

柴达木盆地早渐新世湖相沉积碳氧同位素组成及其古气候意义

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摘要: 青藏高原新生代以来的隆升显著改变了亚洲的构造地貌格局, 而且深刻影响了东亚及全球的新生代季风环流和气候演变过程。位于青藏高原北部的柴达木盆地是研究新生代气候环境演化历史的重要窗口。湖相沉积作为连续的、高分辨率的环境记录载体, 对区域环境响应十分灵敏, 而其中赋存的自生碳酸盐是古湖泊地质历史时期古气候重建研究的重要载体。本文以柴达木盆地北部大红沟剖面上干柴沟组中的湖相泥灰岩为研究对象, 利用X射线衍射、扫描电镜及能谱分析确定其矿物组成为石英、方解石、伊利石、斜长石、绿泥石、钾长石等。通过泥灰岩碳氧同位素组成相关性分析($r^2=0.04$)以及与区域及全球现代湖泊沉积自生碳酸盐碳氧同位素组成对比, 揭示出柴达木古湖在早渐新世(~ 30.8 Ma)仍然为一开放型淡水湖泊。在此基础上, 基于泥灰岩氧同位素组成重建了柴达木古湖早渐新世(~ 30.8 Ma)时期夏季大气降水的氧同位素值($\sim 6.2 \pm 0.7\text{‰}$)。该值较柴达木盆地现代夏季大气降水及河水 $\delta^{18}\text{O}$ (water-SMOW) 值高 $\sim 2.2\text{‰}$, 指示自早渐新世(~ 30.8 Ma)以来柴达木盆地夏季发生了显著降温, 幅度达 $\sim 6.5^\circ\text{C}$ 。这一降温过程受全球变冷及青藏高原隆升共同驱动, 其中全球变冷起到了主导作用。

关键词: 柴达木盆地; 早渐新世; 湖相自生碳酸盐; 碳氧同位素; 古气候

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Carbon and Oxygen Isotopic Composition of Early Oligocene Lacustrine Sediments in the Qaidam Basin and Their Paleoclimate Significance

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Abstract: The Cenozoic uplift of the Tibetan Plateau has profoundly reshaped the tectonic geomorphology of Asia and deeply influenced the Cenozoic monsoon circulation and climatic evolution of East Asia and globally. The Qaidam Basin, located in the northern Tibetan Plateau, serves a window into studying Cenozoic climatic and environmental evolution. Lacustrine sediments, as continuous and high-resolution environmental records, sensitively record regional environmental fluctuations. Among these, the authigenic carbonates preserved within the sediments provide key records for reconstructing paleoclimatic histories of ancient lakes. This study focuses on lacustrine marlstones from the

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Shangganchaigou Formation in the Dahonggou Section, northern Qaidam Basin. The mineral composition of these marlstones is determined by XRD, SEM and EDS analysis methods, which mainly includes quartz, calcite, illite, plagioclase and chlorite. By examining the correlation between $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ values of the marlstones ($r^2=0.04$) and comparing them with regional and global authigenic carbonates from modern lacustrine sediments, this research reveals that the paleo-Qaidam Lake remains an open freshwater lake during the Early Oligocene (~30.8 Ma). Based on this, the oxygen isotopic composition of summer atmospheric precipitation in the paleo-Qaidam Lake during the Early Oligocene (~30.8 Ma) is reconstructed by using the oxygen isotopic composition of the marlstone ($\sim 6.2 \pm 0.7\text{‰}$). This value is approximately 2.2‰ higher than the modern summer atmospheric precipitation and river water $\delta^{18}\text{O}_{\text{(water-SMOW)}}$ values in the Qaidam Basin, indicating the Qaidam Basin has undergone significant cooling since the Early Oligocene (~30.8 Ma), with a magnitude of approximately 6.5 °C. It is proposed that global cooling and the uplift of the Tibetan Plateau cooperate to drive the significant cooling in the Qaidam Basin since the early Oligocene, with global cooling playing a dominant role.

Key Words: Qaidam Basin; Early Oligocene; Lacustrine-authigenic carbonates; Carbon and oxygen isotopes; Paleoclimate

0 引言

印度板块与欧亚板块的碰撞以及随后青藏高原的形成是地球历史上最重要的地质事件之一 (Molnar and Tapponnier, 1975; 钟大赉和丁林, 1996; Zhang *et al.*, 2010; Ding *et al.*, 2022)。青藏高原新生代的显著隆升不仅使亚洲地区的地理地貌及构造格局发生了重大改变, 而且对新生代东亚区域及全球古气候的变迁都有显著的影响 (Raymo and Ruddiman, 1992; 李吉均和方小敏, 1998; 施雅风等, 1999; An *et al.*, 2001; Guo *et al.*, 2002; 张培震等, 2006; Dupont-Nivet *et al.*, 2007; Fang *et al.*, 2020, 2022; Ding *et al.*, 2022)。柴达木盆地位于青藏高原北部, 新生代地层序列连续清晰, 盆地中的河湖相沉积完整记录了新生代青藏高原北部地区的构造活动历史和古环境演变信息 (Fang *et al.*, 2007; Miao *et al.*, 2011; Song *et al.*, 2013, 2018; Ji *et al.*, 2017; Wang *et al.*, 2017, 2022; Nie *et al.*, 2020; Sun *et al.*, 2020; Cheng *et al.*, 2021; Lu *et al.*, 2022; Liu *et al.*, 2023)。要充分理解柴达木盆地早新生代以来构造隆升与气候之间的耦合关系, 选择关键载体尤为关键。作为一个区域的物质交汇点, 湖泊在广泛地接受周缘源区的物质供给的同时, 其自身较高的沉积速率使其在较短时间内就能记录丰富的沉积信息 (汪品先和刘传联, 1993; 沈吉, 2009; 金章东, 2011), 从而使得连续的、高分辨率的湖相沉积对区域环境响应十分灵敏, 而其中赋存的自生碳酸盐岩更是古湖泊地质历史时期古气候和古地理重建研究的重要载体 (刘传联, 1998; Dettman *et al.*, 2003; 吴珍汉等, 2007; 伊海生等, 2007; Kent-Corson *et al.*, 2009; 蓝江湖等, 2013; Han *et al.*, 2014; 张成君等, 2016; 鹿化煜等, 2018; Yang *et al.*, 2022; Zhang *et al.*, 2024)。

基于此, 本研究以柴达木盆地北部大红沟剖面上干柴沟组中的湖相泥灰岩为研究对象, 利用 X 射线衍射、扫描电镜及能谱分析确定其矿物组成, 并开展碳氧稳定同位素分析研究。通过泥灰岩碳氧同位素组成相关性分析以及与区域碳酸盐岩基岩、现代湖泊沉积自生碳酸盐碳氧同位素组成的对比研究, 以揭示柴达木古湖在早渐新世时期的古水文特征, 进而探讨柴达木盆地自早渐新世以来古气候变迁过程及其驱动机制。

