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# 兴蒙造山带的基底属性与构造演化过程

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摘要:为了解兴蒙造山带基底属性和多个构造体系演化与叠加历史,系统总结了近年来在基础地质研究中取得的新成果,并 利用这些成果讨论了兴蒙造山带的基底属性与演化历史,兴蒙造山带是指我国东北地区古生代构造作用影响的地区,这些地 区也遭受了中生代构造作用的叠加与改造.兴蒙造山带主要由微陆块和其间的造山带组成.虽然传统上认为属于前寒武纪结 晶基底的地质体主要已解体为古生代和早中生代,但随着新太古代和古元古代地质体的相继发现,以及新生代玄武岩中幔源 古元古代橄榄岩包体的发现,可以判定兴蒙造山带内微陆块应具有古老的前寒武纪基底,并且壳幔是耦合的.微陆块内部地壳 增生以垂向增生为主,且主要发生在新元古代和中元古代,以及次要的新太古代和古生代.相反,陆块间造山带或岛弧地体的 陆壳则以侧向增生为主,且主要发生在新元古代和市生代.额尔古纳地块与兴安地块的拼合发生在早古生代早期;兴安地块与 松嫩地块的拼合发生在早石炭世晚期;松嫩地块与佳木斯地块的拼合发生在早古生代晚期,中生代早期又经历了裂解与再闭 合的构造演化过程;华北克拉通北缘增生杂岩带与北方微陆块群的最终拼合发生在早古生代晚期,中生代早期又经历了裂解与再闭 合发生在中三叠世,且为剪刀式闭合.晚古生代晚期蒙古一鄂霍茨克大洋板块南向俯冲作用的发生以及早中生代(三叠纪一早 係罗世)的持续南向俯冲,控制了大兴安岭—冀北一辽西地区的岩浆活动,蒙古一鄂霍茨克大洋的闭合发生在中保罗世,晚侏 罗世一早白垩世主要表现为成合后的伸展环境.古太平洋板块中生代的俯冲起始时间为早侏罗世,晚侏罗世一早白垩世早期 东北亚陆缘主要表现为走滑的构造属性和陆缘地体从低纬度到高纬度的构造就位过程,早白垩世晚期一古近纪岩浆作用的 向东收缩揭示了古太平洋板块的持续俯冲和俯冲板片的后撤过程,古近纪晚期日本海的打开标志着东北亚陆缘从活动陆缘 已经转变为沟一弧一盆体系,并且标志着东亚大地幔楔的形成.

关键词:兴蒙造山带;微陆块;地壳增生与再造;古亚洲洋构造体系;蒙古一鄂霍茨克构造体系;环太平洋构造体系;构造地质. 中图分类号:P581 文章编号:1000-2383(2019)05-1620-27 收稿日期:2019-01-02

## Basement Nature and Tectonic Evolution of the Xing'an-Mongolian Orogenic Belt

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Abstract: This paper summarizes recent achievements in basic geological studies in NE China, with the aim of understanding the basement nature and the evolution and overprinting processes of multiple tectonic regimes within the Xing'an-Mongolian Orogenic Belt (XMOB). In this paper, the XMOB only includes the northeastern China which influenced by the Paleozoic orogenic processes, where the overprinting and modification of the Mesozoic tectonic processes took place. The XMOB mainly consists of microcontinental massifs and several orogenic belts between them. Although the so-called Precambrian basement has been dated as Paleozoic and Mesozoic terranes, the new discoveries of the Neoarchean and Paleoproterozoic terranes, together with the Paleoproterozoic mantle-derived xenoliths hosted in Cenozoic basalts, indicate that the microcontinental massifs in the XMOB have the Precambrian basement, and the mantle-crust is coupling within the microcontinental massifs. The crustal ac-

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cretion within the microcontinental massifs mainly happened in the Neoproterozoic and Mesoproterozoic as well as Neoarchean and Paleozoic by vertical accretion. In contrast, the crustal accretion within intercontinental orogenic belts or island arc terranes mainly took place in the Neoproterozoic and Paleozoic by lateral accretion. The amalgamation of the Erguna and Xing'an massifs happened in the Early stage of Early Paleozoic. The collision between the Xing'an and Songnen massifs took place in the late Early Carboniferous. The amalgamation of the Songnen and Jiamusi massifs occurred in the late stage of Early Paleozoic, the breakup and second amalgamation of these two massifs happened in the Early Mesozoic (the Middle Triassic to Early Jurassic). The final collision between the accretionary belt of northern margin of the North China craton and northern massifs took place during the Late Permian to Middle Triassic. The scissor type closure of the Paleo-Asian Ocean finally happened in the Middle Triassic. The southward subduction of the Mongol-Okhotsk oceanic plate happened during the late stage of Late Paleozoic to Early Jurassic, which controlled the magmatic activities in the Great Xing'an Range and northern Hebei-western Liaoning provinces in this period. The closure of the Mongol-Okhotsk Ocean took place in the Middle Jurassic, and the post-closure extensional environment occurred in the Late Jurassic to Early Cretaceous. The onset of subduction of the Paleo-Pacific plate beneath the Eurasia continent took place in the Early Jurassic. The stike-slip tectonic nature in continental margin of northeastern Asia occurred in the Late Jurassic to early Early Cretaceous, resulting in the tectonic emplacement of accretionary complexes in continental margin from low latitude to high latitude. The eastward shrinking of extents of magmatic activities from the late Early Cretaceous to Paleogene reveals the subduction and subsequent roll-back processes of the Paleo-Pacific plate. The opening of the Japan Sea in the Late Paleogene marks the transformation from active continental margin to trench-arc-basin system and the formation of large mantle wedge in eastern Asia.

**Key words**: Xing'an-Mongolian Orogenic Belt; microcontinental massif; crustal accretion and reworking; the Paleo-Asian Oceanic tectonic regime; the Mongol-Okhotsk tectonic regime; the circum-Pacific tectonic regime; tectonics.

中亚造山带是全球最大的增生型造山带,它的 形成与演化以及后期的构造叠加与演化历史一直是 地学领域研究的热点问题之一.中亚造山带东段(国 内通常称为兴蒙造山带)是全球多个构造体系叠加 与演化的经典地区.Şengör et al.(1993)和 Şengör and Natal'in(1996)将该区划分为阿尔泰和满洲两 个构造拼合体,并认为这些地区都是不断向洋迁移 岛弧的系统.本文所指的兴蒙造山带只包括受到古 生代构造作用影响的东北地区,在构造上包括额尔 古纳地块、兴安地块、松嫩地块、佳木斯地块和兴凯 地块以及地块间的构造带和华北克拉通北缘陆缘增 生杂岩带.中生代期间,这一地区又遭受了蒙古一鄂 霍茨克构造体系和环太平洋构造体系的叠加与改造 (Wu et al., 2011).21 世纪以来,随着锆石原位微区 定年技术的广泛应用,兴蒙造山带中微陆块的基底 属性、地质体的时代构成(Wu et al., 2011)、不同构 造体系的形成与演化及其叠加历史等(Ge et al., 2005; Li, 2006; Meng et al., 2010; Xu et al., 2013; 裴福萍等, 2014; Wang et al., 2014, 2017a; 徐备等,2014; Tang et al., 2015, 2016, 2018)研究 已经取得了突出进展,例如,微陆块中传统上认为的 前寒武纪基底(吉林省地质矿产局,1988;内蒙古自 治区地质矿产局,1991;黑龙江省地质矿产局,1993) 基本上已经解体,除少数新元古代地质体外,主体由 古生代和少量早中生代地质体构成(Wilde et al.,

