

单斜辉石矿物结构及化学特征在岩浆储运系统重建中的应用

夏玉洪, 赖绍聪*, 朱韧之*

大陆演化与早期生命全国重点实验室, 西北大学地质学系, 西安 710069

摘要: 破译岩浆储运系统对于揭示幔源岩浆岩成因、演化及相关资源环境效应具有重要的意义。单斜辉石是岩浆岩中的常见造岩矿物, 对其结晶时的物理化学条件极其敏感, 是重建岩浆储运系统的重要载体。本文综述了单斜辉石所能揭露的主要岩石学信息, 包括: 矿物晶体成因、结构特征记录的岩浆过程、结晶物理化学条件 (P - T - H_2O - fO_2)、元素扩散记录的时间尺度、微量元素反映的源区信息。综合利用这些信息, 可以重建幔源岩浆从源区到浅表所经历的岩浆储运系统, 并识别其中的岩浆演化过程, 为幔源岩浆演化提供矿物学尺度的证据。未来研究应致力于提升单斜辉石温压计、湿度计及氧逸度计的精度与适用范围, 并拓展其原位同位素分析体系; 同时, 将研究对象扩展至不同构造背景及不同种类岩石及地外样品, 为深入理解地球乃至行星尺度的岩浆演化过程提供关键约束。

关键词: 单斜辉石; 幔源岩浆岩; 岩浆储运系统; 岩浆演化

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Textural and Chemical characteristics of Clinopyroxene for the Reconstruction of Magma Plumbing System

Xia Yuhong, Lai Shaocong*, Zhu Renzhi*

State Key Laboratory of Continental Evolution and Early Life, Department of Geology, Northwest University, Xi'an 710069, China

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作者简介: 夏玉洪 (1998-), 男, 博士研究生, 从事岩浆岩岩石学研究。ORCID: 0009-0003-5743-8839. E-mail: xiayhgeo@163.com

***通讯作者:** 赖绍聪 (1963-), 男, 教授, 从事岩浆岩岩石学研究。ORCID: 000-0003-0629-2748. E-mail: shaocong@nwu.edu.cn 朱韧之 (1989-), 男, 教授, 从事岩浆岩岩石学研究。ORCID: 0000-0001-5433-9585. E-mail: rzzhunwu@163.com

Abstract: Deciphering magma plumbing systems is of great significance for understanding the petrogenesis and evolution of mantle-derived igneous rocks, as well as the associated resource and environmental implications. Clinopyroxene is a common rock-forming mineral in igneous rocks and is extremely sensitive to the physicochemical conditions of its crystallization, making it a key recorder for reconstructing magma plumbing systems. This paper reviews the key petrological information revealed by clinopyroxene, including: crystal origins, magma processes recorded by textural characteristics, physicochemical conditions (P-T-H₂O-fO₂), timescales recorded by elemental diffusion, and source information reflected by trace elements. By integrating this information, it is possible to reconstruct the magma plumbing system through which mantle-derived magmas travel from source to shallow levels, and to identify the magma evolution processes therein, providing mineral-scale evidence for the evolution of mantle-derived magmas. Future research should focus on improving the accuracy and applicability of clinopyroxene thermobarometers, hygrometers, and oxybarometers, coupled with expanding its in-situ isotopic analytical systems. Extending such investigations to diverse tectonic settings, rock types, and extraterrestrial samples will provide key constraints on magmatic evolution on both earth and planetary scales.

Key words: Clinopyroxene; Mantle-derived igneous rocks; Magma plumbing system; Magma evolution

0 引言

岩浆储运系统 (magma plumbing system) 指从岩浆源区延伸至地壳浅表的垂直岩浆演化系统, 由相互连通的多层级运移通道和岩浆储库构成, 主导熔体上升、储存、分异等关键过程 (Burchardt, 2018; Magee et al., 2018) (图 1)。近些年来的研究表明, 岩浆储运系统中通常会发生多批次的岩浆补给 (Ubide and Kamber, 2018; Zhang et al., 2024a; Xia et al., 2026)、岩浆混合 (Xing and Wang, 2020; Yang

et al., 2022; Shi et al., 2026)、晶体-熔体反应 (Di Stefano et al., 2020; Zhang et al., 2024a; Cheng et al., 2024)、地壳混染 (Streck, 2008; Liu et al., 2026) 等复杂的岩浆演化过程 (图 1)。这些过程被认为是理解岩浆房物理化学状态、熔体变化的关键, 为解释幔源岩石的成因及岩浆动力学背景提供了重要的约束 (Streck, 2008; Cashman et al., 2017; Sparks et al., 2019; Pan et al., 2022; Yang et al., 2022; Ma et al., 2024; Zhang et al., 2024a; Xia et al., 2026)。此外, 精细刻画岩浆储运系统结构与熔体演化过程, 对深入理解关键成矿元素富集及火山喷发机制具有关键意义 (Caricchi and Blundy, 2015; Ubide and Kamber, 2018; Ubide et al., 2023; Zhu et al., 2023; Day et al., 2024; Simpson et al., 2025), 直接关系到资源勘查与火山预警等重大课题, 与人类生存发展息息相关。例如, 发生在岩浆储库中的岩浆补给及反应渗流 (reactive porous flow) 注入促进了 REE、Nb、Zr 等不相容元素的富集, 在塔里木大火成岩省瓦吉里塔格正长岩的稀土和稀有金属富集及矿化过程中扮演了重要角色 (Zhu et al., 2023)。1974 年到 2014 年期间埃特纳火山喷发物中的单斜辉石所记录的 Cr 环带揭示了火山喷发前岩浆储库中最后一批基性岩浆补给事件, 这一补给事件被认为是触发该火山喷发的关键 (Ubide and Kamber, 2018)。

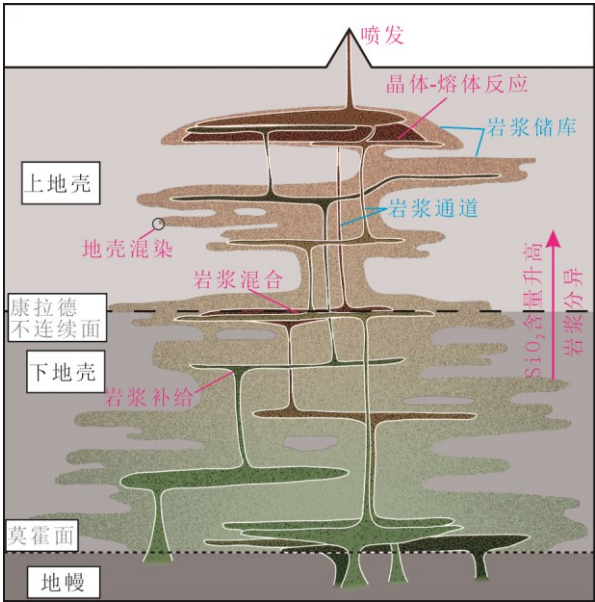


图 1 岩浆储运系统示意图 (修改自 Magee et al., 2018)

Fig. 1 Schematic diagram of the magma plumbing system (modified from Magee et al., 2018)

然而, 幔源岩浆演化过程中的岩浆储运系统却常常受到忽视, 这主要源于两个方面的原因: (1) 一些幔源火山岩中大量地幔捕虏体的出现被认为是岩浆快

速上升的证据，这使得岩浆上升路径中的演化过程易被忽视（Liu et al., 2021; Simpson et al., 2025）；（2）岩石学研究中最常用的全岩地球化学分析得到的是整块岩石被粉碎、均一化后的整体信息，在揭示精细的岩浆过程方面却有很大的局限性（谢元惠等, 2021; Pan et al., 2022; Zhang et al., 2024a）。岩石中的矿物生长往往受控于结晶环境的变化，矿物尺度上的信息（矿物间/单颗矿物内部的成分不均一性、矿物结构特征等）能够很好地记录岩浆演化过程（Davidson and Tepley, 1997; Davidson et al., 2007; Streck, 2008, Ubide et al., 2015, 2019; Pan et al., 2022; Yang et al., 2022; Zhang et al., 2024a; Zhu et al., 2024; Lu et al., 2025; Xia et al., 2026）。因此，从单矿物入手进行研究，能有效规避全岩地球化学手段在破译复杂岩浆过程中的局限性。

