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# 陆源与火山物质的向海输送过程及其控制机制

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摘 要:海洋沉积物的源-汇过程是连接陆地风化、海洋动力系统和全球气候变化的关键纽带,对重建古环境演化具有重要意义.本文综述了陆源与火山物质向海洋的输送过程及其控制机制.陆源物质输运受岩性-气候-海平面-洋流系统共同调控——源岩性质和气候条件通过控制风化作用决定沉积物的产量和理化性质,海平面变化主导沉积物的输送距离,洋流格局决定最终沉积分布.火山物质的输入则受火山活动强度、气候、水文及区域构造背景的多元控制.近年来,地球化学与矿物学示踪技术的发展提升了物源识别能力,但该领域仍面临从定性描述到定量解析的方法学挑战.未来研究需进一步发展多学科交叉方法,以深化对海洋沉积源-汇系统演化规律的认识.

关键词:源-汇系统;河流沉积;风尘沉积;冰川沉积;火山物质沉积;沉积学;气候变化.

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## Transport Processes and Control Mechanisms of Terrigenous and Volcanic Materials to the Ocean

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Abstract: The source-to-sink process in marine sediments is a critical link connecting terrestrial weathering, oceanic dynamics, and global climate change, playing a vital role in reconstructing paleoenvironmental evolution. This paper reviews the transport processes of terrestrial and volcanic materials to the ocean and their control mechanisms. The transport of terrestrial materials is regulated by the lithology-climate-sea level-current system: the nature of the source rocks and climate conditions determine the output and physicochemical properties of the sediments by controlling weathering; sea level changes dominate the distance of sediment transport; and the current pattern determines the final distribution of the sediments. The input of volcanic materials is controlled by the intensity of volcanic activity, climate, hydrology and regional tectonic context. Recent advances in geochemical and mineralogical provenance techniques have significantly improved the identification capacity of sediment sources. However, this field still faces methodological challenges in transitioning from qualitative assessments to quantitative reconstructions. Future research should focus on developing integrated, multidisciplinary approaches to enhance our understanding of the evolution of marine sedimentary source-to-sink systems.

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Key words: source-to-sink system; fluvial deposits; aeolian deposits; glacial deposits; volcanic deposits; sedimentology; climate change.

## 0 引言

海洋沉积物源-汇过程是连接陆地风化、海洋动力系统以及全球气候变化的关键纽带,其研究对于揭示全球碳循环(Leithold et al., 2016; Tao et al., 2023)、古环境演化(Wan et al., 2017; 石学法等, 2023; Yu et al., 2024)以及矿产资源勘探(Garba and Mustapha, 2024)等具有重要的科学意义.在全球气候变化日益加剧的背景下,如何准确预测未来气候变化趋势并制定合理有效的适应性策略成为了地球系统科学的前沿课题,而深入理解古气候、古环境的演化历史是准确预测未来气候变化的前提.近年来,随着地球化学、矿物学等示踪方法的进步以及多学科交叉方法的突破,显著提升了物源识别精度以及物源过程解析能力,推动该领域从定性描述向定量分析发展.

海洋沉积物源-汇过程受多种因素协同控制. 气候系统主要通过控制风化速率和河流输沙量等 影响陆源沉积物供给.温暖潮湿的条件有利于原 始矿物化学分解为风化产物,导致化学风化程度更 高(Kump et al., 2000; West et al., 2005; Clift et al., 2020). 降水是物理侵蚀强度和河流沉积物输 送的关键控制因素,比如最近的研究表明在以高山 短源河流为主的新几内亚北部,降水增加提高了物 理侵蚀速率,导致流域土壤停留时间和沉积物储存 时间减少,从而减少了沉积物化学风化的时间,这 一过程促进了化学风化程度较低的碎屑矿物的输 送(Yu et al., 2024).此外,热带辐合带、ENSO(厄 尔尼诺-南方涛动, El Niño-Southern Oscillation)、 印度洋偶极子等气候因子也会通过影响季风或降 水等方式来影响海洋沉积物的源-汇过程(Yu et al., 2016; Kang et al., 2024; Tang et al., 2024).