2000; Pei et al., 2006; Miao et al., 2007; Wang et al., 2012; Wu et al., 2012; Xu et al., 2012; 郝文丽 等,2014; Zhao et al., 2016); 传统上认为的海西期 花岗岩其主体形成于中生代(Wu et al., 2011);兴 蒙造山带显生宙地壳增生已经被绝大多数学者所接 受(Jahn et al., 2000, 2004; Wu et al., 2000; Chen and Arakawa, 2005; Xiao et al., 2009)等.然而,兴 蒙造山带微陆块中是否存在前寒武纪结晶基底? 该 区地壳增生是否只发生在显生宙?不同构造体系演 化过程及其影响的时空范围如何?这些问题一直存 在争论.针对这些问题,本文系统总结了近年来对兴 蒙造山带基础地质研究取得的最新研究成果与新发 现,探讨了兴蒙造山带中微陆块的基底属性、地壳增 生的时间与方式及其再造过程、多构造体系的叠加 与演化历史,进而构建了中亚造山带东段多构造体 系古生代一中生代的演化与叠加历史.

# 1 兴蒙造山带:古生代一中生代基本 构造格局

兴蒙造山带,南有西拉木伦-长春-延吉缝合 线与华北克拉通相邻,北有中生代蒙古-鄂霍茨克 缝合带与北亚克拉通相连,东侧为环太平洋构造体 系(图 1a).按微陆块及构造带的属性,区内主要构造



Fig.1 Simplified geological map of main tectonic subdivisions in the Xing'an-Mongolian Orogenic Belt 据 Wu *et al*.(2007b)改

单元包括:微陆块和显生宙造山带,后者包括显生宙 岛弧(如多宝山岛弧)、增生或碰撞杂岩(如跃进山杂 岩和那丹哈达地体)(Li, 2006;图 1b).微陆块自西 向东包括额尔古纳地块、兴安地块、松嫩地块、佳木 斯地块和兴凯地块;陆块间的造山带包括位于兴安 地块与松嫩地块之间的多宝山古生代(早古生代-晚古生代早期)岛弧带(Li, 2006)和华北克拉通北 缘的古生代陆缘增生杂岩带(Wu et al., 2007b);中 生代地体为位于佳木斯地块东侧的那丹哈达地体 (张兴洲等,2006; Zhou et al., 2014); 在佳木斯地 块与那丹哈达地体之间为跃进山杂岩,它是二叠纪 构造就位的增生杂岩(Sun et al., 2015; Bi et al., 2016, 2017)(图 1b).兴蒙造山带,在古生代(尤其是晚 古生代)-中生代期间叠加有蒙古-鄂霍茨克构造体 系的改造,该大洋板块晚古生代晚期发生的南向俯冲 作用(Tang et al., 2014, 2016:Li et al., 2017a)、大洋 的闭合(李宇等,2015)以及闭合后的伸展作用(Tang et al., 2015)对兴蒙造山带进行了强烈改造.同时,兴 蒙造山带在中生代期间也经历了古太平洋板块西向 俯冲作用的叠加与改造(Xu et al., 2013).

# 2 兴蒙造山带:微陆块中前寒武纪基 底是否存在?

兴蒙造山带主要由多个微陆块和其间的造山带 组成.对于微陆块中前寒武纪基底是否存在一直存 在争论.这里所说的兴蒙造山带基底是指显生宙之 前的基底.传统上认为这些微陆块中均存在前寒武 纪地质体,如额尔古纳地块中的兴华渡口群、佳疙疸 组,兴安地块上的风水沟河群,松嫩地块中的东风山 群、兴东群、张广才岭群和一面坡群,佳木斯地块中 的麻山群和马家街群以及兴凯地块中的 Nakhimovka组、伊曼群等(吉林省地质矿产局,1988;内蒙古 自治区地质矿产局,1991;黑龙江省地质矿产局, 1993).自从 20 世纪 90 年代中期以来,锆石原位微 区定年技术的广泛应用,对这些传统上认为的前寒 武纪地质体进行了系统定年,结果表明,这些所谓的 前寒武纪地质体,除少数形成于新元古代外,主体形 成于古生代和早中生代(那福超等,2014; Luan et al., 2017b; Wang et al., 2017c),比如,额尔古纳地 块中原定为古元古代的兴华渡口群,主体形成于古

生代,并有少量新元古代地质体(Miao et al., 2007; 葛梦春等,2011; Zhou et al., 2011a; Wu et al., 2012;Ge et al., 2015; 邵军等, 2015);兴安地块中 原定为新元古代的风水沟河群,主体形成于晚古生 代一早中生代(Xu et al., 2012);松嫩地块中原定 为古元古代的东风山群,除少量新元古代地质体外, 主体形成于古生代(Wang et al., 2006, 2014),而 原定为新元古代的张广才岭群、黄松群、一面坡群主 要形成于古生代-早中生代(Wang et al., 2012;郝 文丽等,2014;郝文丽,2018);佳木斯地块中原定为 太古代的麻山群形成于古生代(Wilde et al., 2000; 吴福元等,2001);兴凯地块西缘原定为古元古代的 麻山群,实际形成时代为早古生代(Zhou et al., 2010).从上述研究结果可以看出,传统上认为的前 寒武纪地质体主体均已解体,那么,兴蒙造山带这些 微陆块中是否存在前寒武纪基底?

近年来,随着研究的深入,在不同的微陆块中均 发现了新元古代侵入体,这些侵入体均显示片麻状 构造,变形强烈,如额尔古纳地块中已经发现多期新 元古代(时代为 737~927 Ma)花岗岩和少量的辉长 岩(约 790 Ma)(Wu et al., 2011, 2012; Zhou et al., 2011b; Gou et al., 2013; Tang et al., 2013; Zhao et al., 2016; 赵硕等, 2016a, 2016b);在兴安 地块的北部已鉴别出新元古代侵入体(Zhou et al., 2011b);在松嫩地块东部的小兴安岭-张广才岭也 相继发现了形成于 726~929 Ma 的新元古代花岗 质侵入体(权京玉等, 2013; Wang et al., 2014; Luan et al., 2017a, 2019);在佳木斯地块上不仅发现 了新元古代花岗质侵入体(Yang et al., 2017, 2018),同时也鉴别出新元古代形成的沉积建造 (Luan et al., 2017a);在兴凯地块同样发现了年龄 约为 750 Ma 的花岗质侵入体 (Khanchuk et al., 2010)(图 2).兴蒙造山带古生代沉积建造中碎屑锆 石 U-Pb 定年结果,对这些新元古代地质体的存在 也给予了很好的佐证,如佳木斯地块和松嫩地块东 缘晚古生代早期沉积建造中发现有大量的新元古代 碎屑锆石(Meng et al., 2010),在大兴安岭北段奥 陶纪一泥盆纪沉积建造中碎屑锆石的定年结果同样 证明了大量新元古代岩浆锆石的存在(Han et al., 2011, 2017; 王利民, 2015; Yang et al., 2018).

