果表明, 两个代表性的安山岩样品给出了 452 ± 8 Ma 和 453 ± 8 Ma 的形成年龄, 暗示海南岛西北部存在早古生代火山作用. 这些安山岩的形成时代与冈瓦纳大陆北缘各微陆块记录的早古生代岩浆活动信息相似, 在时空上与金沙江-点苍山-哀牢山构造带和印支地块分布的早古生代岩浆岩具有耦合性. 综合研究认为, 点苍山-哀牢山、Truong Son、Tam Ky-Phuoc Son、Kontum 和海南岛等地区可能发育原特提斯洋的支洋盆.

关键词: 安山岩; 早古生代; 原特提斯洋; 锆石 U-Pb 定年; 海南岛; 岩石学.

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Identification of Early Paleozoic Andesite in Northwestern Hainan Island and Its Geotectonic Significances

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Abstract: In order to decipher the Early Paleozoic tectonic evolution of the northwestern Hainan Island, detailed petrography and zircon U-Pb geochronological analyses of andesites in the Bangxi-Chenxing area were carried out in this study. The andesites are mainly composed of plagioclase, hornblende, biotite and glasses. LA-ICP-MS zircon U-Pb dating results show that two representative andesite samples yield ages of 452 ± 8 Ma and 453 ± 8 Ma, indicating that there is andesitic magmatism in the Bangxi-Chenxing area during the Early Paleozoic. The formation age of the andesites is similar to the time of Early Paleozoic magmatism recorded in the micro-continents along the northern Gondwana. They also show a good coupling with the Early Paleozoic igneous rocks in the Jinshajiang-Diancangshan-Ailaoshan tectonic belt and in the Indochina block. Integrated with regional data, the Diancangshan-Ailaoshan, Truong Son, Tam Ky-Phuoc Son, Kontum and Hainan Island were likely to be branch of Proto-Tethys Oceans during the Early Paleozoic.

Key words: andesite; Early Paleozoic; Proto-Tethys Ocean; zircon U-Pb dating; Hainan Island; petrology.

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0 引言

新元古代 Rodinia 超大陆裂解后逐渐演化产生了原特提斯洋(von Raumer and Stampfli, 2008),其形成时代多被认为是震旦纪至早古生代,并于早古生代晚期由于 Gondwana 大陆的聚合而关闭(钟大赉, 1998; Li *et al.*, 2018a, 2018b).原特提斯洋的遗迹主要呈近东西向(现今方位)分布在东亚各微陆块,如在我国昆仑、塔里木、羌塘、柴达木、祁连、秦岭等地区均保存有大量 510~400 Ma 的岩浆活动、变质作用和沉积作用等地质记录(李三忠等, 2016; 杨鑫等, 2020).前人在华南扬子地块西南缘的昌宁-孟连缝合带、澜沧江缝合带和腾冲-保山地块等地区相继报道了早古生代蛇绿岩、变火山岩和过铝质花岗岩(Wang *et al.*, 2013a, 2013b; Nie *et al.*, 2015; 吴喆等, 2020)等岩浆作用信息,暗示原特提斯洋向东南延伸进入了我国的滇西一带,昌宁-孟连缝合带、澜沧江缝合带、腾冲-保山一带发育原特提斯洋主洋盆或支洋盆.近年来, Liu *et al.* (2018)在点苍山-哀牢山构造带识别出了早古生代富铌玄武岩、高镁安山岩和花岗闪长岩,表明点苍山-哀牢山构造带亦发育有原特提斯洋支洋盆.然而,对于原特提斯洋向东南延伸至何处,目前仍不明了,仍缺少精细的年代学证据,不利于深入认识原特提斯洋的构造演化历史.本文在海南岛西北部的西昌-邦溪一带识别出了早古生代安山岩,锆石 U-Pb 年代学显示其形成于~452 Ma,反映海南岛存在早古生代岩浆活动,其形成时代与点苍山-哀牢山等地区分布的岩浆岩具有可对比性,暗示海南岛可能存在原特提斯洋的支洋盆,这一新认识为深入理解原特提斯洋的构造格局和 Gondwana 大陆的聚合过程提供了新的证据.