1 区域地质背景

1.1 区域地质概况

柴达木盆地位于青藏高原北部, 是该区最大的陆相沉积盆地 (图 1a)。新生代以来受印度板块与欧亚板块碰撞的远程效应影响, 柴达木盆地在周缘三大主控断裂带即阿尔金断裂、东昆仑断裂及祁连山-柴北缘断

裂控制下发生显著沉降, 接受了巨厚的新生代沉积, 盆-山耦合效应十分显著 (Yin *et al.*, 2008; Cheng *et al.*, 2021; Zhang *et al.*, 2025)。现今的柴达木盆地是亚洲内陆地区最为干旱的地区之一, 其年平均气温为 -4°C , 大部分地区年均降雨量小于 100 mm (Zeng *et al.*, 2024)。

依据地层岩性、接触关系、古生物化石组合、磁性地层学等特征, 柴达木盆地新生代地层序列由老到新依次为: 路乐河组 (E_2^1l)、下干柴沟组 (E_2^2xg)、上干柴沟组 ($E_3N_1^1sg$)、下油砂山组 (N_1^2xy)、上油砂山组 (N_1^3sy)、狮子沟组 (N_1^3sh) 和七个泉组 ($Q_{1-2}q$) (青海省地质矿产局, 1991; Fang *et al.*, 2007, 2019; Yin *et al.*, 2008; Lu and Xiong, 2009; Zhang *et al.*, 2010; Zhuang *et al.*, 2011a; Chang *et al.*, 2015; Ji *et al.*, 2017; Cheng *et al.*, 2021)。需要指出的是, 近年来研究者通过对柴达木盆地北部地层开展高精度磁性地层学研究, 将路乐河组沉积的起始年龄限定在 $\sim 30\text{ Ma}$ 或者 $\sim 20.5\text{ Ma}$ (Wang *et al.*, 2017, 2022; Nie *et al.*, 2020), 这与之之前 $\sim 52\text{ Ma}$ 的认识相差甚大, 从而造成了柴达木盆地新生代地层“传统年龄模型”与“新年龄模型”的争议 (Cheng *et al.*, 2021; Lu *et al.*, 2022; Wang *et al.*, 2022; 傅志明等, 2025; 宋博文等, 2025)。而最近学者们开展的碳酸盐岩 U-Pb 定年结果对柴达木盆地新生代地层年代格架依旧存在不同认识 (Ge *et al.*, 2025; Zeng *et al.*, 2025)。考虑到路乐河组初始沉积年龄偏老的“传统年龄模型”与通过地震资料、生长地层和热年代学等地质证据得到的盆地内变形历史更为吻合 (Cheng *et al.*, 2021), 同时大红沟剖面轮藻化石生物地层学研究结果与先前对大红沟剖面开展的高精度磁性地层学年龄格架相一致 (Xing *et al.*, 2023), 特别是最近有学者在柴达木盆地北部红沟剖面路乐河组地层中获得了 $45.0 \pm 2.8\text{ Ma}$ 的火山凝灰质砂岩最年轻锆石年龄 (Zeng *et al.*, 2025), 本研究依旧采用柴达木盆地新生代地层“传统年龄模型”。

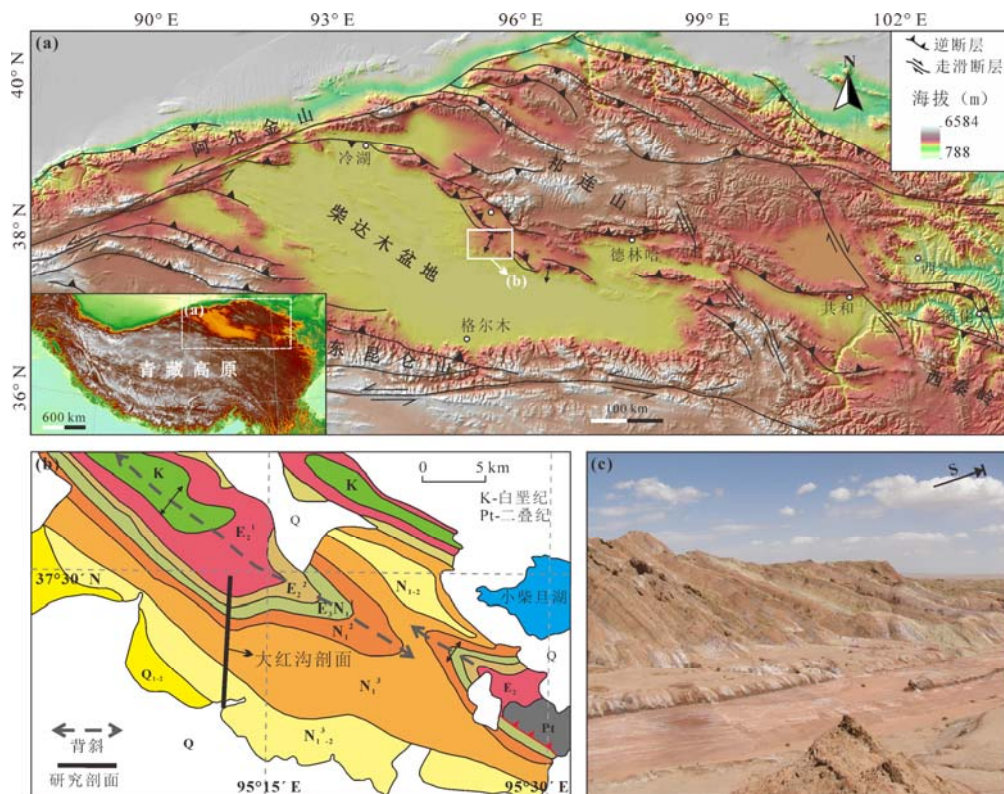


图1 a. 柴达木盆地地理位置 (修改自 Ji *et al.*, 2017); b. 研究区地质简图及大红沟剖面位置 (修改自 Ji *et al.*, 2017); c. 大红沟剖面下段地层野外露头

Fig.1 a. Geographical location of the Qaidam Basin (modified after Ji et al., 2017); b. The simplified geologic map of the study area and location of the Dahonggou section (modified after Ji et al., 2017); c. Field outcrop of the stratum in the lower member of the Dahonggou section

1.2 研究层位及地层时代限定

大红沟剖面（地理坐标：37°29'11"N，95°12'42"E）位于柴达木盆地北部，大红沟背斜南翼（图 1b），出露的地层自老到新依次为路乐河组、下干柴沟组、上干柴沟组、下油砂山组、上油砂山组和狮子沟组。剖面中上干柴沟组岩性以灰绿色—紫红色细砂岩、泥质粉砂岩和泥岩为主，含有少量薄层石膏、泥灰岩夹层。其中砂岩可见平行层理、板状及楔状交错层理、滑塌构造发育，泥岩、泥质粉砂岩则发育水平层理（图 2），为典型的湖泊三角洲-浅湖相沉积（Song *et al.*, 2013; Ji *et al.*, 2017）。此外，先前在大红沟剖面上干柴沟组下段 135 层湖相地层中发现了大量保存完好的植物大化石（Song *et al.*, 2020）。