海平面波动对陆架暴露程度与沉积物输运路径具有重要的控制作用,一方面海平面升降会影响河口以及海岸线位置,进而影响沉积物输运和沉积的过程(Boulay et al., 2007; Huang et al., 2011).另一方面,海平面升降会通过控制陆架暴露程度,进而影响陆架沉积物的风化和侵蚀(Wan et al., 2017; Xu et al., 2018).比如,相对于海平面较高的时期,低海平面时期,河口向海迁移,有利于河流将大量沉积物输送至拉克希米盆地,沉积物粒度相对

更粗(Cai et al., 2022).洋流与地形耦合则决定了沉积物的再分配模式与堆积中心位置,比如晚第四纪的海平面变化和黑潮路径主要决定了冲绳海槽陆源沉积物的扩散和沉积(Dou et al., 2012);黑潮入侵的增强可能是晚第四纪台湾岛沉积物向南海东北部输送的主要原因(Zhang et al., 2022);印度沿岸流的季节性转向控制了印度半岛南端沉积物的冷暖期变化(Goswami et al., 2012; Kang et al., 2024).此外,构造活动,如青藏高原隆升(Clift et al., 2004; Song et al., 2024),通过改变流域侵蚀速率与物源区格局,对沉积通量产生长期影响.

源-汇过程复杂的控制机制决定了不同海域的源-汇系统显示出显著的空间异质性.边缘海系统(如南海(Liu et al., 2016; Cao et al., 2023)、孟加拉湾(Li et al., 2017; Liu et al., 2019)、阿拉伯海(Cai et al., 2022; Fang et al., 2024)等)以陆源碎屑输入为主导,其物源解析需聚焦临近河流及区域环流特征;而重建开阔大洋物源演化则需重点关注风尘输送、火山活动、生物源以及自生组分的贡献;对于高纬海区(如北冰洋(Meinhardt et al., 2016; Dong et al., 2024))沉积物来源则需要重点考虑冰川作用的贡献.这种空间分异特征要求人们在追踪物源时需针对不同海区的环境特征进行差异化分析,进而更准确地解析沉积物"从源到汇"的过程.

由于海洋沉积物源-汇过程受气候、海平面变化、构造活动和洋流演化等多因素的协同控制,使得海洋沉积物不仅是这些过程的产物,也是记录其演化历史的天然档案.通过分析沉积物的地球化学组成、矿物学特征和沉积结构等信息,可以高精度重建古气候波动、海平面升降、构造活动以及洋流格局的长期演化序列,为理解地球系统演变提供关键证据.本文将聚焦河流、风尘、冰川、火山等主要海洋沉积物来源类型,系统解析沉积物向海输送的过程,并阐明不同环境背景下沉积物源-汇过程的主要驱动机制,为海洋沉积物源-汇过程研究提供理论框架.

## 1 陆源物质输入及其控制机制

陆源沉积物,主要包括石英、长石、云母等矿物碎屑.陆源沉积物输入是指陆地岩石经化学风化和

#### 表 1 全球陆源沉积物向海洋输送通量的估算

Table 1 Estimates of global fluxes of terrestrial sediment transport to the ocean

搬运机制	全球通量(Gt/a)
河流输入	25
风尘输入	0.7
冰川输入	2

注:表中数值据Syvitski et al. (2003)以及其中所引用的参考文献.

物理剥蚀等作用产生的碎屑物质,经河流、风以及冰川等外力搬运至海洋,并在洋流的作用下在海洋发生进一步搬运和沉积.陆源沉积物从源到汇的过程主要受到源区岩性、气候条件、海平面变化以及洋流强度/方向的控制.根据搬运方式,陆源沉积物输入类型可分为河流物质输入、风尘输入以及冰川物质输入,各类型陆源物质输入特征及其控制机制各有特点.

#### 1.1 河流输入及其控制机制

全球河流每年向海洋输送的沉积物通量约占海洋接收陆源沉积物总量的95%(Syvitski et al., 2003),是陆源物质向海洋输入的最主要方式(表1),河流向海洋输送沉积物的过程如图1所示.这些沉积物主要由碎屑颗粒(如石英、长石和黏土矿物)、有机质(如陆源植物碎屑和土壤有机碳)以及溶解态组分(如Si、Fe、Al等元素)构成.进入海洋的河流沉积物,是在海洋周围不同的气候、构造和岩性等环境中形成的(Liu et al., 2016),这些环境因素的变化控制了沉积物的物理和化学性质.

不同岩性岩石的抗风化能力不同,使得岩性成

为源区沉积物产量的关键控制因素.酸性岩的抗风化能力相对基性岩更强,化学风化速率相对较慢(Bufe et al., 2022),这种特性直接影响着源区沉积物的供给强度.比如由于印度—太平洋交汇区岛屿岩性以中基性火山岩为主,地处热带,降水量大,水系以高山短源河流为主,易于风化、剥蚀和搬运(Yu et al., 2023).因此,印度—太平洋交汇区岛屿面积虽小,但其河流入海沉积物量巨大,研究表明这些岛屿约占陆地总面积3%,但是其河流沉积物入海通量约占全球36%(Milliman and Farnsworth, 2011).相比之下,在气候条件一定时,花岗岩流域因岩石抗风化能力更强,其沉积物产出效率显著低于玄武岩流域(Ibarra et al., 2016).

