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青海湖地区晚第四纪黄土的物质来源

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摘要: 青海湖地区的晚第四纪黄土记录了湖区晚第四纪以来的环境和气候变化。迄今为止,对于青海湖地区晚第四纪黄土物质来源的研究较薄弱。以青海湖东岸的种羊场晚第四纪风成沉积剖面为主要研究对象,在青海湖区及其周边采集了黄土、风成砂、湖相沉积、河流沉积等样品,结合黄土高原西部临洮黄土样品,对它们的元素组成($<75\ \mu\text{m}$ 的硅酸盐组分)进行了对比研究。初步结果表明:(1) $\text{K}_2\text{O}/\text{Al}_2\text{O}_3$ (摩尔比)和 Zr/Ti 、 Zr/Nb 比值显示青海湖地区的风成沉积显著区别于本区的河流沉积和湖相沉积;(2)青海湖地区的晚第四纪黄土与黄土高原西部临洮黄土的 $\text{K}_2\text{O}/\text{Al}_2\text{O}_3$ 和 Zr/Ti 、 Zr/Nb 比值相一致;(3)青海湖地区的晚第四纪黄土可能来自柴达木盆地。

关键词: 青海湖;黄土;物质来源;元素地球化学。

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Provenance of the Late Quaternary Loess Deposit in the Qinghai Lake Region

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Abstract: Late Quaternary loess deposit is widely distributed in the Qinghai Lake region and adjacent area. The loess deposit independently recorded the environmental and climatic changes during the late Quaternary in the Qinghai Lake region, northeastern Qinghai-Tibetan Plateau. Nearly all studies are focused on reconstructing the environmental and climatic changes recorded by loess deposit in the Qinghai Lake region. However, up to now, the provenance of the loess deposit in the Qinghai Lake region is still poorly understood. Here we present the elemental concentration of the silicate fraction of the eolian deposit ($<75\ \mu\text{m}$) from the ZYC section in the Qinghai Lake region and LT section at Lintao County on the western Chinese Loess Plateau, loess deposits at Guanjiashan Mountain, Chaka town and Wulan County, eolian sands at western and eastern shore of the Qinghai Lake, and fluvial deposits from the Buhuhe river and lacustrine deposits at Erlangjian site in the Qinghai Lake region. The results show that: (1) $\text{K}_2\text{O}/\text{Al}_2\text{O}_3$ (molar ratio) and Zr/Ti , Zr/Nb ratios indicate that the eolian deposits in the Qinghai Lake region and adjacent area can be distinguished clearly from the local deposits represented by river deposits at Buhuhe river and lacustrine deposits at Erlangjian in the drainage basin of Qinghai Lake; (2) There is similarity of the elemental ratios ($\text{K}_2\text{O}/\text{Al}_2\text{O}_3$ and Zr/Ti , Zr/Nb) between the eolian deposits in the Qinghai Lake region and loess deposits at Lintao County. Thus, the eolian deposits in the Qinghai Lake region and Lintao County may have the same source region; (3) Late Quaternary loess deposit in the Qinghai Lake region is probably sourced from the Qaidam basin.

Key words: Qinghai Lake; loess; provenance; elementary geochemistry.

黄土高原黄土物质来源的研究对于揭示亚洲内陆的荒漠化、古大气环流格局等具有重要意义。虽然当前对我国黄土高原主体地区黄土的物源已经有过

多年的研究,但是依然存在很大的争议(Sun, 2002; Pullen *et al.*, 2011; Xiao *et al.*, 2012; Che and Li, 2013)。其中一个突出的分歧是柴达木盆地是否为黄

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土高原的主要源区之一(Sun, 2002; Pullen *et al.*, 2011).这一分歧涉及到搬运黄土的风动力问题:黄土主要是由冬季风搬运还是同时包含西风的贡献?

青海湖作为我国内陆最大的封闭湖泊,是青藏高原东北部极为重要的生态屏障.青海湖的湖相沉积记录了上新世以来特别是晚第四纪青藏高原东北部的环境和气候演化历史(Shen *et al.*, 2005; Yu and Zhang, 2008; Liu *et al.*, 2010, 2012; An *et al.*, 2012; Fu *et al.*, 2013).青海湖地区广泛发育晚第四纪黄土.黄土独立地记录了青海湖地区晚第四纪以来的环境和气候变化过程,可以和湖泊沉积记录的信息相互检验,是青海湖地区古环境研究的重要沉积记录.青海湖地区的黄土主要分布在湖岸阶地、河流阶地或冲积砾石层上,起始年代约为 14 ka(Lu *et al.*, 2011; Liu *et al.*, 2012).青海湖东岸种羊场风成沉积剖面古土壤—黄土(30~100 cm 深)的中值粒径为 10~50 μm ,大于 63 μm 的组分为 11%~46%.