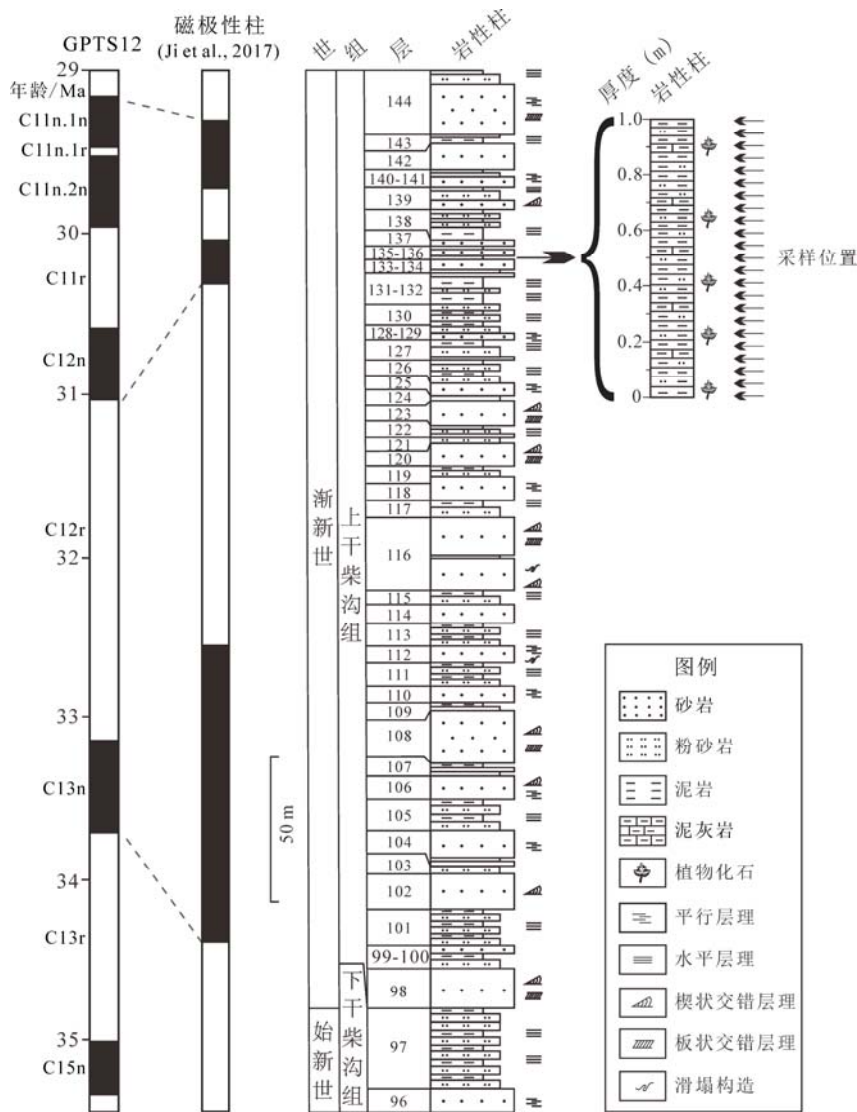


图 2 柴达木盆地大红沟剖面上干柴沟组下段年代地层框架与本次采样位置（修改自 Song *et al.*, 2020）

Fig.2 Chronostratigraphic framework of the lower member of the Shanggancaigou Formation of the Dahonggou section, Qaidam Basin, showing the location of those study samples (modified after Song *et al.*, 2020)

本项目组先前已经对大红沟剖面河湖相地层开展了高精度磁性地层学研究，将上干柴沟组地层时代限定为渐新世一中新世早期（34.2~19.5Ma）（Ji *et al.*, 2017）；基于内插法，将植物大化石的产出层位，即本次研究的泥灰岩样品采样层位的地质年代限定为~30.8Ma（图 2，Song *et al.*, 2020）。

2 样品采集和室内分析

本次研究的泥灰岩样品均采自大红沟剖面上干柴沟组 135 层产植物大化石层位（图 2）。野外采样时选择新鲜、未经风化的岩石样品，共计 23 件。

2.1 X 射线衍射（XRD）和扫描电镜（SEM）分析

选择新鲜样品，经过前处理之后使用荷兰 X’Pert PRO DY2198 型 X 射线粉晶衍射仪进行检测，设置 Ni 片滤波，X 光管、CuK α 辐射电压、电流分别为 kV 40，mA 40。扫描速率 0.02°2 θ /s，分辨率设置为 0.02°2 θ ，在 3°~65°2 θ 范围内进行扫描。此外，还选择典型样品，使用配备 EDAX 能量色散谱仪（EDS）系统的 FEI Apero S 扫描电子显微镜（SEM）进行照相分析。SEM 图像在 20kv 的加速电压、10 的光斑尺寸、约 135 μ A 的发射电流和 10mm 的工作距离下获得。XRD 分析和 SEM 分析均在中国地质大学（武汉）地质过程与成矿预测全国重点实验室进行。

2.2 碳氧同位素分析

首先在无污染环境下将新鲜样品自然风干。接着去除与空气接触的部分，取样品 0.2~0.5 g，使用玛瑙研钵研磨至小于 80 目。挑取 200~400 μ g 粉末样品置于 10 mL 的顶空反应瓶中，盖紧瓶盖，放置于 72℃恒温反应器中。由 Gasbench 自动完成加酸、CO₂ 气体制备、分离和采集步骤、最后将 CO₂ 气体输入 Gasbench II—MAT253 稳定同位素质谱仪进行碳、氧同位素测定。测样分析结果以 VPDB 为标准，标样为 GBW04416 和 GBW04417，测试结果 $\delta^{13}\text{C}$ 误差优于 $\pm 0.1\%$ ， $\delta^{18}\text{O}$ 误差优于 $\pm 0.2\%$ 。碳氧稳定同位素分析在中国地质大学（武汉）地质微生物与环境全国重点实验室完成。

3 实验结果

3.1 XRD 和 SEM 分析

如表 1 和图 3 所示，XRD 分析表明大红沟剖面上干柴沟组泥灰岩矿物组成主要为石英（36.14%~19.26%）、方解石（36.82%~7.06%）、伊利石（33.04%~15.01%）和斜长石（28.36%~9.33%），其次为绿泥石（13.45%~6.41%）和钾长石（16.48%~0%）。其中碳酸盐矿物几乎都为方解石，仅在 1 个样品（135-17）中可见微量的白云石。蒸发岩类矿物含量同样非常低，仅在 2 个样品（135-7、135-22）中发现有低含量的石盐产出。而在扫描电镜下，形态大小不一的黏土矿物，如伊利石，呈片状分布（图 4a），而半自形方解石则呈块状围绕在石英周围生长发育（图 4c）。同时 X 射线能谱分析（EDS）显示泥灰岩样品中方解石中的 Mg 含量均很低，为低镁方解石（图 4b，d-f）。