1 区域地质背景和样品描述

1.1 区域地质背景

海南岛位于华南板块的最南缘(图 1),一般被认为是华夏地块的组成部分.岛内构造线方向主要呈东西向、北东向和北西向,自南向北分布有九所-陵水、尖峰-吊罗、昌江-琼海和王五-文教 4 条隐伏断裂(图 1;广东省地质矿产局, 1988).

岛内分布的已知最老地层为元古界抱板群、石碌群和石灰顶组,主要由一套高绿片岩相-角闪岩

相变质岩系组成(Li *et al.*, 2008).寒武系主要发育在三亚东北部和万宁北部地区,主要由石英砂岩、泥质粉砂岩、粘土岩、千枚岩、硅质岩和碳酸盐岩等组成.奥陶系下部主要由石英砂岩、碳酸盐岩、页岩等组成,上部见粉砂岩、粘土岩、千枚岩、变质细砂岩、碳质板岩夹火山岩,主要分布于邦溪、儋州、屯昌、万宁、五指山、乐东、东方等地区(广东省地质矿产局, 1988).志留系、石炭系、二叠系、白垩系、第四系和少量三叠系也发育于海南岛(图 1).

海南岛岩浆作用异常强烈,形成了大规模不同期次的岩浆岩,其中,侵入岩分布面积约占全岛面积的 51%,火山岩出露面积约为 13%.侵入岩基本上遍布整个海南岛,并以二叠纪和中生代花岗岩类为主(图 1);火山岩以新生代玄武岩为主,主要见于海南岛北部的第四系(图 1).少量元古代(Li *et al.*, 2008)、志留纪-二叠纪(李献华等, 2000; 谢才富等, 2006; 苟琪钰等, 2019)和白垩纪(Zhou *et al.*, 2015)花岗质岩石和基性-变基性岩分布于海南岛抱板群、昌江、琼中、屯昌、三亚等地区.

1.2 样品描述

本文研究的安山岩样品(样品号 CX1621 和 CX1624)分别采自海南岛西北部的邦溪和西昌晨星地区(图 1),它们呈似层状或不规则状分别产于《1:50 万海南省地质图》划分的奥陶系和二叠系中,其中,邦溪地区安山岩呈东西向展布,出露宽大于 50 m,表面风化较严重,呈土黄-土灰色(图 2a),新鲜面呈灰黑色,斑状结构,块状构造(图 2b),斑晶含量约为 80%,主要为斜长石(~60%)、角闪石(~30%)和黑云母(~10%),基质含量约为 20%,主要为细颗粒斜长石(~60%)、玻璃质(~35%)和少量黑云母(~5%);斜长石和角闪石斑晶多呈自形-半自形,颗粒大小多在 200~300 μm (图 2c).西昌晨星地区的安山岩整体呈北东向展布,出露宽大于 20 m,岩石较新鲜,碎裂明显(图 2d),斑状结构,块状构造,斑晶含量约为 90%,主要为斜长石(~50%)、黑云母(~20%)、角闪石(~20%)和石英(~10%),基质含量约为 10%,主要为细粒斜长石(~60%)、黑云母(~30%)和玻璃质(~10%);斑晶斜长石和角闪石多呈残斑状(图 2e);本文的年代学研究结果显示本区安山岩的形成时代为晚奥陶世,暗示本区原划分的二叠系可能需要解体,部分地层的时代至少不晚于晚奥陶世.