不同气候类型通过降水、温度等因素显著调控着河流向海洋的沉积物输送过程.在现有海洋沉积研究中,季风气候和热带雨林气候是探讨河流沉积物输入时受关注较多的两种气候类型,因此,笔者重点针对这两种气候类型对河流沉积物输送的影响机制展开讨论.在年际时间尺度上,季风气候区降水主要集中在夏季(Fazal et al., 2023),风向在夏季由海洋吹向陆地,冬季由陆地吹向海洋.在地质时间尺度上,主要表现为暖期夏季风较强,季风降水多;冷期夏季风较弱,季风降水少(Sarim et al., 2023; Ning et al., 2024).季风气候的这种明显的冷暖期变化特征,主要通过控制源区的剥蚀风化强度以及河流径流量来影响河流向海洋的沉积物输送过程.一般情况下,夏季风较强时,高温多雨,剥蚀和风化速率高,河流径流量大,可向海洋输送更多

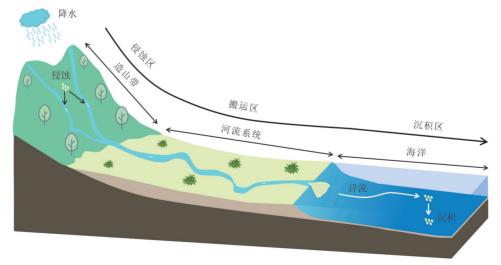


图 1 河流物质输入示意(参考自 Tofelde et al., 2021)

Fig.1 Schematic diagram of fluvial material inputs (referenced from Tofelde et al., 2021)

陆源沉积物(Clift and Jonell, 2021),比如黄河(Xu et al., 2025)、印度河(Clift and Jonell, 2021)等主要 河流均呈现出此特点.值得注意的是,并非所有河 流沉积物输入量都呈现出夏季风盛行期增加,冬季 风盛行期减少的规律.比如位于南冲绳海槽的岩心 记录表明,由于在夏季风减弱时,锋面系统南撤,导 致中国南部降水量增加,源区剥蚀风化增强,河流 净流量增加,从而导致陆源物质输入比例增加;反 之,在夏季风增强时,其陆源物质输入比例减少(陈 金霞等, 2009). 对于热带源区而言, 当发生冷暖期 的气候变化时,其温度和降水同样会发生变化, 该变化控制着源区的剥蚀、风化速率以及河流径 流量,进而控制了河流向海洋的沉积物输送.但 与季风气候不同,即使是在冷期,热带源区的剥 蚀和风化速率可能依旧处于一个较高的水平,河 流径流量可能也一直较大,所以其冷暖期的剥 蚀、风化变化特征可能与季风气候区不同.比如, 对于热带新几内亚岛鸟头半岛北部而言,由于暖 期降水量大,物理侵蚀速率快,并且沉积物被快速 输送至海洋,导致沉积物化学风化的时间较短,进 而导致沉积物在暖期化学风化程度更低,而在 冷期化学风化程度相对更高(Yu et al., 2023).

海平面变化主要通过控制大陆架的裸露来影 响河流向海洋输送沉积物的情况. 宽广大陆架受 海平面影响较大,在高海平面或海平面上升期间, 随着大陆架逐渐被海水覆盖,陆源碎屑物质被向 海搬运的过程中会在陆架区沉积(Prins and Postma, 2000; Kessarkar et al., 2003); 在低海平面或 海平面下降时,大陆架裸露,河流入海口向海延伸 (Huang et al., 2011), 沉积中心向海迁移(Li et al., 2020a).一方面,这会导致裸露陆架上的沉积 物再次被剥蚀、风化并搬运,最终输送至海洋并沉 积(Wan et al., 2017);另一方面,岸线和河口的向 海迁移会导致粒度带整体向海推移,致使同一地 理位置出现沉积物粒度明显粗化的现象(Boulay et al., 2007). 二者共同作用,最终控制着河流输 入沉积物的物理和化学性质,例如巽他陆架 (Hanebuth et al., 2002). 但是,需要特别指出的 是,这种海平面变化对沉积物输送的调控效应主 要集中在具有宽广大陆架的海域,而对于日本岛 周边(Zhao et al., 2020)和新几内亚岛北部海域 (Tang et al., 2024)等陆架狭窄的区域,河流输送 沉积物的特征受海平面波动的影响则相对有限.

在沉积物进入海洋之后,主要随洋流被进一步的搬运.洋流的方向和强度控制了河流沉积物入海后的最终分布和特征.尤其是受季风控制季节性转向的洋流,比如沿岸流.洋流的转向控制了位于洋流路径区域的沉积物来源和特征的季节性变化,在地质时间尺度上可能体现为冷暖期的变化(万世明等,2008; Goswami et al.,2012; Kang et al.,2024).而洋流的强度对沉积物具有重要的分选效应,当洋流搬运能力较强时,沉积在某一位置的沉积物粒度相对可能会更粗;反之,沉积物粒度相对会更细(Wang et al.,2015; 宋震等,2024).