表 1 柴达木盆地大红沟剖面上干柴沟组泥灰岩矿物组成（%）XRD 分析结果

Table.1 Mineral composition (%) of the study marlstone in the Shangganchaigou Formation, Dahonggou section, Qaidam Basin								
样品 编号	矿物含量（%）							
	石英	钾长石	斜长石	方解石	白云石	绿泥石	伊利石	石盐
*135-10	21.45	2.54	18.29	25.01	—	9.43	23.27	—
135-2	28.21	2.28	28.63	10.91	—	9.10	20.88	—
135-3	24.68	—	17.00	28.91	—	8.89	20.43	—

样品 编号	矿物含量 (%)							
	石英	钾长石	斜长石	方解石	白云石	绿泥石	伊利石	石盐
135-4	36.14	—	20.12	10.61	—	10.41	22.71	—
*135-50	29.07	10.20	12.45	20.74	—	8.39	19.15	—
135-6	23.80	—	15.30	27.37	—	10.45	23.08	—
135-7	30.46	2.90	24.37	10.13	—	10.30	20.98	0.86
*135-80	25.51	—	18.16	24.01	—	8.70	23.62	—
135-9	28.88	—	25.32	7.75	—	12.62	25.43	—
135-10	29.44	3.55	16.31	21.70	—	10.18	18.82	—
135-11	23.06	2.45	20.94	7.06	—	13.45	33.04	—
*135-12	22.43	1.81	13.43	34.24	—	8.06	20.03	—
135-13	24.93	4.23	12.53	22.73	—	9.43	26.16	—
135-14	20.50	4.74	16.30	26.05	—	10.55	21.87	—
*135-15	19.26	—	17.37	26.99	—	9.18	27.20	—
135-16	31.75	2.80	16.68	12.13	—	11.01	25.63	—
*135-17	25.40	3.72	9.33	19.31	2.54	10.95	28.75	—
135-18	28.42	3.06	26.52	10.78	—	9.52	21.71	—
135-19	33.59	16.48	17.75	7.95	—	6.41	17.82	—
*135-20	30.35	—	13.85	20.46	—	12.98	22.37	—
135-21	25.55	—	19.78	24.22	—	9.38	21.07	—
135-22	27.95	—	26.68	18.84	—	7.75	16.87	1.92
*135-23	23.48	3.75	13.07	36.82	—	7.88	15.01	—

注：表中带“*”的为具代表性的典型泥灰岩样品，其 XRD 分析图见图 3。

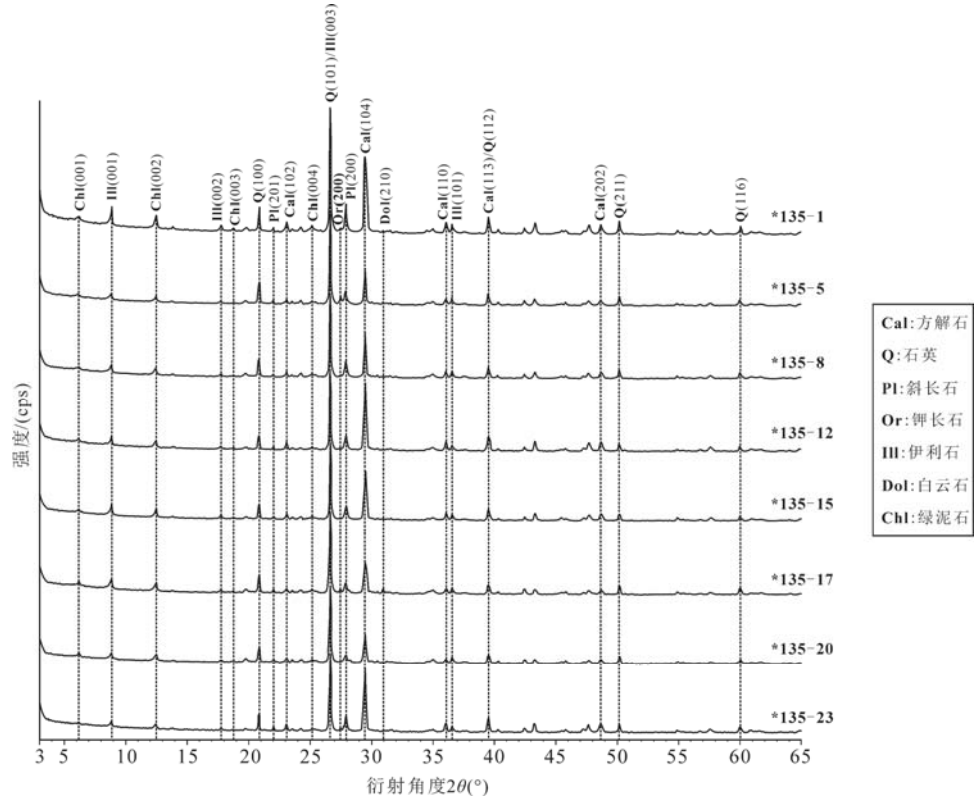


图 3 柴达木盆地大红沟剖面上干柴沟组典型泥灰岩样品 XRD 分析图

Fig.3 XRD pattern of typical marlstone samples in the Shangganchaigou Formation, Dahonggou section, Qaidam Basin