在上述沉积建造碎屑锆石 U-Pb 年代学研究中,除了古生代和新元古代碎屑锆石外,还存在一部分古元古代碎屑锆石(Meng et al., 2010; Han et al., 2011, 2012; Wang et al., 2012, 2014),这就给



Fig.2 Distribution map of Precambrian terranes within microcontinental massifs in the Xing'an-Mongolian Orogenic Belt

我们提出了这样一个问题,在兴蒙造山带这些微陆 块中是否还存在古元古代地质体? 古生代和新元古 代花岗质岩石中锆石的 Hf 同位素模式年龄也暗示 该区应存在古元古代甚至更为古老的陆壳物质 (Tang et al., 2013; Wang et al., 2016, 2017c; Luan et al., 2017b; Yang et al., 2017).虽然 Pei et al.(2007)报道了松辽盆地南部基底钻孔中存在 1 873±13 Ma 和 1 808±21 Ma 的变质岩,但由于 该钻孔位置靠近华北克拉通,因此对其成因还存在 争论,即它是代表松嫩地块的基底?还是来自南部 华北克拉通的推覆体?由此还不能作为判定松嫩地 块存在古元古代地质体的直接证据.然而,随着研究 的深入,在不同的微陆块中相继发现了出露于地表 的古元古代和新太古代地质体,如孙立新等(2013a, 2013b)报道了额尔古纳地块北段存在1837±5 Ma 和1 741±30 Ma的花岗质片麻岩,张超等(2018)报 道了兴安地块东侧 18 亿年的花岗岩,钱程等(2018) 同样在兴安地块东侧发现了 2 579±15 Ma的花岗 岩,本研究组在松嫩地块东缘的张广才岭南段也发 现了 18 亿年的地质体(Luan et al., 2019),这些古 元古代和新太古代地质体的发现,与古生代和新元 古代沉积建造中古元古代和新太古代碎屑锆石的存 在以及古生代和新元古代花岗岩中锆石的古元古代 和新太古代 Hf 同位素模式年龄一起,表明兴蒙造 山带这些微陆块中存在古元古代和新太古代基底. 然而,由于新元古代一古生代一中生代岩浆作用的 多期改造及其花岗岩形成过程中的多期再造,造成 目前只有极少数保存在晚期的花岗岩中.兴蒙造山 带微陆块中存在古元古代基底物质也得到了近年来 岩石圈地幔研究的证实,如 Zhang et al.(2011b)报道 了兴安地块和松嫩地块新生代玄武岩中存在古元古 代Re亏损模式年龄的地幔橄榄岩捕虏体,Wang et al.(2015b)和 Guo et al.(2017a)分别报道了兴凯地块 新生代玄武岩中存在古元古代 Re 亏损模式年龄的地

幔橄榄岩捕虏体,这些结果都暗示在这些微陆块的岩石圈地幔中存在古老岩石圈地幔物质,并且地壳和岩石圈地幔在时代上是耦合的,它们共同揭示兴蒙造山带的这些微陆块曾经具有古老的陆壳和古老的岩石圈地幔物质,这些古老的地壳与岩石圈地幔物质已经被后期的新元古代、古生代和中生代构造岩浆作用所再造、改造或被新增生的地幔或地壳物质所置换(Guo et al., 2017a;Sun et al., 2017).

# 3 兴蒙造山带:地壳的增生与再 造过程

中亚造山带是全球最大的增生型造山带之一,这 一认识已经得到了共识(Sengör et al., 1993; Windley et al., 2007; Xiao et al., 2009),并且基于其中大量岛 弧组合和蛇绿岩的存在以及花岗岩全岩 Sm-Nd 同位 素资料得出,中亚造山带是地球上显生宙地壳增生的 主要场所(Sengör et al., 1993; Jahn et al., 2000).上 述研究主要集中在岛弧地体组合和其中的花岗质岩 石的 Sm-Nd 同位素组成方面(Jahn et al., 2000; Wu et al., 2000; Hong et al., 2004). 然而, 中亚造山带, 尤 其是中亚造山带东段——兴蒙造山带,其中不仅有岛 弧地体组合和陆缘增生杂岩带的存在,更为重要的是 其主体由多个微陆块组成,仅用少量岛弧地体中花岗 质岩石 Sm-Nd 同位素组成得到的认识并不能代表整 个造山带的地壳增生历史,也就是说以往的认识过高 地估算了中亚造山带显生宙的地壳增生量(Kröner et al., 2014, 2017; Sun et al., 2017).

近年来,对兴蒙造山带微陆块中不同时代花岗 质岩石进行了广泛的锆石 Hf 同位素组成分析,以 额尔古纳地块为例,该地块中发育古元古代、新元古 代、古生代和中生代花岗质岩浆作用(图 3),不同时 代花岗岩中锆石 Hf 同位素的二阶段模式年龄统计 结果表明,该微陆块中地壳增生发生的主要时期是





Fig.3 Probability plot of granitoid magmatic event ages within the Erguna massif



图 4 额尔古纳地块花岗岩锆石 Hf 同位素二阶段模式年龄频谱与地壳增生

Fig.4 Probability plot of zircon  $T_{DM2}$  (Hf) ages from granitoids and crustal accretion within the Erguna massif



图 5 额尔古纳地块花岗岩年龄与锆石 Hf 同位素二阶段模式年龄变异图 Fig.5 Plot of zircon T<sub>DM2</sub>(Hf) ages against granitoid ages within the Erguna massif



图 6 多宝山岛弧地体古生代-中生代花岗岩年龄与锆石 Hf 同位素组成变异图

Fig.6 Plots of zircon Hf compositions ( $T_{DM2}$  (Hf) ages and  $\epsilon_{Hf}$  values) against the ages of Paleozoic-Mesozoic granitoids from Duobaoshan island arc terrane

中元古代和新元古代,新太古代次之,古生代更微弱 (图 4).然而,花岗质岩浆作用时代与锆石 Hf 两阶 段模式年龄图解表明,随着花岗岩形成时代变新,锆 石 Hf 同位素模式年龄逐渐变年轻,这很好地揭示 了不同时代花岗质岩石的形成是不同时代源岩再造 的结果(Sun *et al.*, 2017)(图 5).对兴凯地块古生 代一中生代花岗岩中锆石 Hf 同位素组成的研究结 果同样给出了类似的认识:地壳增生主要发生在新 元古代和中元古代,其次为古元古代,而不同时代花 岗岩的形成同样显示出不同时代源岩分别再造的产 物(Zhang et al., 2018).

在兴蒙造山带中,与微陆块相比,岛弧地体组合 出露的面积很少,主要有多宝山古生代岛弧地体和 华北克拉通北缘古生代陆缘增生杂岩带.对多宝山 岛弧地体中发育的古生代和中生代花岗岩锆石 Hf 同位素统计结果表明,这些花岗质岩石所揭示的地 壳增生事件主要发生在新元古代和古生代,新元古 代和古生代早期岛弧地体物质的再造为古生代和中 生代花岗质岩浆的产生提供了物源(图 6).

综上所述,可以看出兴蒙造山带的地壳增生可

以划分成2个构造阶段,即微陆块形成演化阶段的 陆壳增生和古生代造山带形成演化阶段的陆壳增 生,其中微陆块的地壳增生主要发生在新元古代和 中元古代,新太古代和古元古代次之,显生宙地壳增 生量很少,地壳的增生方式可能以垂向增生为主.相 反,古生代造山带(或岛弧地体)的地壳增生主要发 生在古生代和新元古代,地壳的增生方式以侧向 增生为主.