过去青海湖晚第四纪黄土的研究主要关注这套沉积记录的古环境和古气候信息(周笃珺等, 2004; 王建国等, 2005; 赵存法等, 2009; Lu *et al.*, 2011; 胡梦珺等, 2012; Liu *et al.*, 2012; 鄂崇毅等, 2013; 侯光良等, 2013; 尚媛等, 2013; 曾方明等, 2015b),鲜有对于这套沉积物质来源的研究,一定程度限制了青海湖地区风成沉积与古环境变化的研究.

由于青海湖地区紧邻柴达木盆地地下风向并处于柴达木盆地向黄土高原传输粉尘的可能路径之上,青海湖地区晚第四纪黄土的物质来源研究有助于解决柴达木盆地是否可以作为黄土高原黄土的主要源区这一重要分歧.本文以青海湖东岸的种羊场晚第四纪风成沉积剖面为主要研究对象,在青海湖区及其周边采集了黄土、风成砂、湖相沉积、河流沉积等样品,结合黄土高原西部临洮黄土样品,对它们的元素组成进行对比研究,以期探讨青海湖地区晚第四纪黄土的物质来源.

1 样品与方法

为了通过元素地球化学手段示踪青海湖地区晚第四纪黄土的物质来源,在青海湖地区及其周边和黄土高原西部临洮县采集了风成沉积剖面样品、黄土、风成砂、河流和湖泊沉积样品(图 1),共计 20 个.各类样品的信息如下:

(1)风成沉积剖面:青海湖东岸种羊场风成沉积剖面(36°38'N, 100°52'E)(Liu *et al.*, 2012; 曾方明

等, 2015b),剖面深为 120 cm. ZYC8 为现代砂质土壤, ZYC18、ZYC25、ZYC30 为古土壤, ZYC43、ZYC46 为黄土, ZYC57、ZYC60 为风成砂.临洮黄土剖面(35°22'N, 103°55'E)发育典型的晚更新世黄土—古土壤序列(曾方明等, 2007, 2014a; Zeng *et al.*, 2011), LT15 为全新世古土壤, LT36 为末次冰期晚期黄土.

(2)黄土:CK 样品(36°47'N, 99°7'E)取自茶卡镇,黄土披覆于冲积砾石层之上; WL 样品(36°57'N, 98°24'E)取自乌兰县,黄土发育于河流阶地之上; GJS 样品(37°13'N, 98°56'E)取自关角山山坡.

(3)风成砂:HD-1(36°43'N, 100°47'E)和 HD-2(36°44'N, 100°47'E)取自青海湖湖东沙丘; SNH 风成砂(36°59'N, 99°36'E)取自石乃亥乡出露的风成沉积地层下部.

(4)河流沉积: BHH-1 和 BHH-2(37°02'N, 99°44'E)取自布哈河,其中 BHH-1 为现代河漫滩沉积, BHH-2 取自河流阶地砂砾石层.

(5)湖相沉积: ELJ-1 和 ELJ-2(36°39'N, 100°26'E)取自二郎剑发育的湖相地层,其中 ELJ-1 取自地层上部, ELJ-2 取自地层下部.

为了减小粒度效应对元素相对含量的影响,选用的样品粒径小于 75 μm .为了减小有机质和碳酸盐对物源示踪信号的干扰,采用质量分数为 30%的 H_2O_2 和 30%的 HCl 去除有机质和碳酸盐.元素含量在中国科学院青海盐湖研究所盐湖化学分析测试部测定,采用的仪器为荷兰帕纳科公司生产的 Axios X 射线荧光光谱仪,各元素分析误差小于 5%.本文种羊场风成沉积剖面样品的氧化物数据引自文献(曾方明等, 2015b).

2 结果与讨论

2.1 元素组成结果

各类样品的元素分析结果和元素比值见表 1.元素 Si、Al、Fe、Ca、K、Mg、Na 以氧化物的形式表示,单位为百分含量(%).各类样品均以 SiO_2 、 Al_2O_3 、 Fe_2O_3 和 K_2O 为主,这 4 种氧化物的含量超过了 80%.元素 Ba、Mn、Nb、P、Sr、Ti、Zr 含量单位为 10^{-6} , Nb 的含量最低.