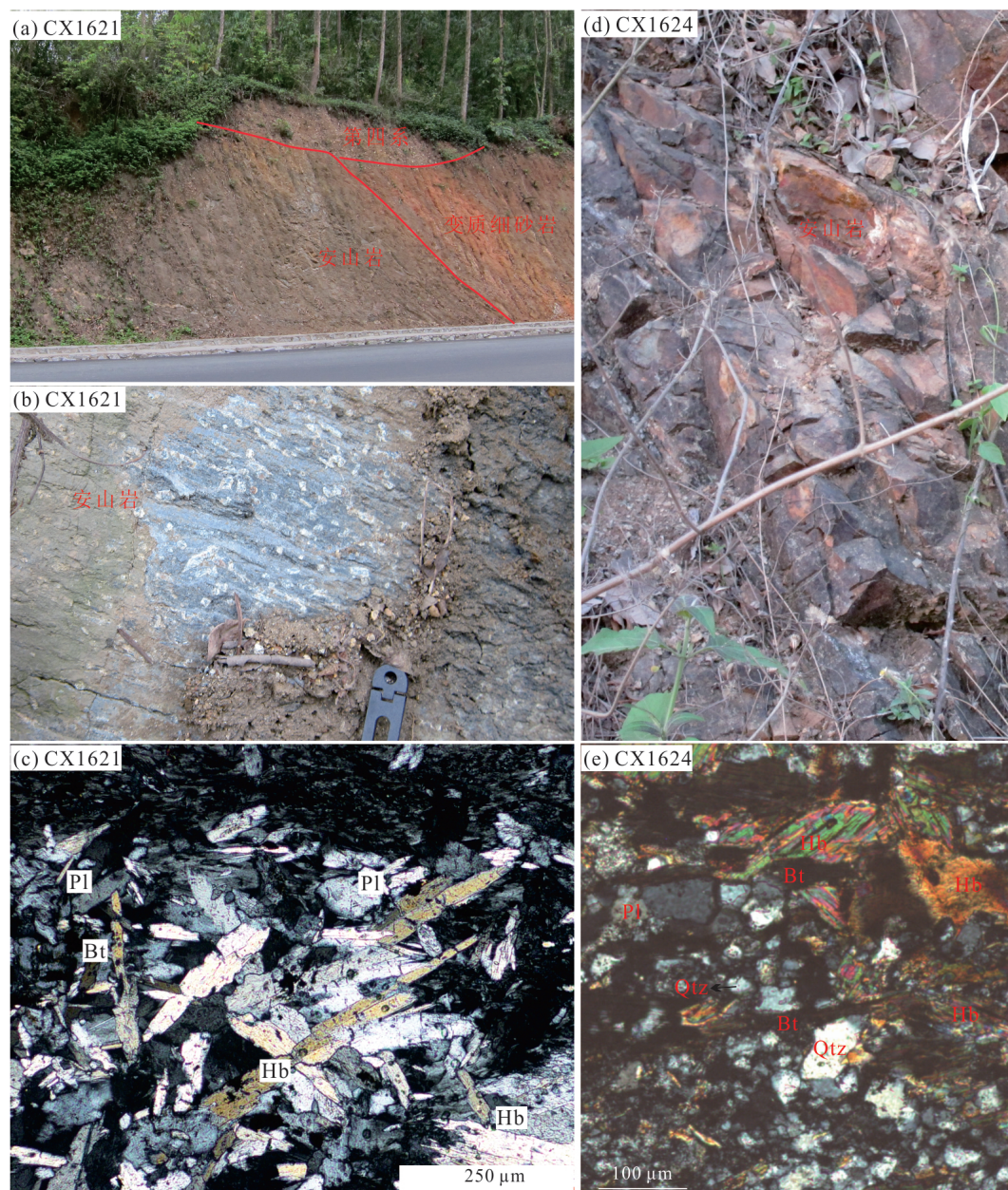


图2 海南岛西北部邦溪-西昌安山岩野外照片(a, b, d)和显微照片(c, e)

Fig.2 Field photos (a, b, d) and photomicrographs (c, e) of the representative andesites from Bangxi-Xichang, northwestern Hainan Island

a. 岩石表面风化较严重; b. 新鲜面为斑状结构; d. 块状构造; c, e. 斑晶主要为斜长石和角闪石. Pl. 斜长石; Hb. 角闪石; Bt. 黑云母; Qtz. 石英

部分锆石呈柱状, 长度一般为 $100\sim 150\ \mu\text{m}$, 内部发育明显的振荡环带(图3), 其 Th/U 比值较高, 均大于 0.2(表1), 为典型的岩浆成因锆石.

样品 CX1621 进行了 20 个锆石点分析测试, 其 Th/U 比值变化于 $0.23\sim 0.90$. 有 16 个测试点给出了 $452\pm 8\ \text{Ma}$ ($\text{MSWD} = 0.23$) 的 $^{206}\text{Pb}/^{238}\text{U}$ 加权平均年龄(图3a), 代表了其形成年龄; 另外有 2 个点给出了 $1\ 428\ \text{Ma}$ 和 $1\ 369\ \text{Ma}$ 的 $^{207}\text{Pb}/^{206}\text{Pb}$ 年龄, 可能代

表了捕获锆石或继承锆石的年龄.