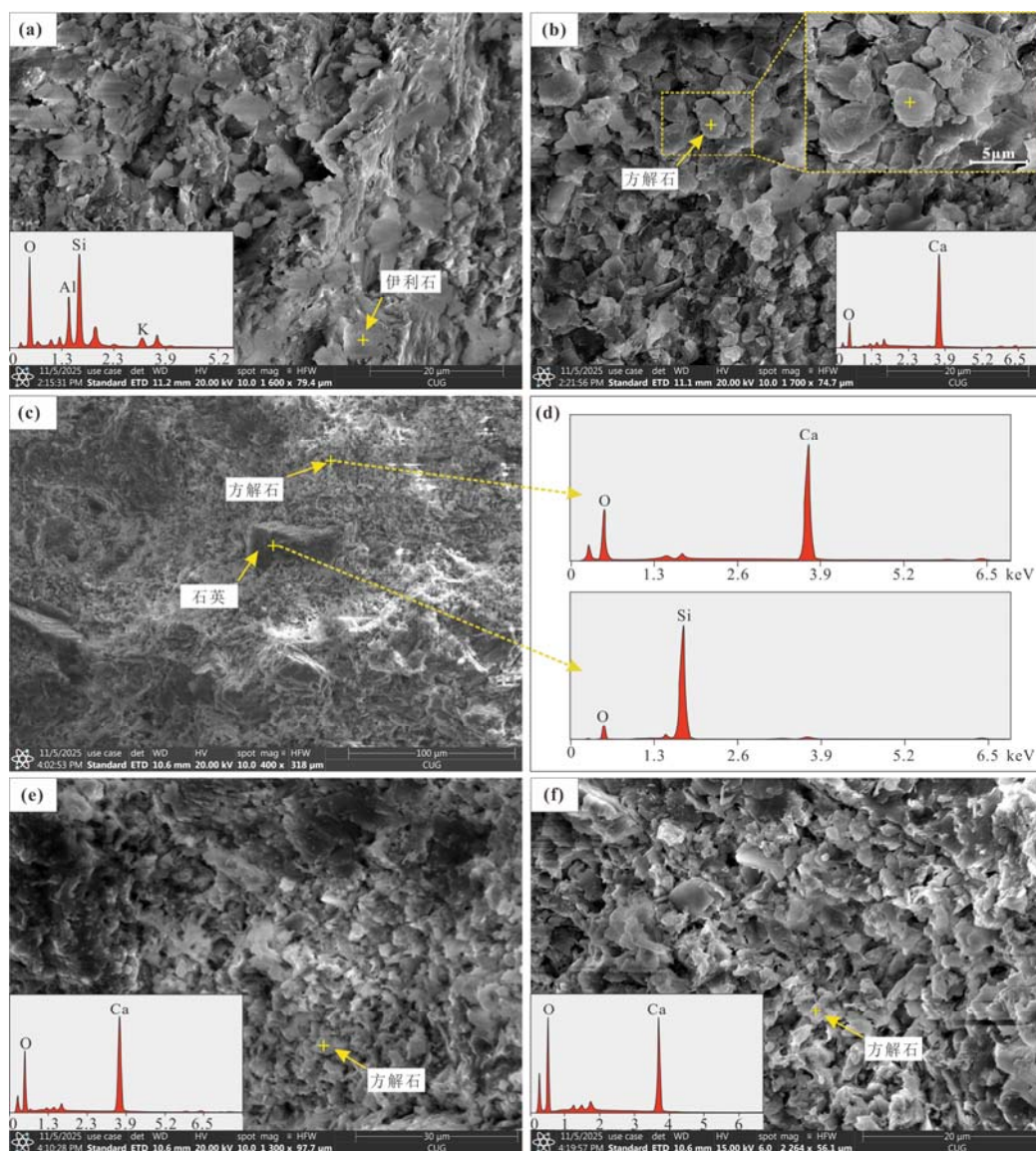


图4 柴达木盆地大红沟剖面上干柴沟组典型泥灰岩样品扫描电镜（SEM）照相和能谱（EDS）分析
 (a) 薄片状伊利石及其EDS图解，样品号135-7，扫描电镜；(b) 半自形方解石及其EDS图解，样品号135-7，扫描电镜；
 (c-d) 围绕在石英周围生长的半自形方解石及其EDS图解，样品号135-20，扫描电镜；(e-f) 半自形方解石及其EDS图
 解，样品号135-20，扫描电镜。

Fig.4 SEM (Scanning Electron Microscopy) imaging and EDS (Energy Dispersive Spectroscopy) analysis of typical marlstone samples in the Shangganchaigou Formation, Dahonggou section, Qaidam Basin
 (a) Flaky clay minerals and EDS diagram, illite, sample 135-7, SEM; (b) Semi-idiomorphic calcite and EDS diagram, calcite, sample 135-7, SEM; (c-d) Semi-idiomorphic calcite growing around quartz and EDS diagram, sample 135-20, SEM; (e-f) Semi-idiomorphic calcite and EDS diagram, sample 135-20, SEM

3.2 碳氧同位素分析

本次采自大红沟剖面上干柴沟组的泥灰岩样品的碳氧稳定同位素测定结果如表2所示， $\delta^{13}\text{C}$ 值为-3.20‰~-6.27‰，平均值为 $-4.57 \pm 0.8\%$ ； $\delta^{18}\text{O}$ 值为-7.28‰~-10.42‰，平均值为 $-8.56 \pm 0.7\%$ 。

表2 柴达木盆地大红沟剖面上干柴沟组泥灰岩 $\delta^{18}\text{O}_{\text{carb}}$ 和 $\delta^{13}\text{C}_{\text{carb}}$ 测试结果

Table.2 $\delta^{18}\text{O}_{\text{carb}}$ and $\delta^{13}\text{C}_{\text{carb}}$ values of marlstone in the Shangganchaigou Formation, Dahonggou section, Qaidam Basin

样品编号	$\delta^{13}\text{C}_{\text{VPDB}} (\text{‰})$	$\delta^{18}\text{O}_{\text{VPDB}} (\text{‰})$
135-1	-4.95	-8.34
135-2	-3.87	-9.64
135-3	-4.70	-8.32
135-4	-3.81	-8.95
135-5	-5.16	-8.74
135-6	-5.24	-8.64
135-7	-4.03	-10.42
135-8	-5.17	-8.50
135-9	-4.38	-8.97
135-10	-5.80	-8.79
135-11	-4.19	-8.71
135-12	-5.96	-9.13
135-13	-4.70	-8.14
135-14	-4.33	-7.83
135-15	-5.26	-8.36
135-16	-3.28	-7.28
135-17	-5.07	-7.97
135-18	-3.52	-7.74
135-19	-3.20	-9.06
135-20	-4.26	-7.81
135-21	-4.15	-8.26
135-22	-3.72	-7.92
135-23	-6.27	-9.40
平均值	$-4.57 \pm 0.8\text{‰}$	$-8.56 \pm 0.7\text{‰}$

4 讨论

4.1 上干柴沟组泥灰岩中碳酸盐的来源

湖泊沉积物中的碳酸盐主要包括自生碳酸盐与碎屑（外源）碳酸盐两部分：自生碳酸盐是湖泊通过无机沉淀、生物活动（包括生物诱发和生物直接成因）作用从湖泊表层沉淀形成的碳酸盐；碎屑碳酸盐则是湖泊流域内的碳酸盐岩经风化、剥蚀，再由搬运介质（如地表径流、风等）直接搬运入湖沉淀而成（汪品先和刘传联，1993；兰敏文等，2022）。在利用碳酸盐岩碳氧同位素重建地质历史时期古气候时，由于碎屑碳酸盐的形成不受湖泊环境控制，因此不具有气候指示意义，反而会干扰古气候的重建（蓝江湖等，2013）；而直接受控于湖泊环境变化的自生碳酸盐才具有气候指示意义。

本次研究的大红沟剖面上干柴沟组泥灰岩中的碳酸盐（方解石）主要为自生成因，证据如下：1）泥灰岩样品的采样层位普遍发育水平层理，沉积环境为稳定的静水环境，水动力条件弱，十分有利于自生碳酸盐（方解石）的沉淀（Leng and Marshall, 2004；兰敏文等，2022）；2）将泥灰岩样品的碳氧同位素组成与源区即柴达木盆地周缘造山带中碳酸盐岩基岩（灰岩和白云岩）的碳氧同位素组成对比（Han *et al.*, 2014；Song *et al.*, 2018），显示它们的碳氧同位素组成差异显著（图 5），表明上述泥灰岩样品中的碳酸盐（方解石）主要为自生成因，几乎没有受到源区碳酸盐岩基岩碎屑的混入影响；3）先前学者们对柴达木盆地新生代地层开展的岩性组成和碳氧同位素研究也表明柴达木盆地新生代沉积中的碳酸盐矿物主要为自生成因