#### 1.2 风尘输入及其控制机制

海洋风尘沉积是指陆源物质通过大气动力搬运后最终在特定海区沉降形成的一类沉积物,其源-汇过程如图 2 所示.其主要来源于干旱-半干旱地区的表层风化物质,在风力作用下被抬升进入大气,并随全球环流系统进行远距离传输.当风速达到临界阈值时,在土壤湿度低、植被少和可侵蚀沉积物供应充足的地区,风尘沉积物就会被释放出来(Maher et al., 2010; Rowland, 2021).这类沉积物主要由小于 10 µm 的细小颗粒组成(Rea and Hovan, 1995; Rea et al., 1998).全球每年约有 2 000 Mt风尘进入大气循环系统,其中约四分之一最终沉降在海洋环境(Shao et al., 2011),尽管其通量仅相当于河流输入量的 5%(Duce et al., 1980),但在远离大陆的远洋深水区,风尘沉积却成为深海沉积物的主要来源(万世明和李安春, 2004; Serno et al., 2014).

风尘输入也主要受岩性、气候和洋流等多重因 素的协同控制.源区岩性和洋流对风尘输入的控制 机制与上述河流物质输入类似,这里不再重复.海 平面变化导致的大陆架裸露并非风尘输入的主要 影响因素.在地质时间尺度上,冷暖期的气候变化 是海洋风尘输入的重要控制因素.冰期较高的干旱 程度、较少的植被以及较强的风,导致风尘输入量 更高(Muhs, 2013; Wan et al., 2020). 冰期全球的 风尘通量比间冰期高 2~5 倍(Maher et al., 2010; Wan et al., 2020). 降水和风等气候因子是影响风尘 输入的关键因素.其中风的强度决定了风尘的粒 度,源区的干旱程度决定了风尘通量(Rea et al., 1985; Rea, 1994; Wan et al., 2020). 风力越强,风 尘的粒度一般会更大(Stuut et al., 2014). 随着风尘 源区干旱程度的增加,风尘通量一般也会相应增加, 比如,菲律宾海西部风尘输入的增加与东亚冬季

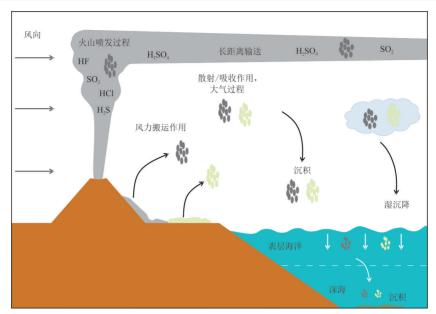


图 2 风尘及火山物质输入示意

Fig.2 Schematic diagram of dust and volcanic material input

图改自 Langmann (2013);图中灰色颗粒物质代表火山喷发的碎屑物质,指示火山物质源-汇过程;淡黄色颗粒物质代表陆源碎屑物质,指示风尘物质源-汇过程

#### 表 2 主要源区现代年平均风尘通量模拟结果对比

Table 2 Comparison of modern annual average eolian flux simulation results for major source regions

			_					-	
文献	非洲		亚洲		美洲		Ma II fil m	4-r A	
	北部	南部	阿拉伯	中部	东部	北部	南部	澳大利亚	全球
Tanaka and Chiba, 2006	1 087	63	221	140	214	2	44	106	1 877
	(58%)	(3%)	(12%)	(7.5%)	(11%)	(0.1%)	(2%)	(6%)	
Werner <i>et al.</i> , 2002	693		101	96				52	1 060
	(65%)		(9.5%)	(9%)				(5%)	
Luo et al., 2003	1 114		119		54			132	1 654
	(67%)		(7%)		(3%)			(8%)	
Zender <i>et al.</i> , 2003	980		415			8	35	37	1 490
	(66%)		(28%)			(0.5%)	(2%)	(2.5%)	
Ginoux et al., 2004	1 430		496			9	55	61	2 073
	(69%)		(24%)			(0.4%)	(3%)	(3%)	
Miller et al., 2004	517		43	163	50	53		148	1 019
	(51%)		(4%)	(16%)	(5%)	(5%)		(15%)	

注:通量数值据 Maher et al.(2010),单位: Tg/a; 括号内为全球占比.

风强度和亚洲大陆干旱程度的增加基本一致(Wan et al., 2012; 孙军等, 2025),而源区降水的增加会导致输送至菲律宾海的风尘通量相对较低(Wang et al., 2020).但需要注意的是,长时间持续极端干旱可能会因源区风尘物质供应匮乏而导致通量减少(万世明和李安春, 2004).