4 微陆块间的拼合历史

#### 4.1 额尔古纳地块与兴安地块的拼合

4.1.1 缝合线的位置 额尔古纳地块与兴安地块 之间缝合线的位置一直是一个争论的问题.传统上 将德尔布干断裂作为二者间的缝合线(黄汲清等, 1977;黑龙江省地质矿产局,1993;任纪舜等,1999; Wu et al., 2003).然而,近年来的工作表明德尔布 干断裂是一个晚中生代(115~130 Ma)形成的伸展 断裂(郑常青等,2009).此外,德尔布干断裂作为一 个地块间的缝合带也缺乏地球物理资料的支持(孙 晓猛等,2011).因此,德尔布干断裂不可能作为陆块 间的缝合带.相反,新林蛇绿岩的发现(李瑞山, 1991)、塔河地区早古生代同碰撞型花岗岩的发现 (Ge et al.,2005;Wu et al.,2011)和头道桥早古生 代高压蓝片岩的发现(Zhou et al.,2015)均暗示新 林一喜桂图-头道桥构造带应是额尔古纳地块与兴 安地块之间的构造缝合线(图 7).

**4.1.2 块体的拼合时间**额尔古纳地块与兴安地 块的拼合时间可以从蓝片岩的变质时间、缝合带中 同碰撞岩浆作用发生的时间以及 2 个陆块上岩浆作 用的时间与性质等方面得到回答.(1)对头道桥蓝片 岩的锆石 U-Pb 年代学研究表明,蓝片岩相变质作 用发生的时间为 510~490 Ma(Miao et al., 2015; Zhou et al., 2015);(2)塔河地区同碰撞型花岗岩的 形成时代为 500 Ma(Ge et al., 2005);(3)额尔古纳 地块与兴安地块均发育 480 Ma 和 460 Ma 的岩浆



图 7 额尔古纳地块、兴安地块和松嫩地块间的缝合线位置

Fig.7 Locations of the suture zones between the Erguna and Xing'an massifs as well as the Xing'an and Songnen massifs 据 Li et al., 2017b;①德尔布干断裂;②喜桂图一塔源缝合带;③黑河一嫩江一贺根山断裂;④黑河一嫩江缝合带 作用,并且显示双峰式火成岩组合,暗示区域伸展环境的存在(Ge et al., 2005; Zhao et al., 2014); (4) 在额尔古纳地块北段兴华渡口群发生的麻粒岩相变质作用约发生在 500 Ma(Zhou et al., 2011a, 2011b).综合上述特征,我们可以得出额尔古纳地块与兴安地块的拼合时间应是早古生代早期.

#### 4.2 兴安地块与松嫩地块的拼合

4.2.1 缝合线的位置 兴安地块与松嫩地块之间 缝合线的位置一直存在争论,争论的焦点集中在缝 合线从扎兰屯向南如何延伸.基于兴安地块北段早 古生代多宝山岛弧的存在以及地块南侧贺根山蛇绿 岩的发现,多数学者认为黑河一嫩江一贺根山构造 带作为兴安地块与松嫩地块间的缝合线(图 7;Wu et al., 2011; Liu et al., 2017). 然而, 徐备等(2014) 和 Xu et al.(2015)根据晚古生代早期沉积建造组合 和前寒武纪地质体的出露位置,将黑河一嫩江一乌 兰浩特-锡林浩特南-艾力格庙构造带作为2个地 块间的缝合线.Li et al.(2017a)根据兴安地块东缘 和东南缘早古生代弧型火成岩的发现,这包括北部 的多宝山弧型火成岩及其斑岩型铜一钼矿床(Li, 2006; Ge et al., 2007; Zeng et al., 2014; Wu et al., 2015, 2018)、扎兰屯南部蘑菇气地区早古生代钙碱 性火山岩(Guo et al., 2009:Li et al., 2017b)和南 部苏尼特左旗一锡林浩特地区早古生代弧型火成岩 (Chen et al., 2000;石玉若等, 2005; Jian et al., 2008),将2个地块间的缝合线确定在该早古生代火 成岩带的东侧,即黑河一嫩江一乌兰浩特一锡林浩 特南一线,该构造线也是重要的生物群界线,其北以 分布图瓦贝(Tuvaella)动物群为特色,见于多宝山、 额尔古纳西、伊尔施和东乌珠穆沁旗等地(内蒙古自 治区地质矿产局,1991;徐备等,2014).此外,结合对 贺根山蛇绿岩的最新定年结果显示其形成时代应在 350 Ma 左右(Zhang et al., 2015),这表明贺根山洋 盆打开的时间应是石炭纪早期,它是陆内裂开的一 个洋盆.综上所述,本文认为黑河一嫩江一乌兰浩 特一锡林浩特南一线应是兴安地块与松嫩地块之间 的缝合线(图 7).

**4.2.2 块体的拼合时间** 兴安地块与松嫩地块之间的拼合时间具有多种观点,包括晚志留世一泥盆纪(Şengör and Natal'in, 1996)、晚泥盆世一早石炭世(邵济安,1991; Hong *et al.*, 1995)、早石炭世晚期(赵芝等,2010)、二叠纪之前(Shi *et al.*, 2004; Sun *et al.*, 2001;童英等, 2010)和三叠纪(Chen *et al.*, 2000; Miao *et al.*, 2004).目前多数学者支持二

者的拼合发生在晚古生代.Li et al.(2013)基于黑 河一嫩江缝合带两侧晚石炭世早期"S型"火成岩的 发现,并结合小兴安岭西北侧从早石炭世海相沉积 到晚石炭世陆相沉积的转变(黑龙江省地质矿产局, 1993),认为兴安地块与松嫩地块之间的拼合发生在 早石炭世晚期(约 320 Ma),这与研究区普遍缺失早 石炭世晚期(Serpukhovian 期)一晚石炭世早期 (Bashkirian and Moscovian 期)沉积是相吻合的(黑 龙江省地质矿产局,1993;内蒙古自治区地质矿 产局,1991).

#### 4.3 松嫩地块与佳木斯地块的拼合

4.3.1 缝合线的位置 牡丹江断裂,与沿该断裂分 布的黑龙江杂岩(包括磨刀石、依兰和萝北3个出露 区)一起,通常被认为是松嫩地块与佳木斯地块间的 缝合线,这已经被绝大多数学者所认可(张兴洲, 1992;黑龙江省地质矿产局,1993;李锦轶等,1999; 赵焕利等,2012).但是对其南延却有不同认识,多数 学者将该缝合带南延到磨刀石蓝片岩出露地,而周 建波等(2013)将该带向南与华北克拉通北缘吉中地 区出露的红帘石片岩带相连(为长春-延吉缝合带 的位置),并统称为吉黑高压变质带.周建波等 (2013)认为吉黑高压变质带是由于松嫩地块与佳木 斯地块的碰撞拼合以及佳木斯地块与华北克拉通的 碰撞拼合所致,其碰撞拼合的时间为晚三叠世一早 侏罗世.本文的认识是:由于华北克拉通北缘长春一 延吉缝合带的变质时间(约 250 Ma; Wu et al., 2007b) 明显早于黑龙江杂岩的构造就位时间 (175~186 Ma; Wu et al., 2007a; Zhou et al., 2009; Aouizerat et al., 2018; Dong et al., 2018), 同时结合华北克拉通北缘近东西向晚三叠世碱性火 成岩和双峰式火成岩带的存在(Xu et al., 2013; Tang et al., 2018),可以判定华北克拉通北缘吉 中一延吉的中高压变质带应是西拉木伦一长春一延 吉缝合带的一部分,应与古亚洲洋的最终闭合有关 (Liu et al., 2017),而与松嫩地块和佳木斯地块的 拼合无关,沿嘉荫一牡丹江断裂分布的黑龙江杂岩 应代表了该缝合线的位置.