单斜辉石是一种广泛出现在从酸性岩到超基性岩中的链状硅氧骨干硅酸盐矿物（图 2），其分子化学通式可表示为 $XY[T_2O_6]$ 。其中，X 在晶体结构中占据面积较大的 M_2 位置，主要为 Na^+ 、 Ca^{2+} 、 Fe^{2+} 、 Mg^{2+} 、 Li^+ 等六到八次配位的大离子半径阳离子；Y 占据 M_1 位置，主要为 Mn^{2+} 、 Fe^{2+} 、 Mg^{2+} 、 Fe^{3+} 、 Cr^{3+} 、 Al^{3+} 、 Ti^{4+} 及三价稀土元素离子（ REE^{3+} ）等六次配位高价态小离子半径阳离子；T 位置则主要为硅氧骨干中的四次配位阳离子，以 Si^{4+} 为主，并伴随少量 Al^{3+} 和 Ti^{4+} 的替代（Cameron and Papile, 1981; Yang et al., 2009; 赵珊茸, 2017; 胡君豪, 2018）。这种特殊的晶体化学结构使单斜辉石相对于斜长石、斜方辉石、橄榄石等其他常见造岩矿物能够容纳更为广泛的元素（如稀土元素、高场强元素、大离子亲石元素及过渡金属元素等）。更重要的是，岩浆运移过程中的演化过程会影响这些不同阳离子的分配行为。因此，单斜辉石是重建岩浆储运系统演化过程的良好研究对象（Müller et al., 2013; Mollo et al., 2018; Zhou et al., 2021; Geng et al., 2022; Neave et al., 2024）。岩浆演化过程不仅反映在成分变化上，还反映在其结构特征，如振荡环带、筛状结构、扇形环带等结构都记录了岩浆储运系统内物理化学条件的改变（详细见第 2 节）。

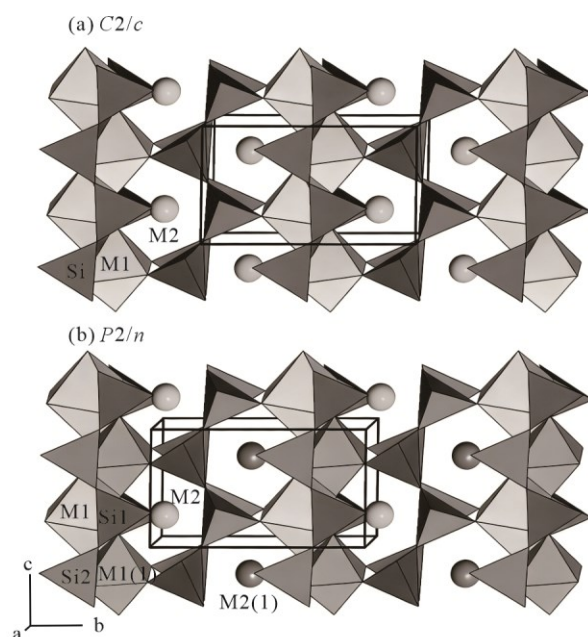


图 2 $C2/c$ (a) 和 $P2/n$ (b) 单斜辉石的晶体结构示意图 (据 Yang et al., 2009)

Fig. 2 The crystal structure of $C2/c$ (a) and $P2/n$ (b) clinopyroxenes (after Yang et al., 2009)

单斜辉石的主量元素成分对温度、压力、熔体水含量等条件极为敏感，因此其能为恢复岩浆演化过程中的结晶环境提供有效约束（详细见第 3 节）。单斜辉石单颗粒内部元素扩散可约束岩浆储库内的时间尺度（详细见第 4 节）。单斜辉石微量元素则能够揭示源区交代介质、源区岩性等岩浆源区信息（详细见第 5 节）。此外，在富含挥发份的碱性岩浆体系中，由于富水体系抑制斜长石的结晶，往往出现单斜辉石而非斜长石（Danyushevsky, 2001）。基于上述优势，单斜辉石在构建幔源岩浆的岩浆储运系统及揭示岩浆演化过程等方面展现出显著潜力，近年来受到了众多研究者的关注。据不完全统计，近十多年来，关于单斜辉石揭示岩浆储运系统的文章发表数量逐步增多（图 3），这些成果极大地丰度和完善了岩浆储运系统理论。本文系统综述了单斜辉石所能记录的岩石学信息，归纳了其在重建岩浆储运系统中的作用，以期国内外岩浆岩相关研究者提供参考。

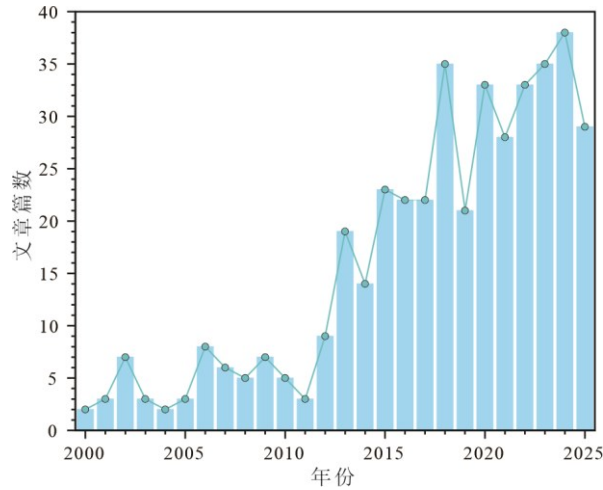


图 3 自 2000 年以来发表的关于单斜辉石揭示岩浆储运系统的文献统计（自 2011 年起，该领域发表文献数量呈逐年上升趋势。数据来源于 Web of Science，搜索关键词为“clinopyroxene”和“magma plumbing system”，数据统计截至 2025 年 12 月 31 日）

Fig. 3 Paper published on “clinopyroxene reveals magma plumbing system” since 2000 (Since 2011, the annual number of publications in this field has shown a increasing trend. Data were sourced from the Web of Science using the search keywords "clinopyroxene" and "magma plumbing system," with statistics collected up to December 31, 2025).

1 火山岩中常见的单斜辉石斑晶类型及其意义

在传统的岩浆岩结构分类方案中，将岩石中具有矿物颗粒大小明显不同的两群的结构称为“斑状结构”（桑隆康等, 2012）。这种命名方案仅根据矿物颗粒大小将颗粒较大的矿物颗粒称为“斑晶”，却并未从成因上去区分这些矿物晶体。近年来，研究者们发现这些“斑晶”矿物可能包含了“自生晶（autocryst）”、“再循环晶（antecryst）”、“转熔晶（peritectic crystal）”、“捕虏晶（xenocryst）”等多种成因的矿物晶体（Cashman and Blundy, 2013; 罗照华等, 2013; 马昌前等, 2020; Di Stefano et al., 2020; Petrone et al., 2022）。这一认识对传统意义上的“斑晶-基质”分类提出了挑战。

在近年来的研究中，研究者们发现火山岩中颗粒较大的单斜辉石同样是常常作为“晶体群”而出现（Xing and Wang, 2020; Petrone et al., 2022; Zhang et al., 2023, 2024a; Xia et al., 2026）。即便在同一期火山喷发形成的岩石中，单斜辉石也常

具有多种来源。其成因差异可通过化学成分的变化得以有效识别。单斜辉石自生晶是严格意义上的“斑晶”，即由寄主岩浆结晶而成，其成分与周围熔体完全平衡，表现在单斜辉石与熔体的 Fe-Mg 交换系数处于平衡范围内（镁铁质岩浆体系中与熔体平衡的单斜辉石 $K_D(\text{Fe-Mg})^{\text{单斜辉石-熔体}}=0.28\pm0.08$, Putirka, 2008）。单斜辉石再循环晶指同源岩浆较早结晶后被晚期岩浆捕获的单斜辉石。其表现出与基质成分不平衡并且保存寄主岩浆早期演化阶段相对高 Mg 的特征。在岩浆演化过程中，再循环晶单斜辉石可能会与寄主岩浆发生相互作用形成低 Mg 的边部。如 Zhang et al. (2024a) 报道的西秦岭马圈沟新生代碱性玄武岩中具有高 Mg 核部及低 Mg 边部的单斜辉石即为再循环晶单斜辉石与寄主岩浆发生反应所形成。单斜辉石捕虏晶为寄主岩浆上升过程中所携带的非同源岩浆系统结晶矿物和地幔/地壳岩石解体产物中的单斜辉石。幔源单斜辉石捕虏晶会保留幔源岩石的矿物化学特征，如西秦岭礼县-宕昌地区好梯新生代黄长岩中的幔源单斜辉石捕虏晶具有高达 90 以上的 Mg# 值，与区域上报道的地幔捕虏体中单斜辉石的 Mg# 值吻合 (Xia et al., 2026)。下地壳的单斜辉石捕虏晶的主量元素成分上以异常高的铁含量显著区别于其他类型单斜辉石晶体，富铁的下地壳单斜辉石捕虏晶在单偏光镜下呈绿色调，当其被寄主富 Mg 岩浆捕获后，晶体继续发生生长形成绿色核部、浅色边部的单斜辉石，即“绿核单斜辉石” (Szabó and Bodnar, 1998; Geng et al., 2022)。微量元素上，下地壳单斜辉石捕虏晶以与长石平衡的 Sr、Eu 负异常为特征 (Geng et al., 2022; Xia et al., 2026)。同一个火山岩系统中出现上述多样单斜辉石晶体类型表明在岩浆运输过程中具有多源的物质供给及持续的演化过程。相较于捕虏体的可挑选性，“再循环晶”、“捕虏晶”等多种来源的矿物晶体则是全岩地球化学研究中几乎不可避免的干扰源，这些晶体可能会对全岩成分进行改造，比如幔源单斜辉石/橄榄石捕虏晶的加入可能会使全岩组分的 Mg# 升高，从而在一定程度上限制了对岩浆演化过程的理解 (谢元惠等, 2021; Zhang et al., 2024a, 2025)。因此，识别这些不同类型晶体对于精确构建完善的岩浆储运系统至关重要。