(Han *et al.*, 2014; Zhang *et al.*, 2024), 且受后期成岩作用的影响微弱 (Kent-Corson *et al.*, 2009; Zhuang *et al.*, 2011b; Song *et al.*, 2018)。

综上所述, 大红沟剖面上干柴沟湖相泥灰岩样品中的碳酸盐(方解石)主要为自生成因, 且受后期成岩作用影响甚小, 其碳氧同位素组成主要受控于其发生初始沉淀时柴达木古湖湖水的碳氧同位素组成, 从而保存了早渐新世柴达木古湖湖水的原始碳氧同位素信息。

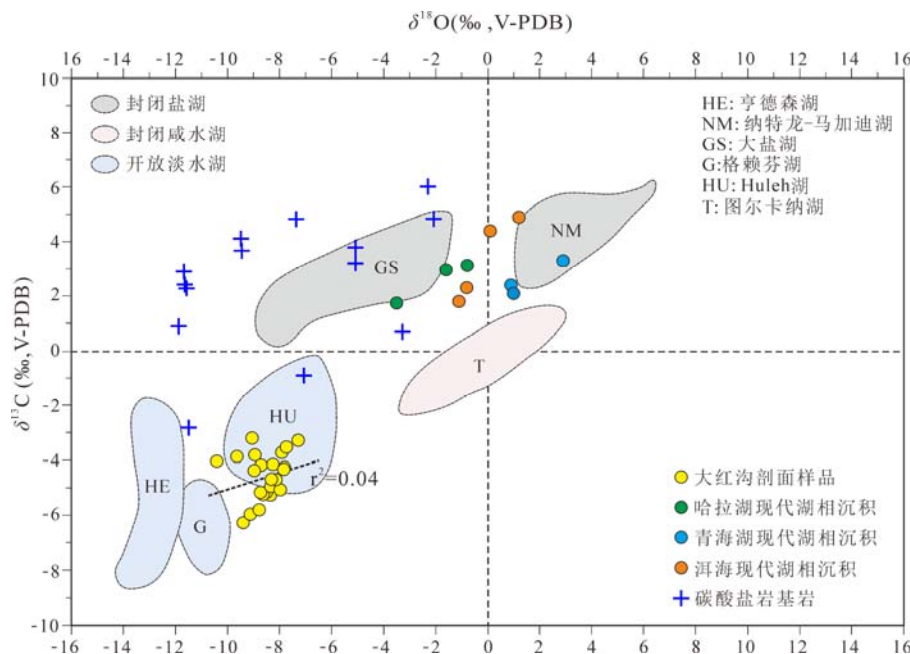


图5 大红沟剖面上干柴沟湖相泥灰岩碳氧同位素组成与区域及全球对比

Fig.5 Carbon and oxygen isotope composition of lacustrine marlstone from the Shangganchaigou Formation, Dahonggou section, and its regional and global comparison

图中碳氧同位素数据来源: 全球不同水文条件下现代湖泊沉积 (Talbot, 1990; Talbot and Kelts, 1990); 青海哈拉湖现代湖泊沉积 (马雪洋, 2017); 青海青海湖、洱海现代湖泊沉积 (Liu *et al.*, 2018); 柴达木盆地周边造山带碳酸盐岩基岩 (Han *et al.*, 2014; Song *et al.*, 2018)

4.2 碳氧同位素相关性分析及其古水文指示

在古湖泊学研究中, 通过分析湖泊沉积物中自生碳酸盐的碳氧同位素组成, 进而分析二者之间的相关性, 是判别古湖泊开放封闭程度、湖水盐度高低的一个重要手段 (Talbot, 1990; Horton *et al.*, 2016)。将本次研究的泥灰岩样品的碳氧同位素组成与区域乃至全球其他不同水文条件下的现代湖泊沉积碳氧同位素组成开展对比研究 (图5), 显示: 1) 大红沟剖面上干柴沟组泥灰岩样品的 $\delta^{13}\text{C}$ 值和 $\delta^{18}\text{O}$ 值之间相关性非常弱 ($r^2=0.04$) (图5), 表明当时柴达木古湖为一水文条件开放、水体交换活跃、水体滞留时间短的淡水湖泊 (Talbot, 1990; Leng and Marshall, 2004; Horton *et al.*, 2016), 这与上述通过 XRD 分析揭示的泥灰岩样品中蒸发岩类矿物 (白云石和石盐) 含量极低的结果一致; 2) 泥灰岩样品 $\delta^{13}\text{C}$ 值和 $\delta^{18}\text{O}$ 值的投点均落入开放的淡水湖泊区间 (图5), 与全球其他地区的典型咸水湖泊现代沉积 (Talbot, 1990) 和中国西北地区典型咸水湖现代沉积 (青海湖、洱海和哈拉湖, 马雪洋, 2017; Liu *et al.*, 2018) 的碳氧同位素组成差异显著 (图5), 同样表明早渐新世时期柴达木古湖仍然为一湿润气候背景下的开放型淡水湖泊。Keith 和 Weber (1964) 提出了利用碳酸盐岩的 $\delta^{18}\text{O}$ 和 $\delta^{13}\text{C}$ 值区分咸水碳酸盐岩和淡水碳酸盐岩的经验公式: $Z=2.048 \times (\delta^{13}\text{C}+50)+0.498 \times (\delta^{18}\text{O}+50)$, 当 Z 值大于 120 时, 为咸水碳酸盐岩; Z 值小于 120 时, 为淡水碳酸盐岩。将上干柴沟组泥灰岩样品的 $\delta^{18}\text{O}$ 和 $\delta^{13}\text{C}$ 平均值 ($\delta^{18}\text{O}=-8.56\text{‰}$, $\delta^{13}\text{C}=-4.57\text{‰}$) 分别代入该公式, 获得

的 Z 值为 113.68, 同样指示当时的湖泊水体为淡水环境。此外, 前人研究表明湖泊中自生方解石中 Mg 的含量也可作为反映湖泊水体盐度的重要指标, 低镁方解石通常产出于低盐度的开放湖泊水体中 (Müller *et al.*, 1972; Mischke *et al.*, 2008; 张菀漪等, 2011)。如上所述, 泥灰岩样品 EDS 分析结果显示其中的自生方解石主要为低镁方解石 (图 4), 进一步支持当时的柴达木古湖仍为低盐度的开放湖泊。