2.2 青海湖地区晚第四纪黄土的物质来源

中国黄土高原风尘堆积的物质来源问题一直是黄土古环境学研究的热点问题(Sun, 2002; Chen *et al.*, 2007; Xiao *et al.*, 2012; 曾方明等, 2014b).在

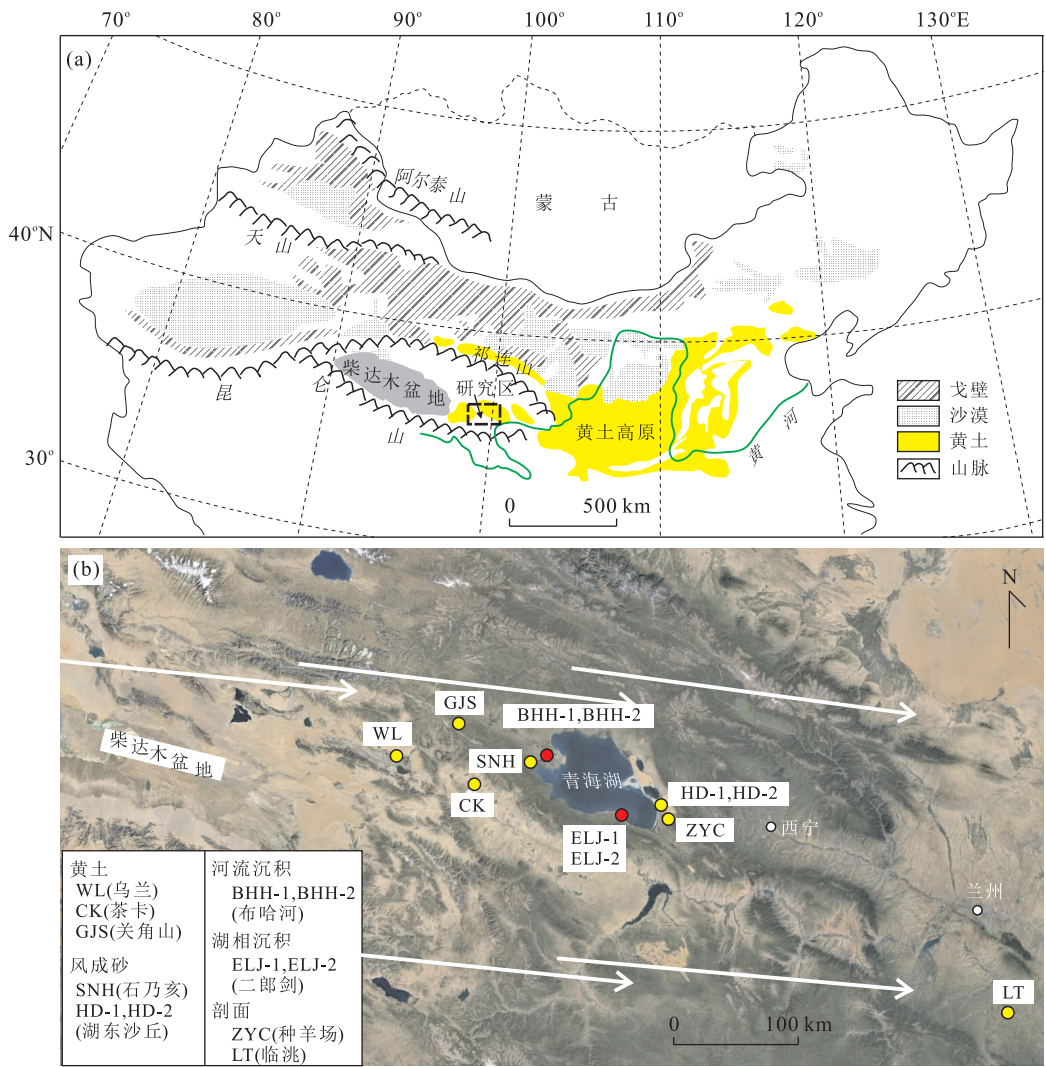


图 1 研究区地理位置(a)和样品分布概图(b)

Fig.1 Location of the study area between the Qaidam basin and Chinese Loess Plateau (a) and spatial sketch of samples (b) 白色箭头为 1979—2008 年春季 200 hPa 西风环流的平均风场,数据来源于 NCEP-2 再分析数据;底图来自 Google Earth

众多的物源示踪手段中,元素地球化学方法可以示踪中国北方和南方风成沉积的物质来源(Sun, 2002; Hao *et al.*, 2010).