2 单斜辉石的结构特征：破译岩浆储库成分变化及动力学过程

火山岩中单斜辉石常表现出多样的结构特征，如振荡环带、筛状结构、骸晶结构、扇形分区等（图 4）。这些多样的单斜辉石结构记录了复杂的生长与改造过程，反映了岩浆储运系统中物理化学条件的改变（Streck, 2008）。例如，笔者在西秦岭好梯新生代黄长岩中识别出具有筛状结构、反应结构的单斜辉石，指示其岩浆储运系统中经历了持续的熔体活动（Xia et al., 2026）。本节总结了单斜辉石中常见的结构及其成因。

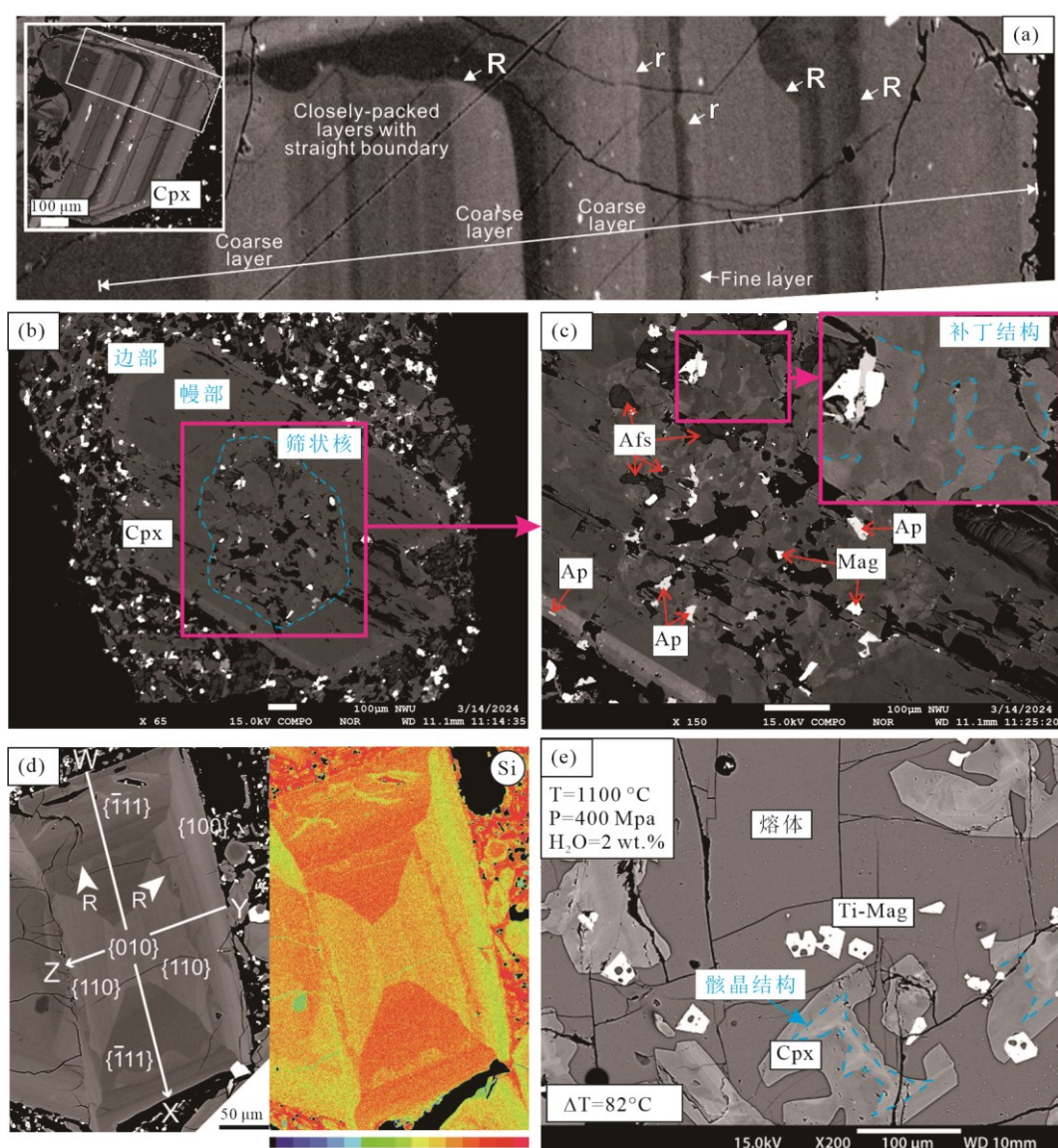


图 4 天然和试验样品中具有不同结构特征的单斜辉石：（a）北天山超钾质煌斑岩脉中具有振荡环带的单斜辉石，不同韵律层之间具有溶蚀界面（据 Xing and Wang, 2020）；（b-c）

西秦岭好梯新生代黄长岩中具筛状结构的单斜辉石，筛状区域具明暗相间的补丁结构及碱性长石、磷灰石、磁铁矿等矿物包裹体（据 Xia et al., 2026）；（d）意大利 Mt. Etna 火山岩中具扇形环带的单斜辉石（据 Ubide et al., 2019）；（e）高程度过冷实验条件下具骸晶结构的单斜辉石（据 Masotta et al., 2020）

Cpx: 单斜辉石; Afs: 碱性长石; Ap: 磷灰石; Mag: 磁铁矿; Ti-Mag: 钛磁铁矿; R/r: 熔蚀界面

Fig. 4 Clinopyroxene with varying textural features in natural and experimental samples: (a) Clinopyroxene with oscillatory zoning in ultrapotassic lamprophyre dyke from the North Tianshan. Resorption surface occurs between different layers (after Xing and Wang, 2020). (b–c) Clinopyroxene with sieve texture in the Cenozoic Haoti melilitites from the West Qinling. The sieved domains exhibit bright and dark patchy structures and contain mineral inclusions such as alkali-feldspar, apatite, and magnetite (after Xia et al., 2026). (c) Clinopyroxene with sector zoning in volcanic rocks from Mt. Etna, Italy (after Ubide et al., 2019). (d) Clinopyroxene with skeletal texture formed under highly undercooling experimental conditions (after Masotta et al., 2020).

Cpx: clinopyroxene; Afs: alkali-feldspar; Ap: apatite; Mag: magnetite; Ti-Mag: Titano-magnetite; R/r: Resorption surface

2.1 振荡环带

振荡环带 (oscillatory zoning) 是斜长石、单斜辉石等矿物中常见的一种结构，表现为背散射照片中明暗程度的周期性变化。Ginibre et al. (2002a, 2002b) 发现斜长石的振荡环带不同生长层间存在不规则的溶蚀界面 (resorption surface)，并表现出剖面上的 An 值、Sr 同位素组成等参数的变化 (Davidson and Tepley, 1997)。这些变化反映了矿物结晶过程中温度、压力、熔体成分、岩浆水含量等关键因素的改变，能够反映岩浆演化的动力学过程及岩浆补给事件 (Ginibre et al., 2002a, 2002b)。当熔蚀界面及成分变化表现较弱，被认为是岩浆房内局部对流的结果，反映了岩浆房内局部岩浆成分变化；当这些现象表现强烈，则反映了外来岩浆的补给及混合作用 (Ginibre et al., 2002a, 2002b)。单斜辉石振荡环带中具有类似的

熔蚀界面 (Xing and Wang, 2020; 谢元惠等, 2021; Zhang et al., 2023)。在化学成分上, 单斜辉石振荡环带表现出不均一性。如云南大理新生代钾质煌斑岩 (Zhang et al., 2023) 与北天山超钾质煌斑岩 (Xing and Wang, 2020) 中振荡环带单斜辉石都表现出 Mg#值、Cr 含量的波动, 证明在其岩浆演化过程中存在多批次不同演化程度岩浆的补给作用。值得注意的是, Cr 在单斜辉石中属相容元素 ($D_{Cr}^{Cpx/melt} = 3-17$, D'Orazio et al., 1998), 对岩浆成分变化极其敏感, 且相较于 Fe、Mg 等元素不会发生明显扩散 (Ubide and Kamber, 2018)。在单斜辉石元素面扫图中, Cr 表现出比其他元素更加明显的生长环带 (Petrone et al., 2022)。因此, Cr 在揭示岩浆补给事件的应用中有着很好的表现 (Petrone et al., 2022; Xia et al., 2026)。