样品 CX1624 进行了 20 个锆石点的分析, 其 Th/U 比值变化范围为 $0.24\sim 1.21$. 其中, 有 15 个测试点给出了 $453\pm 8\ \text{Ma}$ ($\text{MSWD} = 0.54$) 的 $^{206}\text{Pb}/^{238}\text{U}$ 加权平均年龄(图3b), 代表了其形成时代; 另外有 2 个点给出了 $1\ 533\ \text{Ma}$ 和 $1\ 528\ \text{Ma}$ 、3 个点给出了 $1\ 877\sim 1\ 814\ \text{Ma}$ 的 $^{207}\text{Pb}/^{206}\text{Pb}$ 年龄, 推测为捕获锆石或继承锆石的年龄.

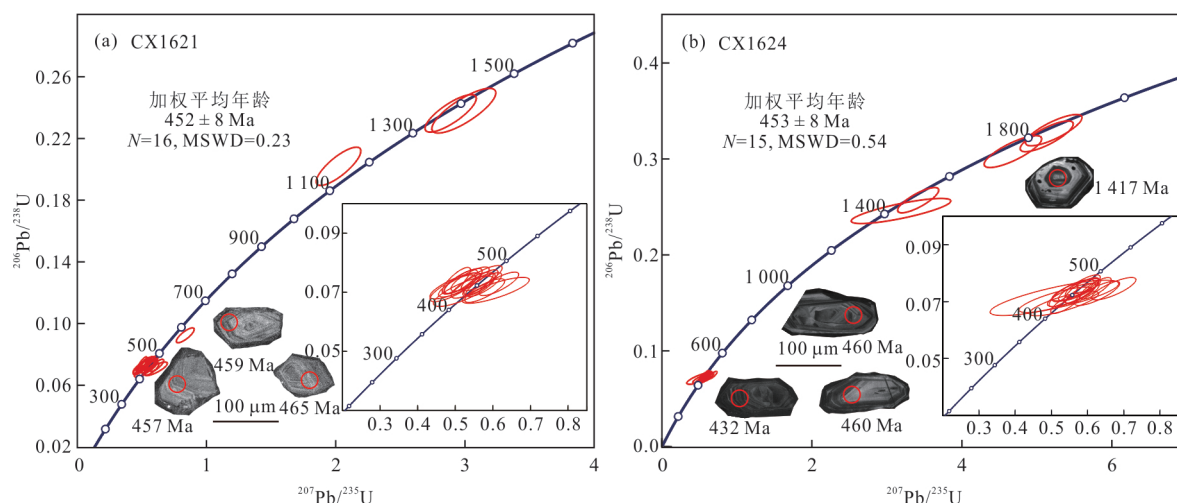


图3 海南岛西北部邦溪-西昌安山岩锆石 U-Pb 年龄谐和图和代表性锆石 CL 图像

Fig.3 Zircon U-Pb concordia diagrams and CL images of the andesites from Bangxi-Xichang, northwestern Hainan Island

4 讨论

4.1 海南岛西北部邦溪-晨星地区火山岩形成时代

对于海南岛西北部邦溪-晨星一带分布的火山岩的形成时代前人开展过一些研究. 张业明等 (1998) 提出海南岛邦溪-军营一带变基性火山岩的形成时代可能为 $\sim 1\,165$ Ma (Sm-Nd 模式年龄). 李献华等 (2000) 认为出露在昌江-琼海断裂带北侧的屯昌晨星和昌江邦溪地区的变基性岩的形成时代为 ~ 333 Ma (Sm-Nd 等时线年龄), 代表了东古特提斯洋的残片. 丁式江等 (2002) 的研究认为屯昌变火山岩形成于 ~ 527 Ma (锆石 U-Pb 交点年龄). He *et al.* (2018) 将邦溪-晨星地区变基性火山岩的形成时代限定为 $330\sim 328$ Ma (全岩 Ar-Ar 和 LA-ICP-MS 锆石 U-Pb 年龄). Li *et al.* (2018b) 获得本区安山质-英安质岩石的年龄为 ~ 350 Ma, 同时还获得了 ~ 433 Ma 的年龄峰值, 并认为后者代表捕获锆石年龄.