综上, 柴达木古湖在早渐新世 (~30.8 Ma) 仍为一湿润气候背景下的开放型淡水湖泊。这与先前基于同层位产出的植物大化石得到的柴达木盆地早渐新世仍然温暖湿润这一古气候重建结果相一致 (Song *et al.*, 2020; Han *et al.*, 2020; Hou *et al.*, 2024)。此外, 前人在柴达木盆地北部红柳沟剖面开展的植物叶片蜡质氢同位素 (δ^2H) 研究也显示早渐新世柴达木古湖表现为气候适宜背景下的湖泊扩张 (Wu *et al.*, 2021), 这与柴达木盆地西部西岔沟剖面记录的柴达木古湖湖平面在上干柴沟组早中期上升到最高点相吻合 (宋华颖等, 2010)。由于晚始新世-早渐新世中国西部的天山、阿尔金山等大型山脉尚未隆升到现今的海拔高度 (Botsyun *et al.*, 2019), 西风带仍然可以携带来自副特提斯洋的大量的水汽传输进入塔里木盆地 (Caves *et al.*, 2015; Bougeois *et al.*, 2019) 和柴达木盆地 (Song *et al.*, 2020), 从而带来充沛的降雨并促使淡水湖泊的扩张。

4.3 泥灰岩氧同位素组成的古气候意义

4.3.1 氧同位素的影响因素

湖相自生碳酸盐氧同位素组成主要受控于湖水氧同位素组成和其发生沉淀时的湖水温度 (Leng and Marshall, 2004); 湖水氧同位素组成与大气降水、入湖水体的氧同位素组成以及湖水蒸发作用的强弱直接相关 (兰敏文等, 2022); 湖水温度与湖相自生碳酸盐 $\delta^{18}O$ 值呈反比, 温度每升高 $1^\circ C$, 自生碳酸盐 $\delta^{18}O$ 值约降低 0.24‰ (Craig and Gordon, 1965)。如上文所述, 本次研究的上干柴沟组泥灰岩样品碳氧同位素组成指示早渐新世柴达木古湖仍为一开放型淡水湖泊, 在此背景下, 蒸发作用对湖水氧同位素值的影响十分微弱; 而入湖河流的同位素组分通常由流域内大气降水的同位素组分决定, 因此湖水氧同位素组成就接近流域内大气降水的氧同位素组成 (Leng and Marshall, 2004)。因此, 在此背景下, 柴达木古湖中自生碳酸盐的 $\delta^{18}O$ 值主要受控于当时大气降水的 $\delta^{18}O$ 值和湖水温度这两个因素。

此外, 由于青藏高原地区来自不同的水汽来源的降水氧同位素组成存在明显差异 (Tian *et al.*, 2007; Yao *et al.*, 2013; Li and Garzzone, 2017), 还需考虑研究区即柴达木盆地早渐新世时期与现今的水汽来源是否存在差异这一因素。而前人的研究已经显示包括柴达木盆地的亚洲内陆地区新生代早期的水汽主要来自西风带的输送 (Caves *et al.*, 2015), 这与柴达木盆地现今仍主体受控于中纬度西风环流是一致的 (Yao *et al.*, 2013; Chen *et al.*, 2019)。

4.3.2 基于氧同位素的古气候重建

先前已经基于大红沟剖面上干柴沟组地层中产出的植物叶片化石重建了柴达木盆地早渐新世 (~30.8 Ma) 时期的年均温 (MAAT) 为 $\sim 11.6^\circ C$ (Song *et al.*, 2020)。Hren 和 Sheldon (2012) 通过统计全球 88 个现代湖泊的表层水温记录, 建立了现代湖泊季节性表层水温与年均温 (MAAT) 的转换函数, 该转换函数为利用湖泊自生碳酸盐形成温度约束现代和地质历史时期陆地气候提供了方法 (Huntington *et al.*, 2015; Xiong *et al.*, 2020), 其中, 夏季 (6-8 月) 湖泊表层水温与年均温的相关性方程为:

$$MAAT = -0.0055 * T_{\text{water}}^2 + 1.476 * T_{\text{water}} - 18.915; r^2 = 0.90 \quad (MAAT < 12 \sim 15^\circ C) \quad (1)$$

据此公式 (1), 将柴达木盆地早渐新世时期的年均温 ($\sim 11.6^\circ C$) 代入公式 (1), 从而获得了柴达木古湖早渐新世时期的夏季表层水温为 $\sim 22.57^\circ C$ 。考虑到先前学者的研究已经表明现代湖泊中自生碳酸盐主要

在夏季的表层湖水中沉淀形成 (Hren *et al.*, 2013; Huntington *et al.*, 2015; Horton *et al.*, 2016), 据此基于泥灰岩样品氧同位素组成和柴达木盆地早渐新世夏季表层湖水温度, 可以重建当时柴达木古湖湖水的氧同位素组成。

采用 Kim and O' Neil (1997) 提出的分馏系数计算公式:

$$1000\ln\alpha_{(\text{calcite-water})}=18.03(10^3T^{-1})-32.42, \text{ 其中 } T \text{ 为绝对温度 } (T_{\text{绝对温度}}=T_{\text{摄氏度}}+273.15) \quad (2)$$

计算得出~22.57°C条件下方解石和水的分馏系数为 1.02896。

根据分馏系数定义公式:

$$\alpha_{(\text{calcite-water})}=(1000+\delta^{18}\text{O}_{\text{calcite}})/(1000+\delta^{18}\text{O}_{\text{water}}) \quad (3)$$

将上述温度条件下对应的分馏系数和泥灰岩样品中 $\delta^{18}\text{O}_{(\text{VPDB})}$ 平均值 ($-8.56\pm0.7\text{‰}$) 分别代入公式 (3), 计算得到~22.57°C条件下柴达木古湖湖水 $\delta^{18}\text{O}_{(\text{water-VPDB})}$ 值为 $-36.47\pm0.7\text{‰}$ 。

最后, 利用 SMOW 和 VPDB 标准之间的转换公式 $\delta^{18}\text{O}_{(\text{SMOW})}=1.03091*\delta^{18}\text{O}_{(\text{VPDB})}+30.91$ (4) 得到柴达木古湖早渐新世时期的湖水 $\delta^{18}\text{O}_{(\text{water-SMOW})}$ 值为 $-6.67\pm0.7\text{‰}$ 。

考虑到自早渐新世以来全球海水 $\delta^{18}\text{O}$ 值以及柴达木盆地纬度的变化, 还需对该数据结果进行校正。由于全球降温和大陆冰盖的形成, 现今地球海洋中海水的氧同位素值相比早渐新世增加了~0.5‰ (Zachos *et al.*, 1994); 前人的研究显示柴达木盆地自早渐新世以来并没有发生显著的纬向移动 (Halim *et al.*, 2003; Yu *et al.*, 2014), 这与先前基于 HadCM3L 模型模拟研究得出的柴达木盆地早渐新世古纬度数据 (37.47°N) 一致 (Song *et al.*, 2020), 因而不需要进行纬度校正。经过综合校正之后最终获得了早渐新世柴达木盆地夏季大气降水的氧同位素为~ $-6.2\pm0.7\text{‰}$ 。将最终获得的柴达木盆地早渐新世夏季大气降水 $\delta^{18}\text{O}$ 值与该盆地现代夏季大气降水和河水 $\delta^{18}\text{O}$ 值 $-8.4\pm2.2\text{‰}$ (朱建佳等, 2015; 朱谱成, 2015; Song *et al.*, 2018) 进行对比 (图 6), 显示柴达木盆地早渐新世夏季大气降水 $\delta^{18}\text{O}_{(\text{water-SMOW})}$ 值比该盆地现代夏季大气降水和河水 $\delta^{18}\text{O}_{(\text{water-SMOW})}$ 值高~2.2‰ (图 6)。