2.2 筛状结构

筛状结构 (sieve texture), 又称海绵结构 (spongy texture), 指矿物内部发育的密集孔状微观结构, 形似筛子。关于单斜辉石筛状结构的成因有以下几种模式: (1) 高程度过冷; 在高程度过冷的条件下, 晶体的化学及结构平衡被打破进而发生反应, 形成以骸晶状或树枝状结构为特征的筛状单斜辉石 (Masotta et al., 2020; Zhou et al., 2021)。(2) 减压部分熔融; 当结晶自深部的单斜辉石或地幔捕虏体中单斜辉石随岩浆上升至更浅部位, 随着压力降低, 单斜辉石可能发生溶解形成筛状结构, 这一过程伴随着 Na_2O 、 Al_2O_3 、 Al^{VI} 含量的降低 (Su et al., 2011; Pan et al., 2018; Neave and MacLennan, 2020)。(3) 晶体-熔体反应; 当岩浆体系中有新的熔体加入时, 熔体与晶体之间会发生反应以达到新的平衡 (Di Stefano et al., 2020; Zhang et al., 2024a; Cheng et al., 2024)。与温度、压力诱导的反应不同, 由晶体-熔体反应诱导形成的筛状结构伴随着新的化学组分的加入, 因此在筛状区域常可见一些微小的矿物包裹体。例如, Zhang et al. (2024a) 报道了西秦岭马圈沟新生代碱性玄武岩中的筛状结构单斜辉石具有钾长石、钛铁矿等矿物包裹体。类似的现象在西秦岭好梯新生代黄长岩中的单斜辉石中也有报道 (Xia et al., 2026)。单斜辉石中的钾含量通常较低, 不足以形成钾长石, 富钾熔体的加入则提供了形成钾长石包裹体的钾来源。与振荡环带类似, 筛状结构的形成同样受控于温度、压力、熔体成分等因素, 能有效揭示岩浆运移过程中热力学

演化、熔体迁移等动态变化。不同的是，振荡环带反映的是矿物晶体结晶过程中物理化学条件的变化；而筛状结构则是矿物晶体形成后，周围环境发生改变的响应。

2.3 骸晶结构及扇形环带

不同于普遍的由中心向边缘逐层生长的模式，近些年来基于实验及自然样品的研究发现单斜辉石晶体存在骸晶状(skeletal texture)、树枝状(dendritic texture)、扇形环带(sector zoning)等不规则结构特征，表明这些单斜辉石的生长经历了更为复杂的过程(Ubide et al., 2019; Masotta et al., 2020; Zhou et al., 2021)。火成岩中的矿物生长受控于近平衡的界面控制(interface-controlled)和不平衡的扩散控制(diffusion-controlled)两种动力学机制(Kirkpatrick, 1981; Hammer, 2008)。晶体生长具体受控于何种机制与过冷度和冷却速率有关(Mollo et al., 2013a; Ubide et al., 2019; Masotta et al., 2020; Zhou et al., 2021)。当过冷度低时，界面控制占主导，熔体中元素扩散速率快于晶体生长速率，此时晶体多呈自形且成分均一(Hammer, 2008 及其引用文献; Mollo et al., 2013a)。在过冷度高的扩散控制下，晶体生长速率超过熔体中元素扩散速率，使得晶体不平衡生长，形成结构不成熟的晶体(Mollo et al., 2013a 及其引用文献)。当过冷度逐渐增高且冷却速率很快时，体系中趋于形成具骸晶结构的单斜辉石，这是由于该条件下熔体中的化学扩散受到限制不足以支持形成的骸晶体继续生长，从而晶体成核效应超过生长效应，使得早期形成的骸晶得以保留(Mollo and Hammer, 2017 及其引用文献; Pontesilli et al., 2019; Masotta et al., 2020)。若体系中温度保持在晶体的目标静息温度(target resting temperature)，单斜辉石晶体的生长速率又会反超成核速率，沿着骸晶继续生长(Masotta et al., 2020)。Masotta et al. (2020)以 Mt. Maletto 火山喷发的粗面玄武岩为原始组分，开展了单斜辉石结晶的热力学实验研究。结果表明，随着过冷度增加，所结晶的单斜辉石具有颗粒由大到小、晶形由自形向半自形—扇形环带—他形(骨骼状、树枝状)的转变。扇形环带单斜辉石可分为漏斗区(hourglass)和棱柱区(prism)，漏斗区($\{-111\}$)平行晶体c轴生长，富Si、Mg而贫Ti、Al，在背散射图中颜色更暗；棱柱区($\{100\}$ 、 $\{110\}$ 、

{010}) 垂直晶体 c 轴生长, 富 Ti、Al 而贫 Si、Mg, 在背散射图中颜色更亮 (Ubide et al., 2019)。这种耦合的元素变化归因于电价平衡: 随着 Al^{3+} 对单斜辉石 T 位置上 Si^{4+} 替代的增加, 需要 Ti^{4+} 替代 M1 位置上的 Mg^{2+} 和 Fe^{2+} 实现电价平衡 (Ubide et al., 2019)。在高过冷度条件下, 熔体中 Al_2O_3 、 TiO_2 处于过饱和状态, 在单斜辉石骸晶中富集, 这一过程导致骸晶周围熔体中的 Al_2O_3 、 TiO_2 被消耗, 因此新生长的部分相对富集 SiO_2 、 MgO (Masotta et al., 2020)。同时, 随着过冷度的增高, 由晶体-熔体界面不平衡阳离子分配引起的晶体内部化学变化也会增大 (Mollo et al., 2013a)。因此, 无论是在实验岩石学还是天然样品的研究中, 具骸晶单斜辉石不同区域间的成分变化均大于低过冷度下形成的扇形环带单斜辉石 (Ubide et al., 2019; Masotta et al., 2020)。Zhou et al. (2021) 提出, 在碱性镁铁质岩浆体系中, 低程度过冷时 ($\Delta T = 13\text{--}45\text{ }^\circ\text{C}$), 单斜辉石所有晶面同时生长, 形成扇形环带; 高程度过冷时 ($\Delta T > 45\text{ }^\circ\text{C}$), 单斜辉石先形成初始骸晶骨架, 随后回填生长形成扇形环带。单斜辉石中的微量元素分配系数也受到过冷状态的影响。过冷状态下单斜辉石中会发生 M_2 位置上 Na^+ 替代 Ca^{2+} 及四面体位置上 A^{3+} 替代 Si^{4+} 等离子交换反应, 为维持电价平衡, 稀土元素、高场强元素等会更容易进入单斜辉石晶格, 其分配系数相应增大 (Mollo et al., 2013a)。

过冷条件通常反映岩浆升降温 (Chen et al., 2025; Ubide et al., 2019)、岩浆储库内高温岩浆受冷的围岩或具有温差的岩浆混合降温 (Xing et al., 2022, Cao et al., 2023)。因此, 发育扇形环带、骸晶结构的单斜辉石可以提供关于岩浆储库热状态演化、岩浆上升速率、火山喷发触发机制等关键信息 (Ubide et al., 2019; Cao et al., 2023; Chen et al., 2025)。

3 单斜辉石恢复结晶物理化学条件

单斜辉石成分对结晶物理化学环境极其敏感, 能够有效约束其结晶时的温度、压力、水含量、氧逸度等条件, 这些物理化学条件对于重建岩浆储运系统至关重要 (Nimis, 1995; Putirka et al., 1996, 2003; Putirka, 2008; Mollo et al., 2018; 张方毅, 2022; Wieser et al., 2025)。

3.1 单斜辉石结晶温度压力

长久以来，岩石学家们对于玄武质岩浆从源区到喷发过程中的运移路径具有争议，即其是直接从深部运移到浅部喷发还是经历了浅部岩浆储库的演化再喷发（Putirka et al., 1996）。火山岩中斑晶矿物的结晶温压计算在恢复岩浆储库深度、构建岩浆储运系统方面具有良好效果，能有效重现岩浆的运移路径。此外，岩浆储运系统中矿物元素扩散可为岩浆储库中的过程提供有效的时间尺度约束，而矿物结晶温度的精准约束是扩散年代学应用中不可或缺的前提（Petrone et al., 2016, 2018; Costa et al., 2020）。目前，通过单斜辉石成分计算结晶温度、压力的应用已较为成熟。研究者们基于单斜辉石晶体结构、热力学方程、机器学习等手段建立了诸多公式来恢复单斜辉石结晶时的温度和压力（图 5）。下文将对常用的单斜辉石结晶温压计算方法进行梳理。