本文对采自邦溪和西昌晨星地区两个代表性安山岩样品的年代学研究分别获得了 452 ± 8 Ma 和 453 ± 8 Ma 的年龄, 同时存在少量 $\sim 1\,400$ Ma 和 $\sim 1\,800$ Ma 的老年信息 (图 3), 这些年龄信息与邦溪地区奥陶系和西昌晨星地区二叠系沉积岩 (本文研究样品的围岩) 中碎屑锆石的年龄分布不一致 (张立敏等, 2017), 说明本文研究样品中的 ~ 450 Ma 锆石并非捕获围岩的. 此外, 样品中的锆石未见含有石英、钾长石、磷灰石、白云母、黑云母等壳源矿物包裹体, 也表明这些锆石并非为捕虏锆石. CL 图像显示, 锆石多具有继承核和振荡

环带, 其中继承核给出了元古代年龄, 环带为 ~ 450 Ma (图 3b), 说明部分锆石是元古代锆石的继承, 而 ~ 450 Ma 应该代表了样品的形成时代. 上述研究成果暗示邦溪-晨星一带分布的火山岩可能为一套混杂岩, 包含早古生代 (~ 450 Ma)、早石炭世 ($350\sim 328$ Ma) 和元古代等不同时代岩石类型.

4.2 大地构造意义

Rodinia 超大陆在新元古代时期裂解后拉开了原特提斯洋的演化序幕 (von Raumer and Stampfli, 2008; Li *et al.*, 2018b), 原特提斯洋的俯冲作用促使游离在冈瓦纳大陆北缘的各微陆块相继拼贴到冈瓦纳大陆. 前人在冈瓦纳大陆北缘的土耳其、伊朗、巴基斯坦、尼泊尔、印支、滇缅泰马等微陆块开展了大量研究工作, 获得了众多新元古代晚期-早古生代岩浆活动信息 (图 4). 比如, Okay *et al.* (2008) 在土耳其西北部识别出了 570 ± 3 Ma 和 569 ± 2 Ma 的花岗质岩石和 569 ± 1 Ma 的花岗质脉体, 并认为这些花岗质侵入岩形成于活动大陆边缘环境, 是对原特提斯洋俯冲的响应. 伊朗中部 Kuh-e-Sarhangi 地区花岗质岩石的形成时代为 $576\sim 536$ Ma (Rossetti *et al.*, 2015)、变基性岩的锆石 U-Pb 年龄为 $554\sim 534$ Ma (Hosseini *et al.*, 2015), 它们的形成被认为与原特提斯洋的斜向俯冲汇聚有关 (Hosseini *et al.*, 2015; Rossetti *et al.*, 2015). Sajid *et al.* (2018) 在巴基斯坦西北部 Utla 和 Mansehra 地区报道了 $480\sim 466$ Ma 的花岗岩, 并认为其形成于安第斯型的活动大陆边缘, 是原特提斯洋俯冲至冈瓦纳大陆北缘的产物. 尼泊尔中西部发育大量早古生代花岗质岩石 ($512\sim 467$ Ma) 及捕虏

表 1 海南岛西北部邦溪-西昌安山岩锆石 U-Pb 同位素测试结果

Table 1 Zircon U-Pb dating results of the andesites from Bangxi-Xichang, northwestern Hainan Island