在化学风化过程中,钾长石比斜长石更抗风化;在含 K 硅酸盐矿物的风化过程中,部分 K 紧密地结合在伊利石粘土矿物的晶格中,从而在风化剖面中不易迁移(Hao *et al.*, 2010).在所有元素中,Al 和 Ti 在天然水中的溶解度是最小的(Sugitani *et al.*, 1996).研究表明,22 Ma 以来的风尘堆积物(Guo *et al.*, 2002)仍处于脱 Ca 去 Na 的初级风化阶段(Liang *et al.*, 2009),元素 K、Al 在这样的环境下是不易迁移元素, K_2O/Al_2O_3 比值的大小由源区物质的组成决定。 K_2O/Al_2O_3 比值是不受粒级分异影响的反映源区物质组成的代用指标(顾兆炎, 1999),适

用于 22 Ma 以来的风成沉积的源区示踪(郝青振, 2001; 梁美艳, 2009).

在风化和搬运过程中,赋存于碎屑岩中的元素 Zr、Nb 基本上不发生改变,可以指示母岩的特征(Bhatia and Crook, 1986).由于在沉积过程中 Zr、Nb 具有相对较低的移动性,Zr、Nb 被认为是具有物源指示意义的元素(Holland, 1978).元素 Zr 富集在锆石中,而锆石作为一种超稳定的矿物通常可以指示风成沉积的物质来源(Xiao *et al.*, 2012; Xie *et al.*, 2012; Che and Li, 2013).Zr/Ti 和 Zr/Nb 比值已经被成功地用来示踪中国南方下蜀黄土的物质来源(Hao *et al.*, 2010).

青海湖地区种羊场剖面黄土—古土壤样品的 SiO_2 含量为 66.9%~70.5%, Al_2O_3 含量为 12.0%~

表 1 青海湖及周边地区黄土、风成砂、河流和湖相沉积以及临洮黄土元素组成

Table 1 Element compositions of loess, eolian sand, fluvial and lacustrine deposits in the Qinghai Lake and adjacent region and loess deposits at Lintao county