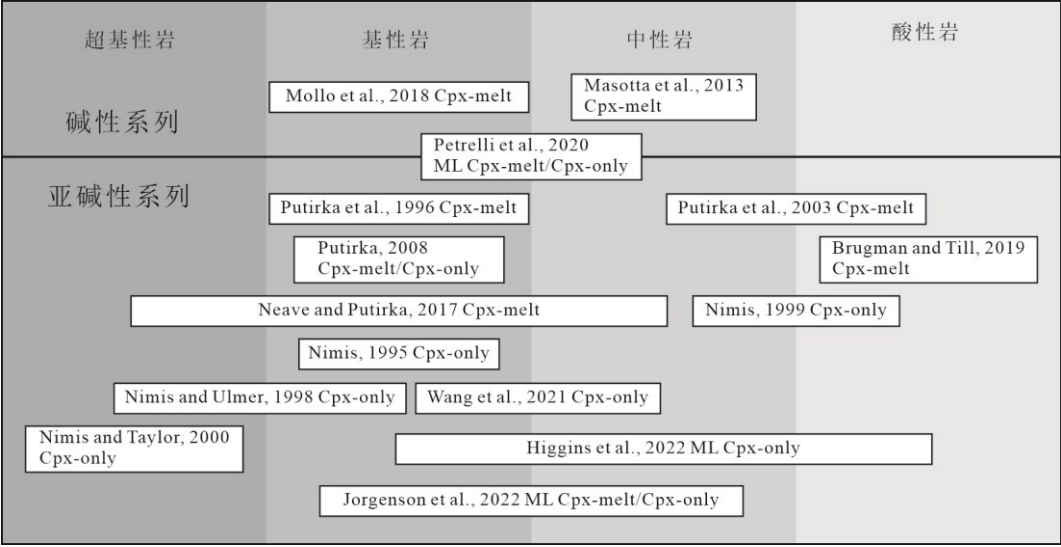


图 5 常用的单斜辉石温压计及其适用范围（“Cpx-melt”代表单斜辉石-熔体温压计；“Cpx-only”代表仅基于单斜辉石成分的温压计；“ML”代表机器学习温压计）

Fig. 5 Common clinopyroxene thermobarometers and their applicable conditions ("Cpx-melt" indicates the clinopyroxene-melt thermobarometer; "Cpx-only" the clinopyroxene-only thermobarometer; and "ML" the machine-learning thermobarometer)

3.1.1 单斜辉石成分温压计

单斜辉石晶胞（cell）与晶位（site）的体积对结晶压力极其敏感（Dal Negro

et al., 1989)。在压力固定的条件下, 单斜辉石晶胞体积 (V_{cell}) 与 M_1 位置体积 (V_{M1}) 呈线性负相关; 在熔体成分固定时, 单斜辉石晶胞体积 (V_{cell}) 及 M_1 位置体积随结晶压力增加而线性降低。这些体积变化涉及单斜辉石 M_1 、 M_2 及 T 位置的化学成分交换, 因此可以应用于恢复单斜辉石的结晶压力 (Nimis, 1995; Nimis and Ulmer, 1998; Nimis, 1999)。Nimis (1995) 基于晶胞体积与 M_1 位置体积之间的关系, 结合实验单斜辉石的晶体结构模拟, 建立了适用于除高铝玄武岩之外的无水玄武质熔体的单斜辉石压力计, 该压力计的压力范围为 0–24 kbar。Nimis and Ulmer (1998) 和 Nimis (1999) 通过收集更加广泛的数据集, 将压力计的适用范围扩展到无水到含水体系, 分别适用于基性到超基性岩和中性到酸性岩。Nimis and Taylor (2000) 利用地幔橄榄岩中单斜辉石与石榴子石之间的 Cr 交换及单斜辉石与斜方辉石之间的 En 组分 ($\text{Mg}_2\text{Si}_2\text{O}_6$) 交换, 分别建立了单斜辉石成分压力计和单斜辉石成分温度计。Putirka (2008) 对 Nimis (1995) 的压力计和 Nimis and Taylor (2000) 的温度计进行了改进。Wang et al. (2021) 应用近些年来报道的涵盖更广地壳压力范围的数据集对以往的公式进行重新校准, 提高了单斜辉石成分温压计的精度。

3.1.2 单斜辉石-熔体温压计

矿物-熔体温压计的核心原理建立在矿物与熔体间平衡反应的热力学关系上, 即通过量化温度、压力发生变化时平衡反应的焓变和熵变来实现 (Mollo et al., 2018)。平衡反应的平衡常数与吉布斯自由能的关系可以表示为:

$$-RT \ln K_{eq} = \Delta G_r^0 \quad (1)$$

公式 (1) 中, R 、 T 、 K_{eq} 、 ΔG_r^0 分别为气体常数、温度、给定平衡反应的平衡常数及参考压力 (P_0) 和温度条件下反应达到平衡时的吉布斯自由能变化。将 ΔG_r^0 用反应过程中的焓变 ($\Delta H_r^{P_0}$)、熵变 ($\Delta S_r^{P_0}$)、体积变 (ΔV_r) 来表示并重新排列公式 (1):

$$\ln K_{eq} = -(\Delta H_r^{P_0} / RT) + (\Delta S_r^{P_0} / R) - ((P - P_0)\Delta V_r / RT) \quad (2)$$

公式 (2) 中 P 表示反应实际压力。取 $P_0=0$, 将压力 (P) 作为因变量, 重新

排列公式（2）：

$$P = -(RT \ln K_{eq} / \Delta V_r) + (\Delta S_r^{P_0} T / \Delta V_r) - (\Delta H_r^{P_0} / \Delta V_r) \quad (3)$$

将 $1/T$ 作为因变量，重新排列公式（2）：

$$1/T = -R(\Delta H_r^{P_0} + P\Delta V_r) \ln K_{eq} + \Delta S_r^{P_0}(\Delta H_r^{P_0} + P\Delta V_r) \quad (4)$$

在建立公式（3）和公式（4）的基础上，通过对已知 K_{eq} 、 P 、 T 的实验数据集进行回归分析，得到剩余未知量并再次带入公式（3）和公式（4），即可计算出自然样品中单斜辉石的结晶温度和压力（Putirka et al., 1996; Putirka, 2008）。

单斜辉石-熔体平衡的压力敏感性主要来自体积变化（Putirka, 2008）。玄武质体系中 Na 和 Al 的偏摩尔体积较大（Lange and Carmichael, 1987），而单斜辉石中 Jd ($\text{NaAlSi}_2\text{O}_6$) 组分的偏摩尔体积较小（Putirka et al., 1996）。反应物与生成物之间的这种体积差，使得 Jd 组分对压力极为敏感（Robie et al., 1979），因此被用于标定单斜辉石-熔体压力计（Putirka et al., 1996）。单斜辉石-熔体平衡反应中 Jd-DiHd ($\text{CaMgSi}_2\text{O}_6$) 交换和 CaTs ($\text{CaAl}^{\text{VI}}\text{Al}^{\text{IV}}\text{SiO}_6$)-DiHd 交换过程中压力变化很小，对温度有着极强的依赖，因此被用来标定单斜辉石-熔体温度计（Putirka et al., 1996）。Putirka et al. (1996) 最初建立了主要针对玄武质岩石的单斜辉石-熔体温压计。为使应用的岩性范围更加宽泛以了解不同性质岩浆的储运系统，后续研究者们逐步扩大了数据集范围，建立了适用于不同体系的单斜辉石-熔体温压计（如 Putirka et al., 2003; Putirka, 2008; Masotta et al., 2013; Neave and Putirka, 2017; Mollo et al., 2018）。其中，Putirka et al. (2003) 建立的单斜辉石-熔体温压计适用于更演化的含水体系。Putirka (2008) 对已发表的温压计进行了综合评估，对 Putirka et al. (1996, 2003) 建立的单斜辉石-熔体温压计进行了重新校准，进一步提高了其精度。此外，Putirka (2008) 还建立了基于单斜辉石与平衡熔体 Al 配分的单斜辉石-熔体压力计 (Eqn. 32c)。Masotta et al. (2013) 利用演化的碱性岩浆体系数据集，对 Putirka et al. (1996) 和 Putirka (2008) 的公式进行校准，建立了专用于演化碱性岩浆体系的单斜辉石-熔体温压计。Neave and Putirka (2017) 通过实验数据集对 Putirka et al. (1996) 和 Putirka (2008) 建立的单斜辉石-熔体压力计进行了评估，发现这些压力在高压条件下具有较高的精度，但是在 <7 kbar 时这些压力计常常高估了单斜辉石的结晶压力。据此 Neave and