**4.3.2 块体的拼合时间** 松嫩地块与佳木斯地块间的拼合时间一直是个争论的问题,目前主要有 2 种看法,一是认为 2 个陆块之间的拼合发生在早古生代晚期,这主要是基于松嫩地块东缘早古生代火成岩的研究(李锦轶等,1999;Wang *et al.*, 2012)以及早期对黑龙江杂岩的研究(张兴洲,1992);另一种观点认为 2 个陆块间的拼合发生在早侏罗世,这种

观点的提出主要是基于对黑龙江杂岩变形时代的研究得出的(Wu et al., 2007a; Zhou et al., 2009; Aouizerat et al., 2018; Dong et al., 2018).上述认 识的差异主要是由于对黑龙江杂岩形成时代认识的 不同(见 4.3.3).许文良等(2012)基于对牡丹江断裂 两侧岩浆事件的对比研究,结合佳木斯地块中三叠 纪岩浆事件的缺乏以及区域二叠纪沉积建造组合属 性,认为松嫩地块与佳木斯地块之间的块体拼合曾 经发生过 2 次,早期拼合发生在早古生代末期(约 425 Ma),晚期拼合发生在早侏罗世.由于后期构造 作用的改造,早期拼合位置较难恢复,而两地块再次 裂开的时间可以通过黑龙江杂岩中沉积建造形成的 最早时间给出限定,约 235~180 Ma 沉积作用 发生(见 4.3.3).

4.3.3 牡丹江洋——中生代早期的一个短命洋 牡丹江洋的提出主要是基于 2 个陆块间出露的黑龙 江杂岩,尤其是依兰地区黑龙江杂岩中 MORB 型玄 武岩和纯橄岩/蛇纹岩的存在(Wu et al., 2007a; Zhou et al., 2009).然而,由于对黑龙江杂岩中构造 块体形成时代的认识不一和牡丹江断裂两侧晚古生 代弧型火成岩的存在,导致了牡丹江洋形成时代的 2 种不同认识,一种观点认为古生代期间,牡丹江洋 一直存在,并且双向俯冲于松嫩地块和佳木斯地块 之下(Dong et al., 2017, 2018),另一种观点认为牡 丹江洋形成于早中生代,是在原来已经拼合的松嫩 地块与佳木斯地块之上重新裂开的(许文良等, 2012).最近,孙晨阳等(2018)系统总结了黑龙江杂 岩中沉积单元碎屑锆石的构成,发现出露于磨刀石、 依兰和萝北3处黑龙江杂岩中的沉积单元最年轻的 碎屑锆石年龄为235~180 Ma,而黑龙江杂岩中白 云母的 Ar-Ar 年龄和变辉长岩中金红石的 U-Pb 年 龄为171~177 Ma(Aouizerat *et al.*, 2018; Dong *et al.*, 2018),这说明黑龙江杂岩中沉积岩的形成时 代应为235~177 Ma,即牡丹江洋的形成时代应是 中三叠世-早侏罗世,由此可以看出位于2个陆块 间的牡丹江洋是一个早中生代的短命洋(图 8).

4.4 兴凯地块与佳木斯地块和松嫩地块的关系

4.4.1 兴凯地块与佳木斯地块的关系 在中亚造 山带东段,传统上将布列亚地块、佳木斯地块和兴凯 地块作为一个整体称为布列亚一佳木斯一兴凯地 块,并且认为它们具有类似的构造属性(Zhou et al., 2010;Sorokin et al., 2017).Zhou et al.(2010) 对兴凯地块西北缘虎头地区出露的麻粒岩相变质岩 研究表明,该区存在泛非期的变质作用,并以此认为 兴凯地块与佳木斯地块具有类似的属性.然而,王枫 等(2016)和Xu et al.(2018)对比了佳木斯地块和兴 凯地块中发育的古生代和中生代岩浆作用,结果发 现兴凯地块中普遍发育早古生代晚期和早中生代岩



图 8 牡丹江洋的形成与黑龙江杂岩构造就位模式

Fig.8 A model for the formation of the Mudanjiang Ocean and tectonic emplacement of the Heilongjiang complex 据孙晨阳等(2018);SNM.松嫩地块;JM.佳木斯地块





据 Xu et al.(2018)

浆事件,而佳木斯地块缺乏这些岩浆事件,为此判定 2个陆块可能不具有类似的构造演化历史,并且认 为兴凯地块西北缘虎头地区出露的麻粒岩相变质岩 可能属于佳木斯地块上的麻山群在敦化一密山断裂 左行平移过程中切割过去的佳木斯地块物质(图9). 上述认识也得到了2个陆块上具有不同的早古生代 沉积建造组合的支持(吉林省地质矿产局,1988;邵 济安和唐克东,1995).

**4.4.2 兴凯地块与松嫩地块的关系** 正如 4.4.1 节 所说,兴凯地块与佳木斯地块至少在早古生代晚期 和中生代早期不具有类似的岩浆作用历史.相反,古 生代和中生代岩浆事件对比表明,松嫩地块与兴凯 地块具有类似的岩浆作用历史,暗示二者具有相似的构造属性(图 9; 王枫等, 2016; Xu et al., 2018). 上述认识也得到了 2 个陆块上均发育约 750 Ma 的 岩浆事件的证实(Khanchuk et al., 2010; Wang et al., 2015b).王枫等(2016)系统总结对比了松嫩地 块与兴凯地块古生代和早中生代岩浆作用,认为敦 化-密山断裂至少发生过 2 次走滑事件,一次发生 在中一晚二叠世-早三叠世,另一次发生在晚侏罗 世-早白垩世早期,由于上述走滑事件的发生,造成 了目前微陆块的分布(图 9b).

## 5 古亚洲洋最终闭合的时间与方式

古亚洲洋最终闭合的位置通常是指沿西拉木 伦-长春-延吉缝合线发生的古亚洲洋闭合.对其 闭合的时间一直存在争论,主要观点有:(1)基于东 北地区泥盆纪稳定陆缘沉积建造的形成,认为古亚 洲洋的最终闭合在早古生代晚期已经完成(徐备等, 2014;Xu et al., 2015);(2)根据二叠纪磨拉石沉积 建造组合的形成或变质作用发生的时间,认为古亚 洲洋的最终闭合发生在中一晚二叠世(吉林省地质 矿产局,1988;Shi, 2006;Wu et al., 2007b;李锦轶 等,2009);(3)根据内蒙古东南部和辽北一吉中地区



图 10 华北克拉通北缘早中生代地层物源变化与古亚洲洋最终闭合过程

Fig.10 Provenance variation of the Early Mesozoic strata in the northern margin of the North China craton and the final closure process of the Paleo-Asian Ocean