北非、东亚和中亚以及阿拉伯的干旱地区是三个主要的现代沙尘源区(表2),均位于北半球(Rea, 1994; Maher *et al.*, 2010).从全球范围来看,太平

洋、北大西洋和西北印度洋是风尘沉积的主要海区(于兆杰,2013).中亚东部的半干旱和干旱地区是太平洋最大的风尘来源,来自该区域在中国下风方向的风尘沉降率超过1000 mg/(cm²•ka)(Rea,1994).但是,除了中亚东部的半干旱和干旱地区,太平洋也会接收来自其他区域的风尘,比如东南信风可以将秘鲁和智利的半干旱至极度干旱地区的沉积物输送至太平洋东南部(Saukel et al.,2011),澳大利亚东南部和新西兰的沉积物会通过

风被输送至太平洋西南部(Stancin et al., 2008).

撒哈拉沙漠向大西洋输送的沙尘总量约为 (970±325) Tg/a, 是大西洋的主要沙尘源(Rowland, 2021). 大西洋表层沉积物中Ti/Al证据表明, 北非地区的风尘是加那利盆地陆源物质的主要来 源(Govin et al., 2012). 在千年时间尺度上,非洲湿 润期、海因里希事件和新仙女木事件等对北非风尘 通量具有显著影响,非洲西北边缘 20 ka 以来的风 尘通量记录显示,在海因里希事件和新仙女木事件 期间风尘通量达到峰值,而在非洲湿润期风尘通量 显著降低(McGee et al., 2013).24万年以来,撒哈 拉沙漠风尘通量的变化主要受岁差周期驱动,与北 半球夏季日照变化高度一致,而受冰期-间冰期旋 回的影响较弱(Skonieczny et al., 2019). 但是,撒哈 拉沙漠并非大西洋风尘沉积的唯一来源.例如位于 北大西洋的 U1313 站位的放射性成因同位素(Sr、 Nd、Pb)证据表明,上新世期间其陆源沉积物主要为 来自北美中纬度地区的风尘(Lang et al., 2014).

东阿拉伯海通过西南季风的湿沉降(降水) 接收了大量的风尘物质(Ramaswamy et al., 2017; Suresh et al., 2021). 地球化学证据表明, 阿拉伯半岛对季风开始(6月)和季风中期(7~8 月)的风尘贡献显著,非洲东北部也在季风中期 贡献了部分风尘物质,亚洲西南部和塔尔沙漠贡 献了季风末期(9月)的风尘物质(Suresh et al., 2021). 西赤道印度洋的 Sr-Nd 同位素证据表明, 其陆源风尘主要来源于阿拉伯半岛(Zhou et al., 2024). 北半球千年时间尺度气候振荡通过影响 季风强度调控了阿拉伯海的风尘输送,海因里希 事件期间风尘输入量减少, Dansgaard-Oeschger期 间夏季风强度增强,风尘输入量增加(Sebastian et al., 2023). 间冰期,夏季降水量增加,源区植 被覆盖扩张,阿拉伯海的风尘供应减少;冰期,高 纬度冰盖扩张导致全球温度梯度增加,季风减弱, 副热带西风急流更偏南、更冷和下沉,源区干旱 加剧,风尘供应又急剧增加(Zhou et al., 2024).

从陆地向海洋的风尘输入给海洋提供了大量浮游植物生长所必需的营养元素(Wan et al., 2020),对海洋生产力以及海洋生物地球化学循环具有重要影响(Jickells et al., 2005).南大洋和太平洋的沉积记录显示风尘铁通量与海洋生产力之间存在很强的相关性(Han et al., 2011; Martínez-García et al., 2014).海洋生产力

的变化又可能会对碳循环产生影响(Han et al., 2011).比如,在寒冷干旱的气候条件下,风尘通量的增加和上升流的增强,会导致大陆边缘营养物质增加,初级生产力升高,进而导致海洋缺氧,促进有机碳保存(Zhai et al., 2018).

## 1.3 冰川物质输入及其控制机制

冰川物质输入是极地和高纬度海域沉积物的重要来源,其主要通过冰山漂移、冰筏碎屑释放以及冰川融水携带的细颗粒物质等方式进入海洋环境,其源-汇过程如图 3 所示.例如,Xiao et al.(2024)基于北极楚科奇海边缘沉积岩心的分析认为,其沉积物主要来自东西伯利亚冰盖、劳伦泰德冰盖等的冰川输入,沉积物主要包括粗粒冰筏碎屑以及与融水相关的沉积物等.在晚更新世期间,北美北极地区的冰盖和欧亚大陆北部的冰盖崩裂了大量的冰山,这些冰山通过弗拉姆海峡漂流到格陵兰海,沉积了氧化铁颗粒和粗粒岩屑等(Darby et al., 2002).