点号	含量 (10 ⁻⁶)		Th/ U	同位素比值						年龄 (Ma)						
	²³² Th	²³⁸ U		²⁰⁷ Pb/ ²⁰⁶ Pb	²⁰⁷ Pb/	1σ	²⁰⁷ Pb/	1σ	²⁰⁶ Pb/	1σ	²⁰⁷ Pb/	1σ	²⁰⁷ Pb/	1σ	²⁰⁶ Pb/	1σ
					²³⁵ U		²³⁸ U		²³⁵ U		²³⁸ U					
样品 CX1621																
CX1621-01	96	424	0.23	0.090 0	0.004 5	2.991 7	0.163 4	0.238 0	0.009 6	1 428	96	1 406	42	1 376	50	
CX1621-02	201	482	0.42	0.050 8	0.003 3	0.521 3	0.033 2	0.073 9	0.002 8	232	152	426	22	459	17	
CX1621-03	117	322	0.36	0.052 3	0.003 2	0.539 5	0.033 5	0.074 5	0.002 7	298	141	438	22	463	16	
CX1621-04	206	246	0.84	0.050 4	0.004 6	0.500 6	0.044 9	0.072 2	0.002 9	213	200	412	30	449	18	
CX1621-05	223	247	0.90	0.053 9	0.004 6	0.522 5	0.053 0	0.070 3	0.003 4	369	194	427	35	438	20	
CX1621-06	364	453	0.80	0.060 8	0.003 3	0.616 1	0.034 7	0.072 9	0.002 6	632	117	487	22	453	16	
CX1621-07	183	359	0.51	0.054 4	0.003 6	0.545 1	0.036 8	0.072 4	0.002 9	391	148	442	24	451	18	
CX1621-08	243	409	0.59	0.065 2	0.003 6	0.835 4	0.045 7	0.092 4	0.003 2	789	119	617	25	570	19	
CX1621-09	626	868	0.72	0.056 9	0.003 0	0.575 7	0.031 5	0.072 6	0.002 7	487	117	462	20	452	16	
CX1621-10	162	297	0.54	0.058 6	0.003 4	0.578 4	0.035 1	0.070 6	0.002 6	550	126	463	23	440	16	
CX1621-11	162	255	0.64	0.062 3	0.006 0	0.607 9	0.058 8	0.070 7	0.002 9	683	208	482	37	440	17	
CX1621-12	283	575	0.49	0.055 2	0.002 8	0.563 2	0.028 4	0.072 6	0.002 5	420	115	454	18	452	15	
CX1621-13	400	1 356	0.30	0.087 3	0.003 9	2.889 5	0.130 9	0.236 3	0.008 0	1 369	86	1 379	34	1 367	42	
CX1621-14	73	173	0.42	0.072 3	0.004 4	2.021 6	0.114 1	0.201 9	0.008 2	994	118	1 123	38	1 186	44	
CX1621-15	230	340	0.68	0.050 7	0.003 5	0.511 9	0.037 2	0.072 5	0.002 7	233	159	420	25	451	16	
CX1621-16	87	185	0.47	0.050 6	0.004 4	0.494 1	0.040 6	0.071 4	0.003 0	233	202	408	28	444	18	
CX1621-17	222	464	0.48	0.050 5	0.002 9	0.514 3	0.029 6	0.073 5	0.002 7	217	135	421	20	457	16	
CX1621-18	186	421	0.44	0.055 3	0.003 4	0.570 3	0.036 0	0.074 8	0.002 8	433	144	458	23	465	17	
CX1621-19	590	712	0.83	0.049 6	0.002 4	0.503 0	0.024 8	0.073 9	0.003 0	176	113	414	17	460	18	