碰撞型花岗岩的形成时间,认为古亚洲洋的最终闭 合发生在晚二叠世一中三叠世(Cao et al., 2013; Wang et al., 2015d)或早一中三叠世(孙德有等, 2004).上述争论之所以存在,追其根本原因就是不 同学者研究对象的不同或研究地区的不同.根据沿 西拉木伦一长春一延吉缝合线晚古生代晚期一早中 生代同碰撞型花岗岩的研究表明,自西向东同碰撞 型花岗岩形成时代具有逐渐变新的趋势(Wang et al., 2015d),二叠纪沉积建造的空间变化也反映了 类似的特征——自西向东二叠纪海相地层具有逐渐 变年轻的趋势(吉林省地质矿产局,1988),这表明古 亚洲洋最终闭合的时间具有自西向东逐渐变年轻的 趋势,东部最终闭合的时间为中三叠世(Wang et al., 2015d; Wang et al., 2018).上述认识也得到了 华北克拉通北缘早中生代地层和物源分析的支 持-----在华北克拉通北缘燕辽地区缺失早三叠世地 层(Meng et al.,待发表),而在华北克拉通东北缘 缺失中三叠世地层,同时早三叠世地层的物源均来 自华北克拉通内部,而晚三叠世地层中已经存在来 自中亚造山带的物源,这说明晚三叠世期间古亚洲 洋已经最终闭合(Wang et al., 2018).结合华北克 拉通北缘早中生代地层的缺失及其物源变化和同碰 撞花岗岩的空间变异,本文认为古亚洲洋最终闭合 的时间发生在中三叠世,并且表现为自西向东逐渐 闭合的剪刀式闭合方式(图 10).

# 6 蒙古一鄂霍茨克构造体系的叠加与 演化:大洋板块的南向俯冲与演化历史

蒙古一鄂霍茨克缝合带是蒙古一鄂霍茨克洋闭 合后的产物,主要分布在东经96°~130°,北纬46°~ 58°的俄罗斯和蒙古境内,西起蒙古中部的杭爱山 脉,东至鄂霍茨克海的乌达海湾,总体呈北东一南西 走向,长约为3000 km,宽约为300 km,北部为西伯 利亚克拉通及其增生边缘,南部为中朝一蒙古板块 及其以北的造山带与地块镶嵌构造区,东部为太平 洋板块(黄始琪等,2014).它是东亚北部一条具有较 长地质历史的造山带,并在东亚大陆形成演化的历 史过程中占有极其重要的位置(李锦轶等,2009).蒙 古一鄂霍茨克洋普遍被认为是古太平洋的巨型海湾 (Zonenshain *et al.*, 1990;Gordienko, 1994;Tang *et al.*, 1995;Zorin, 1999;Parfenov *et al.*, 2001; Shi, 2006),在晚古生代一早中生代期间分隔西伯

利亚克拉通和中朝-蒙古板块.目前对蒙古-鄂霍 茨克大洋形成的时间(Tomurtogoo et al., 2005; Donskaya et al., 2013)、俯冲演化历史(Sengör et al., 1993; Zorin, 1999; Parfenov et al., 2003; Bussien et al., 2011; Donskaya et al., 2013) 以及大洋 闭合的时间(Zonenshain et al., 1990;Enkin et al., 1992;Zorin, 1999;Parfenov et al., 2001)还有诸多 争论,但是该大洋板块北向俯冲于西伯利亚板块之 下(Zorin, 1999)和自西向东剪刀式闭合方式 (Kravchinsky et al., 2002; Metelkin et al., 2007) 得到了绝大多数地质学家认可.蒙古-鄂霍茨克大 洋板块是否存在南向俯冲作用?如果存在,它的演 化历史及其影响的时空范围如何? 这些问题至今没 有得到很好的解决.鉴于此,本文基于近年来在我国 境内火成岩和区域成矿作用的研究成果,主要讨论 该构造体系对我国东北乃至华北的影响,这对于揭 示区域成矿背景和指导找矿具有重要的现实意义.

# 6.1 蒙古一鄂霍茨克大洋板块南向俯冲作用的起始时间

以往的研究多数集中在蒙古一鄂霍茨克大洋板 块北向俯冲作用过程,这主要是因为蒙古一鄂霍茨 克缝合带的主体位于俄罗斯和蒙古境内.与之相比, 对于蒙古一鄂霍茨克大洋板块南向俯冲过程的研究 较少.Zorin(1999)认为蒙古-鄂霍茨克大洋板块在 早石炭世向南俯冲于中朝-蒙古板块之下.然而,随 着近年来对蒙古国地质研究程度的提高,刘翼飞等 (2010)认为蒙古国阿林诺尔钼矿赋矿岩体的成岩成 矿作用是蒙古-鄂霍茨克大洋板块于中三叠世(约 229 Ma)向其南侧大陆俯冲构造体制下地壳伸展作 用的产物;位于中蒙古地块的晚二叠世-三叠纪 (260~235 Ma) Hangay 岩基以 I 型含角闪石花岗 闪长岩为主(Jahn et al., 2004; Li et al., 2013), 研 究表明该岩基形成于安第斯型大陆边缘弧环境 (Tomurtogoo et al., 2005; Orolmaa et al., 2008). 而对于我国来说,额尔古纳地块与蒙古一鄂霍茨克 缝合带相邻,位于缝合带的东南侧,前人认为该地块 以元古宙和古生代岩浆事件为主,中生代岩浆作用 微弱(内蒙古自治区地质矿产局,1991),因而忽略了 蒙古一鄂霍茨克缝合带对我国东北地区的影响.随 着现代同位素测年技术的发展和精度的提高,在额 尔古纳地块乃至整个东北地区鉴别出大量中生代岩 浆事件(Wu et al., 2011; Xu et al., 2013; Tang et al., 2014, 2015, 2016),认识到蒙古-鄂霍茨克缝 合带对我国东北地区存在重要影响,如太平川斑岩



图 11 蒙古一鄂霍茨克构造体制与环太平洋构造体制早侏罗世一新生代演化模式

Fig.11 The evolutionary model for the Mongol-Okhotsk and circum-Pacific tectonic regimes during the Early Jurassic-Cenozoic

据 Tang et al.(2018)修改; SC.西伯利亚克拉通; XMOB.兴蒙造山带; SAOB.锡霍特一阿林造山带

型铜钼矿床被认为是形成于晚三叠世蒙古一鄂霍茨 克洋向其南侧的额尔古纳地块俯冲形成的陆缘弧环 境(陈志广等,2010);大兴安岭地区早中生代花岗 岩、火山岩形成于与蒙古一鄂霍茨克大洋板块俯冲 有关的活动大陆边缘环境(Wu et al., 2011;佘宏全 等,2012;王伟等,2012;Xu et al., 2013;Wang et al., 2015c; Tang et al., 2016).上述研究已经表 明,蒙古一鄂霍茨克大洋板块存在南向俯冲作用,根 据对额尔古纳地块上晚古生代岩浆作用的研究,认 为南向俯冲作用的起始时间在额尔古纳地块西北侧 至少发生在晚二叠世(Li et al., 2017a).