样号	岩性	Ba (10 ⁻⁶)	Mn (10 ⁻⁶)	Nb (10 ⁻⁶)	P (10 ⁻⁶)	Sr (10 ⁻⁶)	Ti (10 ⁻⁶)	Zr (10 ⁻⁶)	SiO ₂ (%)	Al ₂ O ₃ (%)	Fe ₂ O ₃ (%)	CaO (%)	K ₂ O (%)	MgO (%)	Na ₂ O (%)	K ₂ O/Al ₂ O ₃ * (%)	Zr/Ti	Zr/Nb
ZYC8	土壤	453.111	197.006	13.934	324.064	147.355	3 818.946	404.664	70.536	11.738	2.283	0.872	2.627	0.970	1.722	0.242	0.106	29.04
ZYC18	古土壤	466.961	257.434	15.712	370.591	135.879	4 110.997	303.240	68.882	12.480	3.610	0.829	2.652	1.401	1.771	0.230	0.074	19.30
ZYC25	古土壤	493.507	192.925	16.611	330.316	128.783	4 250.000	257.579	69.797	13.304	3.010	0.711	2.967	1.122	1.575	0.241	0.061	15.51
ZYC30	古土壤	515.893	203.684	16.624	267.994	125.633	4 255.133	238.729	70.429	13.126	3.179	0.753	2.910	1.171	1.711	0.240	0.056	14.36
ZYC43	黄土	497.850	242.575	14.881	351.007	132.331	3 897.996	278.210	66.928	12.598	3.558	0.802	2.728	1.432	1.727	0.234	0.071	18.70
ZYC46	黄土	479.095	188.542	14.328	229.557	140.073	3 972.136	300.738	70.544	12.018	2.485	0.835	2.733	1.012	1.790	0.246	0.076	20.99
ZYC57	风成砂	461.648	207.918	14.502	159.270	166.929	3 400.130	425.733	69.447	10.344	2.171	1.037	2.399	1.023	1.950	0.251	0.125	29.36
ZYC60	风成砂	463.094	205.617	14.177	133.082	167.888	3 555.022	426.951	70.041	10.368	2.100	1.046	2.433	0.961	2.000	0.254	0.120	30.12
LT15	古土壤	483.296	198.310	15.366	174.439	145.532	3 834.865	289.525	70.847	12.286	2.706	0.827	2.700	1.110	1.776	0.238	0.075	18.84
LT36	黄土	498.240	170.544	14.261	120.600	153.293	3 755.878	274.598	70.628	11.642	2.215	0.853	2.676	0.955	1.888	0.249	0.073	19.26
CK	黄土	494.261	195.861	14.517	132.495	158.748	3 657.730	390.924	71.541	10.913	2.095	0.937	2.476	0.964	1.870	0.246	0.107	26.93
WL	黄土	532.197	192.101	16.438	170.229	121.740	4 321.781	221.823	70.493	14.105	2.808	0.676	3.243	1.199	1.548	0.249	0.051	13.49
GJS	黄土	481.277	274.787	15.927	270.695	139.605	4 193.302	354.186	69.553	12.343	3.592	0.906	2.580	1.415	1.855	0.226	0.084	22.24
HD-1	风成砂	496.285	154.176	9.572	105.055	178.347	2 660.502	222.920	69.665	9.432	1.404	0.854	2.180	0.814	2.171	0.250	0.084	23.29
HD-2	风成砂	447.814	162.474	12.605	107.225	171.671	2 930.469	403.625	69.770	8.829	1.534	0.826	2.039	0.712	2.189	0.250	0.138	32.02
SNH	风成砂	481.589	212.506	13.072	158.098	166.588	3 153.683	295.566	68.643	11.196	2.514	0.976	2.526	1.229	1.893	0.244	0.094	22.61
BHH-1	河流沉积	466.551	192.329	17.611	133.383	158.391	4 333.907	719.999	70.761	10.931	2.192	0.781	2.350	0.938	1.949	0.233	0.166	40.88
BHH-2	河流沉积	494.661	183.116	15.115	177.608	145.457	3 842.529	349.564	70.276	11.907	2.513	0.657	2.548	1.107	1.828	0.232	0.091	23.13
ELJ-1	湖相沉积	470.430	246.337	14.397	121.810	152.082	3 756.060	484.210	72.615	11.835	2.389	1.060	2.580	1.087	2.087	0.236	0.129	33.63
ELJ-2	湖相沉积	511.306	215.386	17.378	118.480	142.143	4 228.275	395.961	71.432	12.899	2.484	0.885	2.842	1.109	1.826	0.238	0.094	22.79

注: * K₂O/Al₂O₃ 比值为摩尔比(K₂O的分子量为 94.20, Al₂O₃ 的分子量为 101.96)。

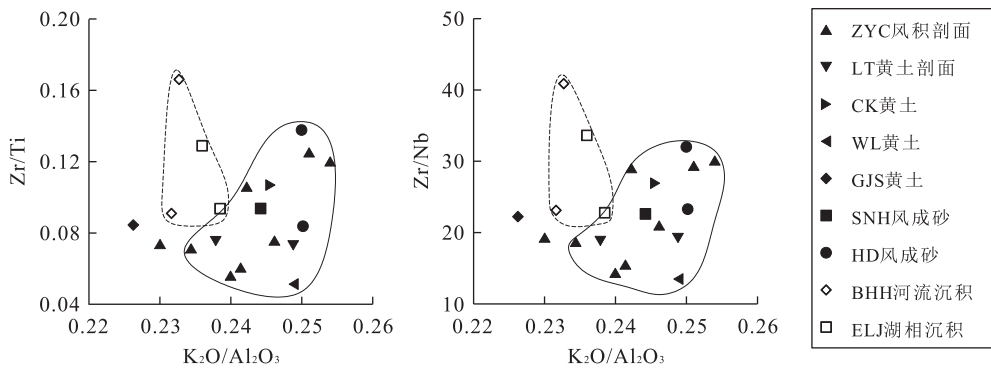


图 2 各类样品的 K_2O/Al_2O_3 -Zr/Ti 和 K_2O/Al_2O_3 -Zr/Nb 比值

Fig.2 Plots of element ratios K_2O/Al_2O_3 vs. Zr/Ti, and K_2O/Al_2O_3 vs. Zr/Nb of different samples