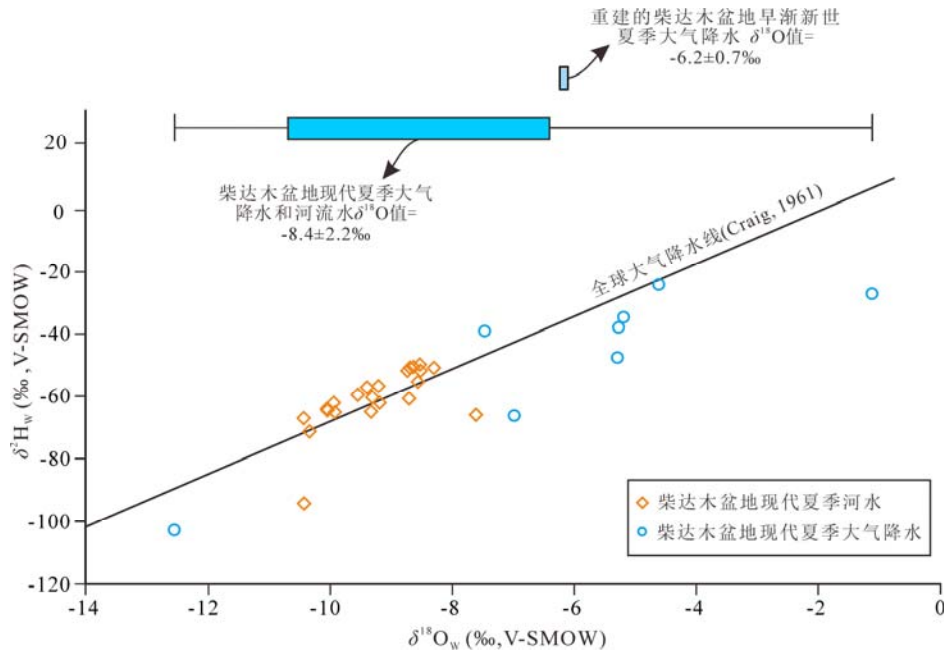


图 6 柴达木盆地现代夏季大气降水和河水氢氧同位素组成及与重建的早渐新世夏季大气降水氧同位素值对比

Fig.6 Modern summer atmospheric precipitation and river water hydrogen and oxygen isotope composition in the Qaidam Basin and their comparison with the reconstructed Early Oligocene summer atmospheric precipitation oxygen isotope values

图中柴达木盆地现代夏季大气降水和河水 $\delta^{18}\text{O}$ 、 $\delta^2\text{H}$ 数据来自：朱建佳等（2015）、朱谱成（2015）和 Song *et al.*（2018）；全球大气降水线 $\delta^{18}\text{O}$ 、 $\delta^2\text{H}$ 数据来自：Craig（1961）。

除温度引起的方解石与水之间氧同位素组成约-0.24‰/°C的分馏效应外（Craig and Gordon, 1965），大气降水的年平均 $\delta^{18}\text{O}$ 值与地表气温之间还呈显著正相关（~-0.58‰/°C）（Rozanski *et al.*, 1993）。因此，温度变化对湖相自生碳酸盐 $\delta^{18}\text{O}$ 值的综合影响为~-0.34‰/°C。因此，柴达木盆地现代夏季大气降水和河水 $\delta^{18}\text{O}$ 值发生的~-2.2‰的负偏指示自早渐新世至今柴达木盆地夏季发生了显著降温，降温幅度达~6.5°C，这与基于大红沟剖面上干柴沟组产出的植物叶片化石重建的柴达木盆地自早渐新世（~30.8 Ma）至今的降温幅度（~7.5°C）基本一致（Song *et al.*, 2020）。基于深海底栖有孔虫壳体 Mg/Ca 比和氧同位素重建的全球海洋底层海水温度变化记录显示，自早渐新世（30.8 Ma）以来全球底层海水温度降低了~6.2°C（Lear *et al.*, 2000; Hansen *et al.*, 2008）。而同样基于深海底栖有孔虫壳体氧同位素重建的全球平均温度记录显示，自早渐新世（30.8 Ma）以来全球平均温度降低了~7.5°C（Westerhold *et al.*, 2020）。此外，除了海洋温度变化记录，陆相古土壤结核团簇同位素古温度 T（ Δ_{47} ）记录显示，与柴达木盆地处于相近纬度的美国西部地区自渐新世以来的降温幅度同样达 6-8°C（Fan *et al.*, 2014）。由此可知，本次得出的柴达木盆地早渐新世以来夏季的显著降温幅度与地球自早渐新世以来的降温幅度基本保持一致。而前人开展的长时间尺度化学风化记录和湖泊古水文变迁研究同样显示柴达木盆地自早渐新世以来的气候演变主要受控于全球变化（Song *et al.*, 2013; Guo *et al.*, 2018; Fang *et al.*, 2019; Ye *et al.*, 2020）。与此同时，考虑到青藏高原东北部新近纪以来经历了多期构造隆升事件（Li *et al.*, 2014; Miao *et al.*, 2022），叠加在全球变冷的气候背景之上，青藏高原隆升对柴达木盆地自早渐新世以来的显著降温同样有重要影响。综上，全球变冷和青藏高原隆升共同驱动了自早渐新世以来柴达木盆地夏季的显著降温，其中全球变冷起到了主导作用。

5 结论

（1）柴达木盆地大红沟剖面上干柴沟组湖相泥灰岩中的碳酸盐（方解石）主要为自生成因，保存了早渐新世柴达木古湖湖水的原始碳氧同位素信息。

（2）通过泥灰岩碳氧同位素组成相关性分析以及与区域及全球现代湖泊沉积自生碳酸盐碳氧同位素组成开展对比研究，揭示出柴达木古湖在早渐新世（~30.8 Ma）仍然为一湿润气候背景下的开放型淡水湖泊。

（3）基于泥灰岩氧同位素组成重建了柴达木盆地早渐新世（~30.8 Ma）夏季大气降水的氧同位素值。该值比柴达木盆地现代夏季大气降水和河水 $\delta^{18}\text{O}$ 值高~2.2‰，指示自早渐新世至今柴达木盆地夏季发生了显著降温，幅度达~6.5°C。全球变冷和青藏高原隆升共同驱动了早渐新世以来柴达木盆地夏季的显著降温，其中全球变冷起到了主导作用。

此外，需要强调的是，尽管本文依旧采用“传统年龄模型”，而目前学界对柴达木盆地新生代年代地层格架依然存在较大争议，希冀今后学界开展更深入系统的研究以取得共识。

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