# 6.2 蒙古一鄂霍茨克大洋板块早中生代(T-J<sub>1</sub>)持续南向俯冲

近年来,随着对我国东北地区中生代岩浆作用 研究的深入,尤其是大量高精度锆石 U-Pb 年代学 资料的获得,使传统上认为东北地区广泛存在的海 西期岩浆作用(吉林省地质矿产局,1988;内蒙古自 治区地质矿产局,1991;黑龙江省地质矿产局,1993) 得以修正,东北地区所谓的海西期花岗岩主体形成 于中生代(Wu et al., 2011).中蒙古地块和额尔古 纳地块上大量早中生代(早三叠世一早侏罗世)钙碱 性火成岩(尤其是早侏罗世钙碱性火山岩组合)和同 期斑岩型矿床的发现以及它们呈南西一北东向的带 状展布特征均揭示了蒙古一鄂霍茨克大洋板块南向 俯冲作用的发生(图 10 和 11a; Tang et al., 2014, 2016; Wang et al., 2015c);兴安地块东缘早中生代 钙碱性火山岩的发现进一步揭示蒙古一鄂霍茨克大 洋板块南向俯冲作用影响的空间范围至少到达松辽 盆地以西地区(Li et al., 2017c).

#### 6.3 蒙古一鄂霍茨克大洋闭合的时间:中侏罗世

对于蒙古一鄂霍茨克洋的剪刀式闭合方式几乎 得到了绝大多数学者的认可.但是,就其最终闭合的 时间还存在较多争论,Zorin(1999)和 Parfenov et al.(2001)认为该洋是在早一中侏罗世闭合;根据古 地磁数据,Kravchinsky et al.(2002)和 Cogné et al. (2005)认为蒙古一鄂霍茨克洋主要是在侏罗纪期间 闭合,东部的闭合时间可持续到晚侏罗世一早白垩 世;Zonenshain et al. (1990)和Sengör and Natal'in (1996)认为蒙古一鄂霍茨克洋西部于三叠纪闭合, 东部于晚侏罗世闭合;而 Enkin et al. (1992)和 Scotese(2001)通过解析古地磁数据,认为蒙古一鄂 霍茨克洋在晚侏罗世并没有闭合,而是于早白垩世 闭合.Guo et al.(2017b)对我国境内漠河盆地和俄 罗斯境内上阿莫尔盆地沉积相的时空变异研究,认 为蒙古一鄂霍茨克洋的最终闭合应发生在晚侏罗世 最晚期一早白垩世早期.本研究组在大兴安岭北段 发现了具有S型花岗岩地球化学属性的中侏罗世 (约168 Ma)白云母二长花岗岩,该区同期花岗岩具 有类似的地球化学属性,暗示该期花岗岩形成于陆 壳加厚背景(赵海滨等,2005;李宇等,2015).结合燕 辽地区海房沟组之下广泛发育的一个区域不整合 (即燕山运动 A 幕)以及自北向南的构造推覆作用 (张岳桥等,2007; Zhang et al., 2011a),本文认为蒙 古一鄂霍茨克大洋闭合发生在中侏罗世(图 11b). 上述认识也得到了俄罗斯学者对蒙古一鄂霍茨克带 研究结果的支持,即中侏罗世地层以区域不整合的 形式覆盖在之前的地层之上,这一地质现象在该带 东西两端均是如此,并且他认为自中侏罗世之后,该 带不存在洋盆,但存在海盆.因此,沿该带仍存在晚 侏罗世一早白垩世的海相沉积(Sorokin, 未发表). 6.4 蒙古一鄂霍茨克大洋闭合后晚中生代的演化

历史——区域伸展

进入到晚侏罗世一早白垩世早期阶段,在大兴 安岭和冀北一辽西地区广泛产出碱性一亚碱性过渡 性质的火山岩和碱性流纹岩,前者以大兴安岭北部 的塔木兰沟组为代表(约162 Ma;孟恩等,2011),南 部以满克头鄂博组为代表(王建国等,2013),在冀 北一辽西地区则以髫髻山组和蓝旗组为代表(165~ 157 Ma;赵越等,2004;胡健民等,2007);后者以大 兴安岭北部的吉祥峰组(约142 Ma)和南部的玛尼 吐组(约142 Ma)碱性流纹岩(王建国等,2013)以及 冀北一辽西地区的张家口组为代表(约135 Ma;张 宏等,2005).这2套(晚侏罗世和早白垩世早期)火 山岩主要发育在断陷盆地中,结合其碱性火山岩组 合,它们共同揭示了区域伸展环境的存在,并分别与 燕山运动 A 幕和 B 幕之后的伸展环境相对应.与这 2 期岩浆事件密切相关的成矿作用主要表现为浅成 低温热液成矿(135~155 Ma; Ouyang et al., 2013).此外,这2期岩浆事件形成的时间具有自北向 南逐渐变新的趋势(Zhang et al., 2008, 2010),暗示 这2期岩浆事件的形成应与蒙古一鄂霍茨克构造体 系的演化有关,我国松辽盆地以东地区以及韩国、日 本均缺失这些岩浆事件进一步证实它们的形成与蒙 古一鄂霍茨克构造体系演化有关,而与环太平洋体系 无关(图 11c; Xu et al., 2013; Tang et al., 2018).

# 7 环太平洋构造体系的叠加与演化: 中生代俯冲起始时间与中一新生代俯 冲演化历史

## 7.1 古太平洋板块中生代在欧亚大陆下俯冲的起 始时间:早侏罗世

古太平洋板块在欧亚大陆下俯冲的起始时间一 直存在争论,包括二叠纪(Li and Li, 2007;Sun et al., 2015;Yang et al., 2015)、三叠纪(Zhou et al., 2014)、早侏罗世(Wu et al., 2007a;Xu et al., 2009,2013)和白垩纪(Chen et al., 2008)不同观 点.由于二叠纪期间在我国东北仍然存在古亚洲洋 板块的俯冲作用(如华北克拉通北缘东段;Guo et al., 2016;Wang et al., 2018),如何鉴别二叠纪期 间古大洋板块的性质,它是古亚洲洋还是古太平洋 目前并没有得到解决.因此,本文在讨论古太平洋板 块在欧亚大陆下俯冲起始时间时,只涉及中生代.

东北亚陆缘三叠纪火成岩主要分布在华北克拉 通北缘和苏鲁一朝鲜京畿造山带的北西侧,前者主 要由碱性火成岩和双峰式火成岩构成,并呈近东西 向带状展布,反映其形成应与古亚洲洋最终闭合后

的伸展环境相对应(Xu et al., 2013; Tang et al., 2018),后者主要由碱性岩(如石岛岩体; Yang et al., 2007)和双峰式火成岩(如蚂蚁河岩体;裴福萍 等,2008)构成,并呈平行于苏鲁造山带的北东一南 西向带状展布,其形成应与苏鲁造山带快速折返过 程相对应.此外,在我国吉黑东部和俄罗斯远东地 区,晚三叠世地层是一套被动陆缘沉积建造组合 (Zhang et al., 2014),它们与晚三叠世的 A 型流纹 岩(Xu et al., 2009)和双峰式侵入岩(Wang et al., 2015a)一起共同揭示了三叠纪期间东北亚陆缘是一 个被动陆缘的构造属性.相反,进入到早侏罗世,佳 木斯地块东缘早侏罗世钙碱性火山岩的发现(Wang et al., 2017b)以及早侏罗世火成岩自陆缘向陆内火 成岩成分的极性变化(Yu et al., 2012;Guo et al., 2016; Wang et al., 2017a), 很好地揭示了古太平洋板 块俯冲作用的开始(图 11a),这也得到了东北亚陆缘 早侏罗世增生杂岩(如日本的美浓地体、我国黑龙江 杂岩)的支持(Xu et al., 2013; Tang et al., 2018).基 于韩国南部大量早一中侏罗世火成岩的存在和我国 黑龙江杂岩中最新金红石 U-Pb 和云母 Ar-Ar 定年 结果(172~175 Ma; Dong et al., 2017; Aouizerat et al., 2018),可以判定古太平洋板块在欧亚大陆下的 俯冲作用可能持续到中侏罗世(Tang et al., 2018).