Putirka (2017) 对 1 atm 到 20 kbar 压力范围的实验数据进行校准, 建立了精度更高的单斜辉石-熔体压力计, 适用于含水和无水体系下的超镁铁质岩到中性岩。Mollo et al. (2018) 将已发表的经典压力计 (Eq. P1, Putirka et al., 1996; Eq. A, Putirka et al., 2003; Eqs. 30, 31, 32a, 32b, Putirka, 2008; Eq. 9, Neave and Putirka, 2017) 和温度计 (Eq. T1, Putirka et al., 1996; Eq. B, Putirka et al., 2003; Eqs. 32d, 33, 34, Putirka, 2008) 应用于镁铁质碱性岩浆体系数据集, 得到的结果误差较大 (压力和温度标准误差最高分别达 4.06 kbar 和 86 K)。Mollo et al. (2018) 对表现最佳的压力计 (Eq. A, Putirka et al., 2003) 和温度计 (Eq. 33, Putirka, 2008) 进行了重新校准, 得到了适用于镁铁质碱性岩浆体系的单斜辉石-熔体温压计。在高硅体系中, 单斜辉石的四面体位置主要为 Si 而缺乏 Al (Salviulo et al., 2000), Brugman and Till (2019) 根据单斜辉石-熔体 CaTs-DiHd 交换建立了适用于高硅低铝体系的单斜辉石-熔体温度计。尽管目前已经有大量单斜辉石-熔体温压计被建立出来, 但这些温压计的精度仍然有待进一步提高。Wang et al. (2025) 从热力学角度分析了单斜辉石-熔体温压计的误差来源, 发现温度计和压力计的误差主要来源分别为成分数据误差及热力学模型参数, 这为后续建立更高精度单斜辉石温压计提供了有效思路, 即以高质量数据集及更严谨热力学模型为基础。

单斜辉石-熔体温压计的应用前提是找到处于化学平衡的单斜辉石及熔体成分。常用的平衡熔体获取方法有如下几种: (1) 选择单斜辉石斑晶周围未结晶的玻璃作为平衡熔体 (Neave and Putirka, 2017)。玻璃成分最能代表真实平衡熔体成分, 不过其在结晶度较高的岩石中较为缺乏。(2) 以全岩成分近似代表平衡熔体。需要注意的是, 由于潜在的捕虏体、捕虏晶等外来组分的干扰, 全岩成分是否能代替真实的平衡熔体需要进一步考量 (Zhang et al., 2024a; 2025)。当全岩成分受到外来组分强烈改造时, 可以通过向全岩组分添加或减去橄榄石、单斜辉石、斜长石等矿物成分, 以恢复真实的平衡熔体成分 (Armienti et al., 2013)。

(3) 从全球地球化学数据库 (如 GEOROC) 中获取与单斜辉石成分相平衡的熔体数据 (如 Xing and Wang, 2020)。

在找到潜在平衡熔体后, 需对其平衡状态进行检验 (Putirka et al., 1996; Putirka, 1999, 2008)。Roeder and Emslie (1970) 基于平衡方程

$\text{MgO}^{\text{Ol}} + \text{FeO}^{\text{liq}} = \text{MgO}^{\text{liq}} + \text{FeO}^{\text{Ol}}$ 建立了橄榄石与熔体之间的 Fe-Mg 交换系数 ($K_D(\text{Fe-Mg})^{\text{Ol-liq}}$)，以此检验橄榄石与熔体的平衡状态。后来，矿物与熔体间的 Fe-Mg 交换系数后来也被广泛用于检验单斜辉石、斜方辉石等其他矿物与熔体之间的平衡状态 (Putirka et al., 2003; Putirka, 2008)。在镁铁质岩浆体系中，处于平衡状态的 $K_D(\text{Fe-Mg})^{\text{cpx-liq}}$ 值为 0.28 ± 0.08 (Putirka, 2008)。然而，当岩浆体系中的过冷状态控制单斜辉石快速生长时，Fe 和 Mg 的分配系数比值 ($D_{\text{Fe}}/D_{\text{Mg}}$) 趋于恒定 (Mollo and Hammer, 2017)。如前面 2.3 小节中所述，随着过冷度提高，单斜辉石-熔体体系中伴随着众多元素的不平衡分配，这意味着用 Fe-Mg 交换系数不能可靠地监测晶体快速生长过程中的平衡状态 (Mollo and Hammer, 2017; Mollo et al., 2018; 罗雕等, 2022)。据此，Mollo et al. (2013b) 提出通过对比基于单斜辉石成分及熔体成分预测的单斜辉石端元 (DiHd, EnFs, Jd, CaTiTs)，当两者相差 (如 $\Delta \text{DiHd} = \text{DiHd}_{\text{预测}} - \text{DiHd}_{\text{实测}}$) 趋近于 0 时，则视为单斜辉石与熔体达到平衡。单斜辉石端元实测值和熔体成分预测值的计算方法可参考 Putirka et al. (1996) 和 Putirka (1999)。随着过冷度增加，单斜辉石会倾向于富集 Al^{IV} 、Ti、Na，亏损 Fe^{2+} 、Ca、Mg，从而导致 DiHd 成分的占比降低，而 Ts、Jd、En、Fs 成分占比增加 (Mollo et al., 2010, 2011, 2012)。因此，基于单斜辉石预测与实测端元成分差值的方法能够有效监测过冷条件下的平衡偏移 (Mollo et al., 2018)。

3.1.3 机器学习单斜辉石温压计

近年来，大数据及机器学习与地质学研究的结合逐渐趋于紧密，其在矿物温压计方面的应用同样得到了快速发展 (Petrelli et al., 2020; 李晓彦和张超, 2023; Petrelli, 2024)。机器学习的特点是无需依赖热力学反应等已定义的概念模型，而是直接基于大型数据集来揭示数据内在规律 (Petrelli et al., 2020; Petrelli, 2024; 邓李涛等, 2025)。Petrelli et al. (2020) 收集了 1403 组实验产生的单斜辉石及平衡熔体成分，采用线性回归、梯度提升、极端随机树、随机森林、K-近邻、决策树六种机器学习算法进行拟合，建立了适用范围为 0–40 kbar、952–1882 K 的单斜辉石-熔体温压计。此外，该研究还建立了仅依靠单斜辉石成分、独立于熔体和

温度的压力计，克服了依靠交换反应的压力计对共存平衡熔体的依赖。Higgins et al. (2022) 使用随机森林算法拟合出了单斜辉石成分温压计并将该方法应用于 Mount Liamuiga 火山，有效揭示了岩浆喷发前的温度、压力及成分演化路径。Jorgenson et al. (2022) 进一步评估了使用极端随机树模型建立单斜辉石温压计的注意事项，包括单斜辉石-熔体不平衡数据的剔除，模型超参数的优化，以及集成算法造成的模型误差。在实际研究中，机器学习单斜辉石温压计应用日益广泛，其与基于热力学平衡的传统温压计结合应用，也取得了良好效果。如 Liu et al. (2023) 应用 Neave and Putirka (2017) 建立的单斜辉石-熔体温压计和 Higgins et al. (2022) 建立的机器学习温压计，计算了中亚造山带东南部狼山弧侵入杂岩中单斜辉石的结晶温度与压力，两种方法得到的温度和压力范围分别为 951–1012°C、2.8–7.6 kbar 和 938–1057°C、1.9–8.0 kbar，在误差范围内具有较好的一致性。

3.1.4 定性约束单斜辉石结晶温压的经验公式

除了上述定量计算公式，还有一些定性约束单斜辉石结晶温度、压力的经验公式。Soesoo (1997) 通过对大量单斜辉石实验数据的主量元素成分进行多元统计分析，建立了适用于 1100–1300°C、0–20kbar 的单斜辉石结晶温压判别经验公式（图 6a–b）。Velde and Kushiro (1978) 发现在 Na-Al-Si-O 熔体中，随着压力升高，Al 的配位环境会从四次配位 (Al^{IV}) 逐渐向六次配位 (Al^{VI}) 转变。Wass (1979) 在碱性玄武岩中识别出四类形成于不同压力条件的单斜辉石，并发现随着结晶压力降低， $\text{Al}^{\text{IV}}/\text{Al}^{\text{VI}}$ 比值显著升高，由此将单斜辉石 $\text{Al}^{\text{IV}}/\text{Al}^{\text{VI}}$ 用作约束结晶压力相对大小的可靠指标（图 6c）。不过，Wass (1979) 强调该方法仅适用于同一岩浆系统内不同单斜辉石的对比，以消除不同地区岩浆成分差异带来的干扰。在实际研究中，这些定性的经验公式具有两方面的意义：（1）在因平衡熔体难以获取、定量温压计不适配样品岩石体系等原因而无法定量计算结晶温压的情况下，可以采用经验公式对结晶温压的相对大小进行定性约束，这对于理解岩浆储运系统中的演化过程同样具有积极意义。（2）为了提升结果可靠性，可将经验公式定性约束与定量计算结果进行对比验证（如蔡晓芸等, 2024; Xia et al., 2026）。