CX1621-20	487	586	0.83	0.056 3	0.002 7	0.568 4	0.029 8	0.073 4	0.002 8	461	107	457	19	456	17	
样品 CX1624																
CX1624-01	318	790	0.40	0.060 3	0.002 7	0.621 5	0.028 8	0.074 7	0.002 5	613	94	491	18	464	15	
CX1624-02	65	89	0.73	0.110 9	0.006 1	4.680 6	0.253 1	0.307 7	0.011 0	1 814	100	1 764	45	1 729	54	
CX1624-03	259	392	0.66	0.064 8	0.003 6	0.672 2	0.040 5	0.074 5	0.002 6	769	112	522	25	463	16	
CX1624-04	112	179	0.63	0.092 0	0.010 4	3.188 8	0.439 2	0.245 8	0.008 6	1 533	219	1 454	107	1 417	45	
CX1624-05	725	659	1.10	0.055 9	0.003 6	0.575 9	0.037 0	0.074 2	0.002 5	456	146	462	24	461	15	
CX1624-06	45	66	0.68	0.052 4	0.013 5	0.508 4	0.117 2	0.071 3	0.004 1	306	500	417	79	444	25	
CX1624-07	74	152	0.49	0.056 7	0.004 7	0.569 0	0.044 4	0.074 2	0.002 6	480	183	457	29	461	15	
CX1624-08	90	147	0.61	0.114 8	0.005 3	5.225 1	0.248 5	0.330 0	0.010 8	1 877	84	1 857	41	1 838	52	
CX1624-09	81	230	0.35	0.055 7	0.003 9	0.560 0	0.038 2	0.073 9	0.002 6	443	157	452	25	460	16	
CX1624-10	116	163	0.72	0.053 7	0.007 9	0.513 0	0.067 5	0.070 6	0.002 9	367	294	420	45	440	18	
CX1624-11	271	730	0.37	0.053 8	0.002 7	0.533 5	0.026 0	0.072 1	0.002 6	365	108	434	17	449	16	
CX1624-12	232	674	0.34	0.056 1	0.002 8	0.539 0	0.027 1	0.069 3	0.002 3	454	109	438	18	432	14	
CX1624-13	312	476	0.66	0.058 6	0.003 2	0.585 8	0.030 8	0.073 0	0.002 4	550	119	468	20	454	15	
CX1624-14	109	323	0.34	0.057 0	0.003 0	0.582 4	0.031 9	0.074 0	0.003 0	500	121	466	20	460	18	
CX1624-15	277	292	0.95	0.055 4	0.006 6	0.534 8	0.063 2	0.069 4	0.002 9	432	267	435	42	432	17	
CX1624-16	181	149	1.21	0.095 0	0.005 1	3.417 7	0.181 0	0.257 0	0.008 7	1 528	102	1 508	42	1 475	45	
CX1624-17	230	408	0.56	0.062 6	0.004 6	0.626 0	0.041 3	0.072 9	0.002 8	694	158	494	26	454	17	
CX1624-18	187	696	0.27	0.055 2	0.003 1	0.586 5	0.032 0	0.075 5	0.002 6	420	121	469	20	469	15	
CX1624-19	135	182	0.74	0.055 7	0.004 7	0.556 2	0.044 3	0.072 3	0.002 7	439	187	449	29	450	16	
CX1624-20	40	163	0.24	0.112 9	0.005 1	5.131 5	0.233 9	0.323 6	0.010 6	1 846	83	1 841	39	1 807	52	