## 7.2 晚侏罗世一早白垩世早期:东北亚陆缘走滑的 构造属性与古太平洋板块斜向俯冲

晚侏罗世一早白垩世早期东北亚陆缘的构造属 性一直存在争论.传统上认为古太平洋板块俯冲仍 然控制了东北亚陆缘的构造演化,该区仍处于活动 陆缘的构造属性(Maruyama et al., 1997; Seton et al., 2012; Zhu et al., 2017). 然而, 近年来对东北亚 中生代岩浆作用的研究结果表明,除少数增生地体 (如饶河杂岩)外,中国东北的东部(松辽盆地以东地 区)、俄罗斯远东、日本和韩国普遍缺失 160~ 135 Ma的岩浆作用(Xu et al., 2013),这似乎与板 块俯冲的活动陆缘背景相矛盾.此外,从陆缘增生地 体的生物学证据和碎屑锆石物源分析结果均显示, 目前位于东北亚陆缘的这些增生地体其原始位置是 位于低纬度地区(邵济安和唐克东,1995; Zhou et al., 2015),日本美浓地体到达海沟的时间约是 190 Ma, 地体最终构造就位时间约是 175 Ma (Isozaki, 1997),这与早一中侏罗世古太平洋板块 的俯冲作用相吻合.那么,这些陆缘增生杂岩是何时 从低纬度走滑到高纬度地区的?由于后期构造的叠 加改造,目前很难鉴别地体从低纬度到高纬度的走

滑构造,然而分布在陆缘的郯庐断裂带是我国东部 重要的走滑断裂带,近年来,对该断裂带北部的研究 表明,约160 Ma 是该断裂带重要的走滑时期(孙晓 猛等,2016;Zhu et al., 2018),除此之外,断裂带南 部和北部还存在约 139 Ma 的走滑事件(Zhang et al., 2018).上述走滑事件的定年结果与日本中生代 地体所发生的变质作用时间(160 Ma 和 140 Ma; Isozaki, 1997)是相吻合的.结合饶河增生杂岩的最 终构造就位时间(137~130 Ma; Zhou et al., 2014),可以判定东北亚陆缘地体从低纬度到高纬度 的走滑事件发生在 160~140 Ma.这与东北亚陆缘 普遍缺乏晚侏罗世一早白垩世早期岩浆事件相一 致.综上所述,可以得出在晚侏罗世-早白垩世期 间,东北亚陆缘与古太平洋板块之间处于一种走滑 的构造属性,并与古太平洋板块以小角度斜向俯冲 作用相联系(图 11c).

## 7.3 早白垩世晚期一古近纪:古太平洋板块俯冲一 后撤过程

进入到早白垩世晚期(约130 Ma),东亚陆缘早 白垩世晚期(130~110 Ma)岩浆作用广泛分布,在 陆缘区形成一套钙碱性火山岩组合(如华北克拉通 东北部分布的二股砬子组和果松组以及佳木斯地块 东部产出的皮克山组),陆内为双峰式火成岩组合 (如松辽盆地中的营城子组和大兴安岭分布的上库 力组和伊列克得组),火成岩组合的空间变异揭示了 古太平洋板块在欧亚大陆下俯冲作用的发生(图 11d) (Xu et al., 2013; Tang et al., 2018), 这也得 到了俄罗斯远东哈巴杂岩和黑龙江省东部饶河增生 杂岩早白垩世构造就位的支持(Zyabrev and Matsuoka, 1999; Zhou et al., 2014; Wang et al., 2017b). 与早白垩世晚期相比, 早白垩世最晚期 (110~100 Ma)-晚白垩世岩浆作用的空间分布范 围逐渐向东收缩,此时陆缘主要由一套钙碱性火成 岩组合构成(如延边地区分布的部分屯田营组火山 岩;Xu et al., 2013),而陆内主要为一套碱性玄武 岩(如阜新碱锅玄武岩、辽南曲家屯玄武岩和胶东大 西庄玄武岩;Xu et al., 2013).进入到古近纪,东北 亚陆缘岩浆作用范围进一步向东收缩,该期岩浆作 用主要分布在中国东北的最东部(如三江盆地和珲 春地区;王智慧等,2016)、俄罗斯远东的东锡霍特一 阿林和萨哈林岛(44~40 Ma; Liao et al., 2018).从 早白垩世晚期到晚白垩世一古近纪,东北亚陆缘岩 浆作用的空间范围逐渐向东收缩(图 12;孙明道, 2016),这揭示了大洋板片俯冲角度逐渐变陡,即板



图 12 中国东北及俄罗斯滨海边疆区白垩纪一古近纪岩浆活动时空变异

Fig.12 Spatial-temporal variations of the Cretaceous-Paleocene magmatisms in NE China and Russian Far East 据孙明道(2016)

片逐渐后撤(roll-back)的过程(图 11e).

## 7.4 日本海打开标志着从活动陆缘到沟一弧一盆 体系的开始

虽然目前对日本海形成的时间还存在争论(最 大的时间跨度达到侏罗纪-新第三纪;Fukuma et al., 1998;Baba et al., 2007),但多数学者认为日本 海的扩张开始于 20 Ma 之前(Lallemand and Jolivet, 1986;Tamaki, 1995).海沟后撤已成为日本海形成的 主要构造模式(Seno and Maruyama, 1984).日本海的 打开,标志着东北亚陆缘已经从早白垩世-古近纪的 活动大陆边缘环境转变成沟-弧-盆体系.同时,俯 冲带的快速后撤,导致俯冲板片在地幔过渡带的滞 留,进而标志着东亚大地幔楔的形成(图 11f).

## 8 结论

(1)兴蒙造山带主要由微陆块和其间的造山带 组成,微陆块中存在前寒武纪结晶基底,并具有壳幔 耦合特征.这些前寒武纪地质体已经被后期构造一 岩浆热事件多期改造与再造.

(2) 兴蒙造山带微陆块中的陆壳增生主要发生 在中元古代和新元古代以及次要的新太古代和古生 代,且以垂向增生为主;而陆块间的造山带或岛弧地 体陆壳增生发生在新元古代和古生代,且以横向 增生为主.

(3)兴蒙造山带中微陆块间的拼合主要发生在 古生代,古亚洲洋的最终闭合时间为中三叠世,且为

### 剪刀式闭合过程.

(4)蒙古-鄂霍茨克构造体系对我国境内的影 响至少始于晚二叠世,南向俯冲作用发生于晚二叠 世-早侏罗世,大洋闭合于中侏罗世,晚侏罗世-白 垩纪主要表现为造山后的伸展环境.该构造体系影 响的空间范围主要在松辽盆地以西和华北克拉通 北缘地区.

(5)古太平洋板块中生代期间在欧亚大陆下俯 冲起始的时间为早一中侏罗世,晚侏罗世一早白垩 世早期东北亚陆缘处于走滑的构造属性,早白垩世 晚期一古近纪岩浆作用的向东收缩揭示俯冲板片的 后撤过程,日本海的打开标志着东北亚陆缘已经从 活动陆缘转换成沟一弧一盆体系,并且标志着东亚 大地幔楔的形成.

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