冰川物质输入主要受气候、洋流以及冰川动 力学等的共同影响. 气候变暖会加速冰川消融, 更多冰川融水进入海洋;同时,冰体流动性增加 甚至解体,从而将冰川沉积物输送至海洋(黄晓 璇等, 2018; 张鑫悦等, 2023). 洋流系统将海冰 及其中的沉积物从冰架上搬运到整个北冰洋,海冰 漂移构成了从北极沿岸和陆架向整个海盆输送碎 屑物质的主要途径(Meinhardt et al., 2016). 洋流强 弱和方向的变化还会控制冰川物质在海洋的分布 及特征.比如,北极群岛的碳酸盐岩碎屑被冰山和 海冰以冰筏碎屑的形式在波弗特环流(Beaufort Gyre)的作用下被输送至北冰洋并沉积(Meinhardt et al., 2016). 间冰期较强的波弗特环流可以将携带 冰筏碎屑的冰山等搬运至西北冰洋并沉积,而冰期 较弱的波弗特环流无法将粗的冰筏碎屑搬运至西 北冰洋(黄晓璇等, 2018).冰川动力学决定了沉积 物的侵蚀效率和物质输运量.在末次冰期旋回中, 湿底型的白令冰川至少三次前进至大陆架边缘,以 极高平均速率(可达 10 m/ka)向大陆架和大陆坡输 送了至少 925 km³的冰川沉积物 (Montelli et al., 2017).冰盖消退或崩解可能会导致大量海冰或冰山携 带沉积物进入深海(Kaparulina et al., 2016). 这些 控制机制的相互作用使得冰川物质输入在千年 至轨道时间尺度上呈现显著的变异性,并为重 建古气候和冰盖演化提供了关键沉积学指标.

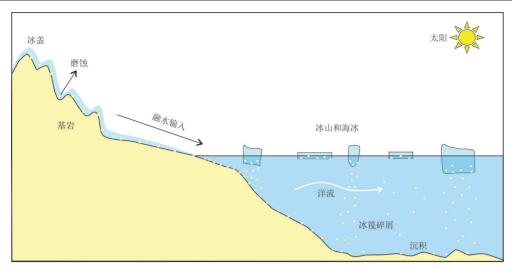


图 3 冰川物质输入示意(参考自 Kaparulina et al., 2016)

Fig.3 Schematic diagram of glacial material input (referenced from Kaparulina et al., 2016)

## 2 火山物质输入及其控制因素

火山沉积物可分为陆上火山喷发物质和洋中 脊或海底火山喷发物质.本文重点讨论来自陆地 火山活动的物质.火山物质指的是由火山活动产 生的各类碎屑物质,包括火山灰、火山玻璃及火山 碎屑岩等,其形成与搬运过程构成了一个复杂的地 球表层物质循环系统,其源-汇过程如图2所示.这 些物质主要通过火山喷发直接进入大气圈或水体 环境,随后经历物理化学风化、搬运和沉积等一系 列地质作用.火山碎屑颗粒的形态特征可提供有 关破碎过程和形成环境的重要信息(Carey and Schneider, 2011). 火山灰是由岩浆和火山口围岩 物质的破碎作用形成的,平均粒径小于2 mm (Zimanowski et al., 2003),其成分包括细小的火山 玻璃颗粒、矿物碎片等.火山玻璃是火山喷发出的 熔岩在迅速冷却过程中未能及时结晶而形成的玻 璃质结构岩石.它通常缺乏规则的形状,可能呈现 珍珠状、气孔状或不规则状.Bao et al.(2023)在南 海东部深海盆地发现的火山玻璃以棕色为主,伴生 少量白色和无色类型,表明安山岩可能是主要 源岩,其中白色火山玻璃具有典型的流动结构.

在喷发过程中,火山物质被喷射至不同高度的大气层,强烈的火山喷发甚至能将火山灰输送至平流层(Robock, 2000).这些悬浮的火山碎屑随后在大气环流系统的驱动下进行扩散传输,最终通过干沉降(重力作用)或湿沉降(随降水)过程从大气中移除(Brown et al., 2012),并在海洋和

陆地等地表环境中发生沉积.值得注意的是,沉降 在陆地表面的火山碎屑会经过风化侵蚀作用,随 后被河流、冰川等外营力二次搬运至海洋环境.