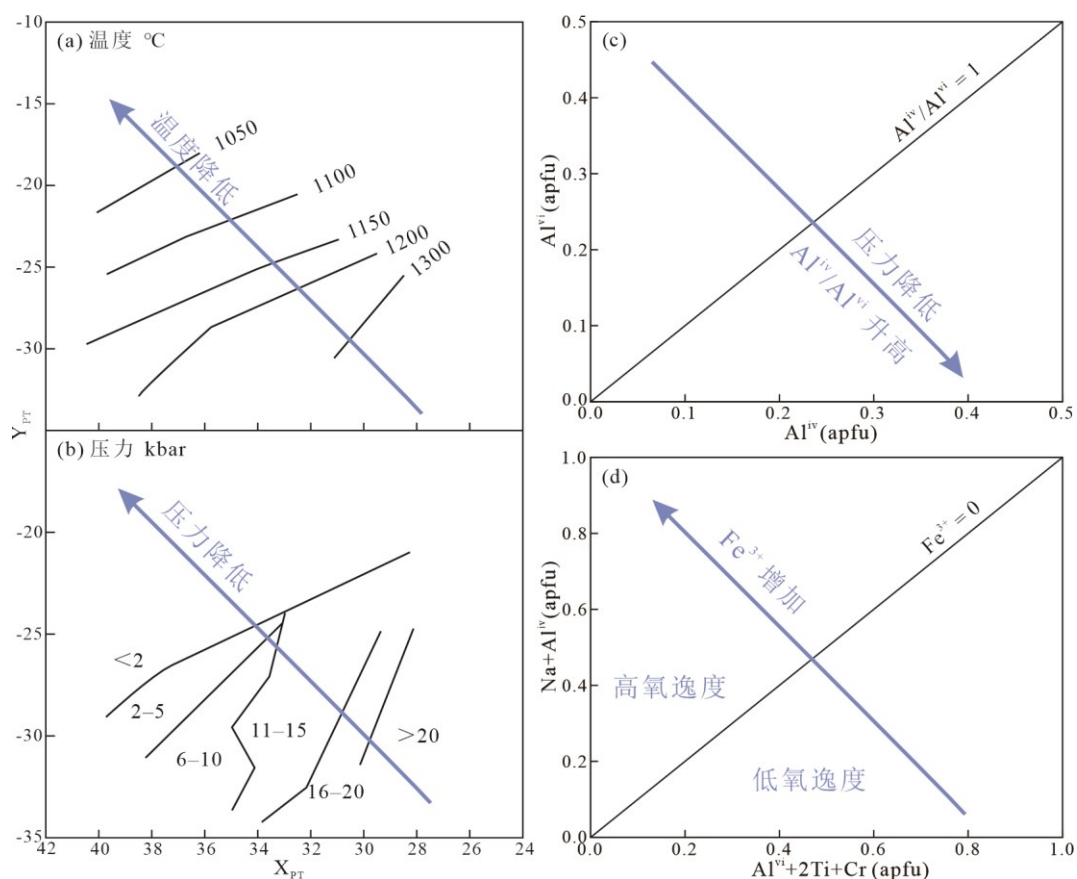


图 6 (a-b) 根据主量元素成分多元统计分析得出的单斜辉石结晶温度、压力判别经验公

式，其中 $X_{PT} = 0.446SiO_2 + 0.187TiO_2 - 0.404Al_2O_3 + 0.346FeO (tot) - 0.052MnO +$

$0.309MgO + 0.431CaO - 0.446Na_2O$, $Y_{PT} = -0.369SiO_2 + 0.535TiO_2 - 0.317Al_2O_3 + 0.323FeO$

$(tot) + 0.235MnO - 0.516MgO - 0.167CaO - 0.153Na_2O$ (据 Soesoo, 1997) (c) 根据

Al^{IV}/Al^{VI} 值的单斜辉石结晶压力相对高低判别图 (据 Wass, 1979); (d) 单斜辉石 $Na +$

$Al^{IV} - Al^{VI} + 2Ti + Cr$ 氧逸度相对高低判别图 (据 Schweitzer et al., 1979)

Fig.6 (a-b) Empirical discrimination diagrams for estimating crystallization temperature and

pressure of clinopyroxene based on multivariate statistical analysis of major element

compositions. $X_{PT} = 0.446SiO_2 + 0.187TiO_2 - 0.404Al_2O_3 + 0.346FeO (tot) - 0.052MnO +$

$0.309MgO + 0.431CaO - 0.446Na_2O$, $Y_{PT} = -0.369SiO_2 + 0.535TiO_2 - 0.317Al_2O_3 +$

$0.323FeO (tot) + 0.235MnO - 0.516MgO - 0.167CaO - 0.153Na_2O$ (after Soesoo, 1997). (c)

Diagram showing relative crystallization pressures of clinopyroxene based on Al^{IV}/Al^{VI} ratios

(after Wass, 1979). (d) Diagram showing relative oxygen fugacities based on $Na + Al^{IV} - Al^{VI} + 2Ti$

$+ Cr$ in clinopyroxene (after Schweitzer et al., 1979).

3.2 单斜辉石成分约束熔体水含量

水是岩浆体系中主要的挥发份之一,对其含量进行定量计算对于约束地球深部动力学过程、岩浆演化、火山喷发、成岩成矿等方面具有重要意义(Ni et al., 2017; Wei et al., 2026; 林嘉敏等, 2026)。Armienti et al. (2013) 根据意大利 Etnean 熔岩中的单斜辉石成分及实验数据,建立了适用于压力为 0.8 GPa 以下的粗面玄武岩体系的单斜辉石-熔体湿度计。Perinelli et al. (2016) 通过增加更多的实验样品及天然样品中的单斜辉石成分,对 Armienti et al. (2013) 建立的单斜辉石-熔体湿度计进行了改进,扩展了该模型的应用范围。Wei et al. (2026) 应用机器学习的方法对 316 个已知平衡熔体水含量的实验单斜辉石成分进行了分析,建立了目前最新的单斜辉石湿度计。Wei et al. (2026) 将该方法应用到峨眉山大火成岩省,揭示了其富水的地幔源区。除此之外,还可利用傅里叶红外光谱或二次离子探针直接测试单斜辉石中的水含量,再利用水在单斜辉石和熔体中的分配系数反算熔体水含量(Lloyd et al., 2016; Wang et al., 2020)。

3.3 单斜辉石成分约束氧逸度

氧逸度是岩浆体系中除温度、压力之外的又一重要热力学参数,影响着地幔部分熔融、岩浆分异、元素富集成矿等众多地质过程(张方毅, 2022; 辛雨等, 2023; Zhang et al., 2024b, 2025)。单斜辉石与熔体间钨的分配系数与氧逸度强烈相关,基于这一相关性, $D_V^{Cpx/melt}$ 氧逸度计的建立得以实现(Canil, 2002; Wang et al., 2019; 张方毅, 2022)。张方毅(2022)使用其新校正的 $D_V^{Cpx/melt}$ 氧逸度计对南秦岭早古生代一套不同演化程度的碱性岩的氧逸度进行了计算,发现随着结晶压力降低及演化程度增加,氧逸度具有升高的趋势,表明在岩浆储运系统中存在着氧逸度的变化。岩浆储运系统中氧逸度的变化可能涉及单斜辉石、橄榄石、石榴子石等矿物分离结晶(Cottrell and Kelley, 2011; Tang et al., 2018)、不同种类挥发分的脱气作用(Mathez, 1984; Holloway, 2004; Jugo et al., 2010)等机制,这一方面还有待于进一步约束。岩浆体系中 Fe 价态($Fe^{3+}/\Sigma Fe$ 或 Fe^{3+}/Fe^{2+})也是反映氧逸度的关键指标(张超和李晓彦, 2025)。单斜辉石中的 Fe^{3+} 含量相对高低可

以通过电荷亏损与电荷过剩的平衡来推断，M₂ 位置的 Na⁺替代 Ca²⁺及四面体位置的 Al³⁺替代 Si⁴⁺会造成电荷亏损，而八面体位置的 Al³⁺、Cr³⁺、Fe³⁺、Ti⁴⁺替代 Mg²⁺或 Fe²⁺则会造成电荷过剩（Schweitzer et al., 1979）。因此，可以根据单斜辉石 Na+^{IV}Al-^{VI}Al+2Ti+Cr 图判别氧逸度的相对高低（图 6d）（Schweitzer et al., 1979）。但需注意，该方法仅是一种定性约束。目前，电子探针 Fe 价态原位分析方法已逐渐成熟（Cao et al., 2025; 张超和李晓彦, 2025）。Cao et al. (2025) 利用电子探针“峰肩法”，以 9 个不同成分的单斜辉石为对象成功标定了其 Fe³⁺/ΣFe 比值，这些标样开发为后续实现从 Fe 价态角度定量约束大量天然样品的氧逸度提供了重要支持。