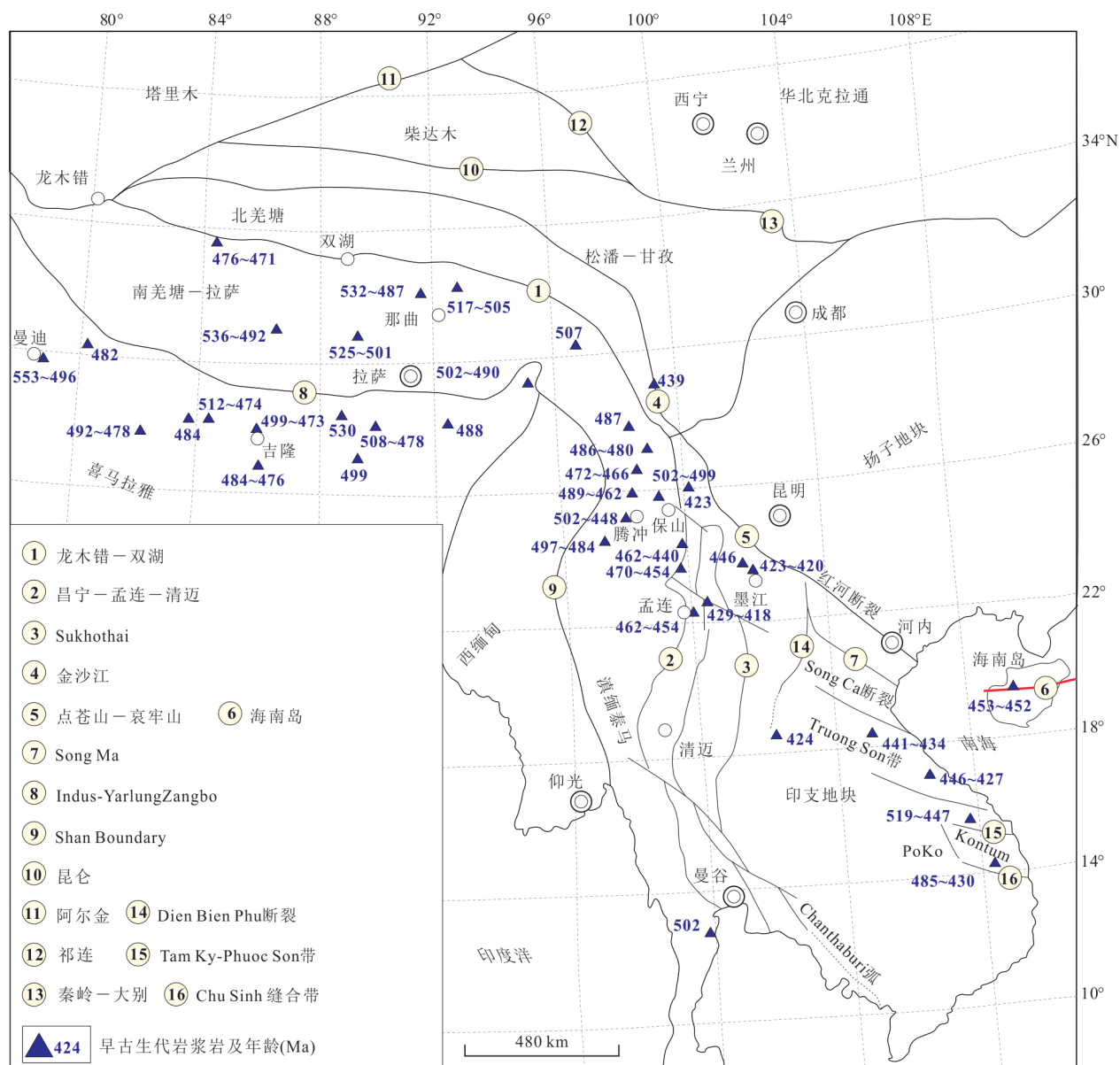


图 5 东南亚原特提斯洋构造格局简图

Fig.5 Sketched tectonic framework of Proto-Tethys Ocean in SE Asia

改编自 Nie *et al.* (2015) 和 Wang *et al.* (2021a); 年龄数据来自 Guynn *et al.* (2012) 等近百篇文献, 不逐一列出

先发生在冈瓦纳大陆西北端(如土耳其和伊朗地区);(2)龙木错—双湖—昌宁—孟连缝合带/构造带记录的大量早古生代岩浆岩多被认为是原特提斯洋俯冲—碰撞过程中的产物,推测龙木错—双湖—昌宁—孟连缝合带/构造带代表了原特提斯洋主洋盆的位置. Jian *et al.* (2009) 在金沙江蛇绿岩中识别出了 439 ± 4 Ma 的斜长角闪岩捕虏体; Liu *et al.* (2018) 在点苍山—哀牢山构造带报道了 423 Ma 花岗岩闪长岩、446~420 Ma 玄武岩和 435 Ma 安山岩,且这些岩石分别具有 I 型花岗岩、富铌玄武岩和高镁安山岩的地球化学特征,具有此种地球化学

特征的岩石组合往往代表了与俯冲作用相关的构造背景;Wang *et al.* (2021a, 2021b) 在印支地块的 Truong Son、Tam Ky-Phuoc Son 和 Kontum 地区厘定出了一系列早古生代(450~423 Ma)岩浆岩,包括具有 N-MORB 地球化学特征的辉长岩和埃达克质岩石,指示为岛弧环境;本文在海南岛西北部识别出了~452 Ma 安山岩(图 3),它们在时空上与金沙江—点苍山—哀牢山构造带、Truong Son、Tam Ky-Phuoc Son 和 Kontum 等地区分布的早古生代岩浆岩具有耦合性;此外, Fergusson and Colquhoun (1996) 通过沉积岩石学研究认为海南岛早古生

代火山-碎屑沉积岩与 Tasmania 地体早古生代沉积建造具有相似性,指示为一个与俯冲作用有关的岛弧环境.因而,我们推测金沙江-点苍山-哀牢山构造带、Truong Son、Tam Ky-Phuoc Son、Kontum 地区和海南岛可能代表了原特提斯洋的支洋盆(图 5),海南岛在晚奥陶纪世(~ 452 Ma)已处于向冈瓦纳大陆北缘汇聚阶段.

5 结论

海南岛西北部安山岩的锆石 U-Pb 年龄为 452 ± 8 Ma、 453 ± 8 Ma,形成于早古生代,它们在时空上与扬子地块西南缘的金沙江-点苍山-哀牢山构造带以及印支地块 Truong Son、Tam Ky-Phuoc Son、Kontum 等地区分布的早古生代岩浆岩具有耦合性.金沙江-点苍山-哀牢山、Truong Son、Tam Ky-Phuoc Son、Kontum 等地区和海南岛可能代表了原特提斯洋的支洋盆.

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