火山碎屑物质经历的多阶段搬运与沉积过程, 使得海洋沉积物中的火山物质不仅保留了原始喷 发的特征信息,同时也承载了后期地质改造的印 记.这一完整的物质循环系统的运行效率受到火 山活动强度、气候条件以及洋流等多重因素的综合 影响与调控.火山活动越强,向海洋输送的火山物 质可能会越多.以南海东部深海沉积岩心为例,由 于源区在间冰期的火山喷发强度显著高于冰期,这 导致了沉积物在间冰期呈现ε<sub>Nd</sub>值偏高而<sup>87</sup>Sr/<sup>86</sup>Sr 比值偏低的特征,即火山物质在间冰期相对含量增 加(Bao et al., 2023). 此外, 气候和洋流对火山碎 屑物质搬运与沉积过程的具体控制机制可参见 上一节的详细论述.值得注意的是,火山活动对 气候变化也会有一定影响,比如有研究表明,北 半球长期冷却趋势(始于约3.5 Ma前)使气候系 统接近冰盖形成的临界点,而2.65 Ma前火山活 动的突然增强(叠加北纬高纬度地区的低日照条 件)为冰盖快速扩张提供了最终触发条件,促使 北半球冰期迅速发展(Prueher and Rea, 2001).

火山灰的大量堆积可能会改变沉积物的粒度特征,掩盖原有的沉积物粒度变化规律.火山灰层的粒径由盛行风强度、火山喷发的火山灰粒径以及火山喷发强度等共同决定(Carey and Sigurdsson, 2000).火山活动为海底沉积环境引入了大量富铁火成岩矿物,如橄榄石、辉石、角

闪石、钙铝榴石、黑云母等(Li et al., 2020b). 此外,在海洋环境中,火山玻璃风化会释放出Fe、Mn、Al、Mg、Ca等元素,可能对海洋的地球化学循环具有重大意义(Chen et al., 2014).

大量火山物质进入海洋会影响海洋生态环境,并可能引发全球气候变化以及生物灭绝(Li et al., 2020b).为应对火山活动突发性和不确定性的挑战,建立有效的监测和预警机制至关重要.在监测和预警工作中,需要综合运用地震、气体、地表形变、遥感、无人机及次声波等多种监测技术手段,以实时捕捉火山活动信号;同时构建多源数据融合的预警平台,借助大数据和人工智能优化预警的准确性与及时性.

## 3 地球化学与矿物学示踪技术

Sr-Nd 同位素组成和黏土矿物组成是追踪沉积物来源的可靠指标,已被广泛用于南海(Cai et al., 2020; Jin et al., 2022)、菲律宾海(Seo et al., 2014; Xu et al., 2018)、阿拉伯海(Yu et al., 2019; Kang et al., 2024)、孟加拉湾(Liu et al., 2019)等海域的沉积物来源判别,本文将着重讨论这两种物源示踪方法.

Sr-Nd 同位素主要受源岩岩性和年龄的控制 (Kessarkar et al., 2003),而对地表过程不敏感(Hu et al., 2020; Duan et al., 2023), 所以基于 Sr-Nd 同 位素组成可以追踪沉积物的物源.但是,需要注意 的是,由于Sr比Nd更具流动性且在风化过程中易 从土壤中流失,其同位素组成更易受化学风化程度 控制(Goldstein and Jacobsen, 1987; Kessarkar et al., 2003). 从元素地球化学行为来看, Rb和Sr表现 出显著差异:首先,二者因物理化学性质的差异会 选择性地进入不同硅酸盐矿物.含Sr矿物(如斜长 石和方解石)的化学稳定性通常低于含Rb矿物(如 云母类和钾长石),使其在风化过程中更易分解,从 而导致 Sr 更易从矿物中析出(Feng et al., 2009). 其 次,Sr在岩石化学风化过程中会被释放到溶液中 (Brass, 1975). 因此, 化学风化不仅使矿物粒度减 小,还会引起残留矿物中Rb/Sr比值增大,进而促进 风化产物中<sup>87</sup>Sr/<sup>86</sup>Sr比值增大(Feng et al., 2009).

此外,碎屑沉积物的 $^{87}$ Sr/ $^{86}$ Sr 比值也会受到粒度效应的控制,而 $^{6}$ M则主要反映沉积物来源的变化(Duan *et al.*, 2023; Gao *et al.*, 2025). 这主要与Rb和Sr的不同地球化学性质以及Sm和

Nd的相似地球化学性质有关,由于细粒组分中云母含量较高,而云母中Rb/Sr比值较高,因此细粒组分往往富集放射性成因<sup>87</sup>Sr(Feng et al., 2009).非洲西北部沿岸深海沉积物的研究也表明粒度对Sr同位素组成的影响是显著的,细粒组分比粗粒组分<sup>87</sup>Sr/<sup>86</sup>Sr比值更高,而Nd同位素则基本不受粒度影响(Meyer et al., 2011).