4 单斜辉石 Fe-Mg 扩散：刻画时间尺度

单颗矿物上的元素变化除了受熔体演化影响，还受控于矿物尺度上的元素扩散。当矿物经历不同熔体作用时，会处于不平衡状态，从而诱发矿物的元素扩散以达到再平衡。扩散发生时，矿物从核到边不同位置的元素浓度变化为与时间有关的非稳态扩散，这一过程可用菲克第二定律（Fick's second law）来表示（Fick, 1855; Chakraborty, 2008）：

$$\partial C / \partial t = (\partial / \partial x)(D \partial C / \partial x) \quad (5)$$

其中，C 为元素浓度，t 为扩散时间（s），x 为距离（m），D 为扩散系数（m²/s）。在实际应用中，某一矿物中某一元素的浓度可以通过电子探针测得，扩散剖面的距离及测点间的距离也易获取。温度是影响扩散的主导因素，因此，只要正确计算扩散发生时的温度，即可获得准确的扩散系数，结合矿物剖面的元素浓度及距离，便能有效刻画岩浆储库中演化过程的时间尺度（Costa et al., 2020; 陈祖兴等, 2020; 陈厚彬等, 2022; 田野和侯通, 2022; 李扬等, 2024）。

在单斜辉石中，常用于扩散研究的元素是 Fe-Mg 体系。单斜辉石中 Fe-Mg 体系具有比橄榄石、斜长石等矿物中 Fe-Mg 体系更高的活化能，即其对温度的变化更敏感，这使得单斜辉石 Fe-Mg 扩散能够灵敏记录温度变化较小的扩散过程（Dimanov and Sautter, 2000; Petrone et al., 2016, 2018）。Müller et al. (2013) 报道了单斜辉石中 Fe-Mg 扩散的扩散系数为：

$$D^{\text{Fe-Mg}} = 2.77 \pm 4.27 \times 10^{-7} \exp(-320.7 \pm (16 \text{ kJ/mol})/RT) \text{ m}^2/\text{s}$$

(6)

其中 R 为气体常数, T 为温度 (K)。如 3.1 小节中所述, 目前有一系列成熟的方法约束单斜辉石的结晶温度, 为确定扩散系数提供了有力支撑。Petrone et al. (2016, 2018) 通过对 Stromboli 火山中具有复杂环带单斜辉石进行 Fe-Mg 扩散计算, 发现单斜辉石在高温岩浆储库中具有数天到 9 年的短暂停留时间, 而在低温岩浆储库中则具有达数十年的更长的停留时间。Pelullo et al. (2024) 通过 Fe-Mg 扩散对意大利南部 Neapolitan 火山区 Ischia 岛喷发物中的单斜辉石进行了时间尺度的约束, 揭示了在数小时内发生了基性岩浆对粗面质岩浆的补给事件, 而后续的驻留时间可达十年, 表明被扩散环带所记录的基性岩浆注入并未直接出发火山的喷发。

5 单斜辉石微量元素对于岩浆源区性质的约束

在全岩地球化学研究中, 微量元素能够有效揭示岩浆源区性质。幔源岩石中重稀土元素的含量及比值受源区石榴子石的出现与否控制 (Prelević et al., 2012)。尖晶石相地幔中 La/Dy 与 Yb 分异程度不明显, 且 Yb 含量较高, 而石榴子石相地幔中 Yb 含量低, La/Dy 与 Yb 呈现明显分异, 因此 La/Yb、Dy/Yb 比值及 Yb 含量能够有效区分地幔源区的尖晶石相与石榴子石相 (Miller et al., 1999; Prelević et al., 2012)。由于大离子亲石元素在流体中相对活跃, 稀土元素和高场强元素则在熔体中更加活跃, 因此 Ba/La、Ba/Yb、La/Yb、Th/Nb 等比值可区分源区的流体交代及熔体交代作用 (Pearce and Peate, 1995; Elliott et al., 1997; Class et al., 2000)。幔源碳酸岩通常表现出富集 Sr、Ba、LREE 等元素, 亏损 Nb、Ta、Zr、Hf、Ti 等元素 (Hoernle et al., 2002; 杨航, 2023; Xia et al., 2024)。幔源碳酸岩的出现可以指示源区的碳酸盐熔体交代作用, 因此 $(\text{La/Yb})_{\text{N}}$ 、Ti/Eu、Hf/Sm、Zr/Hf 等元素比值可用于识别交代介质为碳酸盐熔体还是硅酸盐熔体 (Rudnick et al., 1993; Coltorti et al., 1999; Dai et al., 2017, 2018; Zhang et al., 2020)。

如前所述, 相比起橄榄石和斜长石, 单斜辉石对于稀土元素和高场强元素具有更高的相容性 (D'Orazio et al., 1998; Sun and Liang, 2014; Mollo et al., 2018)。

因此单斜辉石的微量元素成分能够有效的记录熔体成分信息。需要强调的是，低 $Mg\#$ 单斜辉石可能受到地壳混染或岩浆分异的影响，因此通常通过高 $Mg\#$ 自生晶单斜辉石来反映地幔源区信息。在获得单斜辉石微量元素含量的基础上，结合元素在单斜辉石与熔体间的分配系数进行反算，就可得到与之平衡的熔体微量元素组成，进而约束岩浆源区性质（如 Prelević et al., 2012; Xing and Wang, 2020; Zhang et al., 2023; Ma et al., 2025; Xia et al., 2026）。Xing and Wang (2020) 计算了北天山超钾质煌斑岩脉中 $Mg\#$ 值大于 85 的单斜辉石自生晶的平衡熔体，发现其平衡熔体具有变化的 La/Yb (13.6–37.9) 和 Dy/Yb (1.56–3.02) 比值及变化的 Yb 含量，表明其来源于石榴子石相和尖晶石相的混合地幔源区。Ma et al. (2025) 发现青藏高原东南缘新生代超钾质火山岩中单斜辉石自生晶的平衡熔体表现为低 Ba/La 比值及高 La/Yb 比值，表明其源区经历了熔体交代而非流体交代，此外，平衡熔体低的 Hf/Sm 比值及 Ti/Eu 比值指示交代介质为碳酸盐熔体。相较于全岩成分分析的长周期性，单斜辉石原位微量元素分析具有快速高效的优势。因此通过单斜辉石成分反演地幔源区性质可以作为地幔岩石研究的一个有效手段。

6 总结与展望

随着岩石学理论体系的不断发展和分析测试技术的不断进步，开展更精细的矿物学工作成为迫切需要并逐渐可行。作为从酸性岩到超基性岩中常见的造岩矿物，单斜辉石可以很好地记录温度、压力等结晶时的关键信息，并保存岩浆演化过程中各种复杂过程留下的痕迹。通过研究火山岩中的类型多样的单斜辉石，可以约束岩浆源区性质、识别出不同深度的岩浆储库、重建岩浆储运系统、并精细刻画岩浆演化过程。除有助于理解地球上幔源岩石的成因及演化过程外，单斜辉石也是月球、火星等地外星体样品中的常见矿物（Lou et al., 2023; Guo et al., 2024; Li et al., 2026），这为破译地外星体的起源与演化也提供了重要的突破口。例如，Luo et al. (2023) 在嫦娥五号月壤玄武岩碎屑中识别出三类单斜辉石，对其岩浆储运系统、热演化及驻留时间等进行了精细刻画。Li et al. (2026) 在月球陨石中识别出具有 Cr 突变的单斜辉石，证实月球岩浆储运系统中存在岩浆补给等动态演化过程。值得注意的是，由于地外样品较为稀缺且颗粒细小，矿物学原

位分析工作也更具可行性。

对于未来的研究展望主要有以下几个方面：（1）温压计精度的提升及体系的扩展。现有的单斜辉石温压计的精度仍有提升空间，且针对硅饱和长英质岩浆体系和硅极端不饱和碱性体系的温压计仍然较为缺乏。因此，建立精度更高、适用体系更广的单斜辉石温压计是下一阶段研究需要关注的方向。（2）单斜辉石湿度计、氧逸度计的扩展。相较于温压计，高精度单斜辉石湿度计和氧逸度计的约束仍然较少，亟需进一步校准与完善，从而实现对岩浆储运系统物理化学状态演化的更精准约束。（3）单斜辉石原位微区同位素分析的突破。开发更多元的单斜辉石原位同位素分析方法（如 O、Li、Mg 等）及相应的标准物质，将有效示踪岩浆储运系统内部的多元物质来源，揭示岩浆混合与同化混染等过程。（4）研究对象的扩展。在应用层面，未来需将研究扩展至更多样的构造背景、不同岩性乃至地外样品（如月球、火星）的对比研究，这将有助于深入理解地球系统及地外星体内部的岩浆演化机制。

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