K_2O/Al_2O_3 为摩尔比值

13.3%, K_2O 含量为 2.65%~2.97%; 临洮晚第四纪黄土-古土壤样品的 SiO_2 含量为 70.6%~70.8%, Al_2O_3 含量为 11.6%~12.3%, K_2O 含量为 2.68%~2.70% (表 1)。二者的物质成分非常相似。

在 K_2O/Al_2O_3 -Zr/Ti 和 K_2O/Al_2O_3 -Zr/Nb 图 2 中, 青海湖地区种羊场剖面的晚第四纪风成沉积与青海湖区的风成砂(湖东沙丘和石乃亥风成砂)以及乌兰、茶卡的黄土相似, 而且种羊场剖面的风成沉积与黄土高原西部临洮的黄土也相似, 从而表明青海湖地区的晚第四纪黄土、风成砂与乌兰、茶卡的黄土甚至临洮的黄土具有相同的源区。关角山黄土(样品号 GJS)可能受到了边坡作用的影响, 本地物质输入较多, 从而与乌兰、茶卡的黄土的元素比值不太一致。

青海湖地区的布哈河河流沉积和二郎剑的湖相沉积基本上为流域内河水作用搬运而来的物质, 代表了青海湖区的近源沉积。 K_2O/Al_2O_3 -Zr/Ti 和 K_2O/Al_2O_3 -Zr/Nb 分析结果(图 2)显示, 布哈河河流沉积和二郎剑的湖相沉积具有相对较低的 K_2O/Al_2O_3 比值和相对较高的 Zr/Ti、Zr/Nb 比值。青海湖地区种羊场剖面的晚第四纪风成沉积显著区别于青海湖区的河流沉积和湖相沉积, 从而表明青海湖地区的晚第四纪黄土沉积具有和青海湖区近源沉积不同的源区。相对于布哈河河流沉积和二郎剑湖相沉积, 青海湖地区的晚第四纪黄土沉积为远源沉积。

柴达木盆地分布的雅丹地貌以及上部地层的年代学结果表明柴达木盆地具有显著的侵蚀地貌特征(Bowler *et al.*, 1987; Han *et al.*, 2014; Lai *et al.*, 2014; Zeng and Xiang, 2015c)。研究认为 2.8 Ma 以来, 柴达木盆地成百上千米的垂直地层已经被风力侵蚀掉了(Kapp *et al.*, 2011)。然而, 也有研究认为

柴达木盆地 0.1 Ma 以前的地层因为受到了上部盐壳的保护得到了较好的保存, 柴达木盆地的侵蚀并没有那么严重(Han *et al.*, 2014)。上述研究对于 0.1 Ma 以前柴达木盆地的侵蚀存在不同的观点, 但是对于 0.1 Ma 以来柴达木盆地的侵蚀在认识上是一致的。碎屑锆石的 U-Pb 年龄谱表明柴达木盆地是黄土高原黄土(Pullen *et al.*, 2011)和西宁盆地黄土(Che and Li, 2013)的重要源区。Sr-Nd-Pb 同位素结果表明西宁盆地的风尘堆积很可能来自柴达木盆地(Zeng *et al.*, 2015a)。

柴达木盆地受西风环流控制, 近地面风向为西风(Sun, 2002; Kapp *et al.*, 2011), 处于青海湖地区的上风向地带, 从理论上来说为青海湖地区的黄土提供物源补给是合理的。因此, 通过上述分析, 我们认为青海湖地区的晚第四纪黄土与柴达木盆地关系最为密切, 很可能柴达木盆地就是它的源区。然而这种初步推论尚需要下一步在柴达木盆地采集样品开展工作来进一步证实。而且黄土作为一种上地壳混合均一的物质(Taylor *et al.*, 1983), 其元素所携带的源区信号可能在物质循环混合过程中被均一掉。因此, 本次研究通过元素地球化学手段得出的结论需要其他物源示踪手段(如碎屑锆石、硅酸盐矿物的 Nd-Pb 同位素)来进一步证实。

3 结论

(1) K_2O/Al_2O_3 和 Zr/Ti、Zr/Nb 比值结果显示青海湖地区晚第四纪黄土为远源沉积, 显著区别于青海湖地区布哈河和二郎剑所代表的流域内近源物质。

(2) 柴达木盆地可能是青海湖地区晚第四纪黄土的主要源区。

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