不同岩石类型在风化过程中会形成不同 类型的黏土矿物,因此,特定地区的母岩类型 和风化历史会形成独特的黏土矿物"指纹".伊 利石通常被视为原生矿物,其含量较高则表明 在寒冷干旱的气候条件下,大陆风化作用中的 水解过程减弱,而直接的物理侵蚀作用则相对 增强;蒙脱石是一种次生矿物,由原生铝硅酸 盐和铁镁硅酸盐在温暖潮湿条件下经化学风 化而成;高岭石很容易在单硅铝化土壤中被发 现,主要受大陆水解强度的控制(Liu et al., 2007). 在印度西北喜马拉雅前沿地区的 Pinjor 组,伊利石主要来源于长石和云母的风化;绿 泥石主要来自中基性结晶岩和低级变质岩;高 岭石主要是酸性环境下花岗岩和基性岩的风 化产物(Chaudhri and Singh, 2012). 台湾地区 由于广泛出露第三纪沉积岩,尤其是砂岩、页 岩和板岩,具有较强的物理风化作用,导致河 流沉积物中伊利石和绿泥石富集(Li et al., 2012). 类似地,青藏高原东部隆起的古生代-中 生代沉积岩和少量侵入受挤压的火成岩,构成 了红河和湄公河水系中上游的主要基岩类型, 其黏土矿物组成以伊利石为主,绿泥石和高岭 石次之,蒙脱石含量较少(Liu et al., 2007).

Sr-Nd 同位素与黏土矿物组成作为两种经典且互补的物源示踪技术,在海洋沉积动力学重建及古环境演变研究中具有不可替代的重要价值.需要特别指出的是,沉积体系中的地球化学和矿物学信号往往记录了多因素耦合作用的综合地质信息.然而由于陆源物质输入的复杂性,现阶段相关研究多局限于主要物源的定性判别,而无法精准定量示踪功学物源可能的物质来源.除了地球化学和矿物学物源示踪技术外,还有很多其他的方法,比如有机地球化学标志物等.在未来的研究中,可以多指标耦合分析,并结合主成分分析、机器学习等现代技术,以提高物源判别精度.

## 4 总结与展望

海洋沉积物源-汇过程是涉及物质产生-搬运-沉积的系统性地质过程,根据物质来源可将海洋沉积物划分为陆源物质、火山物质、生物源沉积物、自生沉积物以及宇宙源沉积物等主要类型.在前人研究的基础上,本文着重探讨了陆源和火山物质的输入过程及其控制机制.

陆源物质输入根据搬运方式又可分为河流物 质输入、风尘输入和冰川物质输入,其源-汇过程 主要受源岩性质、气候条件、海平面变化以及洋 流强度/方向控制.具体而言,源岩性质和气候条 件共同决定了源区的风化强度和类型,进而影响 陆源物质的产量及理化性质;海平面变化主要调 控沉积物的搬运距离,由此对沉积物通量、粒度 特征等理化性质产生显著影响;洋流的强度和方 向则决定了陆源物质在海洋环境中的最终沉积 分布格局.除了上述共性控制机制外,各类陆源 物质输入又有其独特的控制机制:降水是河流物 质输入的直接气候控制因素,显著影响着沉积物 的剥蚀过程与河流的搬运能力;源区干旱程度和 风力条件是风尘沉积的主要控制因素,分别主导 沉积物通量和粒度特征;冰川动力学是冰川输入 特有的控制机制,主要控制沉积物的侵蚀效率和 沉积物通量等.火山碎屑物质向海洋的输入过程 则受到火山活动强度、气候条件、洋流及区域构 造背景等多重因素的综合影响与调控.其中,区 域构造背景决定了火山的分布及其喷发频率,火 山活动强度决定了火山物质的供给量,气候条件 和洋流主要控制火山物质的搬运和沉积过程.

深入理解沉积物源-汇过程及其控制机制是 开展海洋沉积与古环境研究的前提,正因受到 各种控制机制的约束,海洋沉积物得以成为记录古环境演化的良好档案.通过"将今论古"与 "以古示今"研究范式的有机结合,古气候环境 重建为理解地球系统演变规律和预测全球气候 变化提供了重要理论支撑.但需注意的是,现存 研究在机制解析层面仍存在双重局限:其一,沉 积响应作为多因素非线性作用的综合产物,现 有方法体系难以实现各控制端元的定量解耦与 贡献度分离;其二,物源示踪技术虽已建立矿物 和地球化学等多种方法,但在多物源混合沉积 区仍面临成因多解性判识难题,现有解译多停 留于定性分析层面.突破上述瓶颈需着力构建多 学科协同研究体系,将传统沉积学方法与数值模 拟等新技术相结合,以推动源-汇过程定量分析.

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